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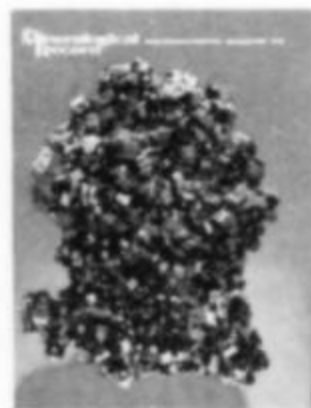
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COVER: VANADINITE group, 17 cm tall, from
Mibladen, Morocco. Collection of Stephen Stolte
and Achim Karl (*Mineralien Fossilien Galerie*,
Frankfurt); photo by Harold and Erica Van Pelt.

Dangers to Science from Species Dealers

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With the remarkable increase in the description of new mineral species in recent years has come a similar increase in the number of people who provide such material. Species dealers may do harm in two principal ways: the dissemination of incorrect information, and the physical damage to specimens, our topic here. Unfortunately, the few benefits to mineralogy from such commercial activity are compromised by the fragmentation and practically random distribution of these scientifically significant specimens. This is a concern widely shared in the professional curatorial community.

There has always been some partitioning and distribution of rare species; it benefits the science when such materials are deposited in large institutional research museums in various countries, thus preserving the rarest minerals in several regions of the world. However, the present wanton fragmentation by today's commercial dealers and collectors is a genuine disservice to the science of mineralogy. Although this is not a new activity, the scope of it has become staggering! The usual scenario is that dealers/collectors obtain a specimen of a rare mineral, and then break it into numerous small fragments, sometimes microscopic, which are either sold or exchanged. Despite the vigorous and indignant denials of some species dealers and collectors, this is a very prevalent practice, as those who collect or study rare species know too well. Such distribution is based mainly on a desire for profit, a goal unrelated to the advancement of mineralogy. Far too often fine samples are broken up into meaningless chips, some so small that they are packed in gelatin capsules! There seems to be no lower limit for such fragmentation; it seems to be a "seller's market" with no other controls. This is a sad situation, which is a result of irresponsible behavior. Not all species dealers are guilty, but those few with fine reputations are already well-known and need no exoneration here.

The subdividing of samples into tiny chips is harmful because it destroys associations, and thus much useful information about assemblages and parageneses which contain clues as to the conditions under which the mineral formed. It is crucial to remember that a new mineral description is but the *first* investigation of this naturally occurring material. Almost inevitably, other diverse studies will follow in subsequent years. The original description, done by descriptive mineralogists, may someday be followed by other investigations concerning the relations of the chemical system in which the mineral formed, the relative stability of the associated minerals, crystal structure analysis, thermochemical investigations, measurement of physical constants such as elasticity, compressibility, etc. For many new minerals, all of these factors are unknown, and when the specimens are irresponsibly broken down, the perpetrators might well be ensuring that such studies can never

be done. For these studies, there must be complete assemblages, and substantial quantity, not merely tiny chips, broken from others to make many samples and more profit. Even if the fragments could be gathered together, any apparent information derived from such jigsaw puzzle activity would not meet the criteria of scientific inquiry. Hence, the fragmentation actions of species dealers will impede scientific investigations, and in some cases preclude them.

What is the answer to the problem? There might not be one; appeals for responsible behavior fall on ears slightly deafened by the ringing of cash registers. There are no laws, no rules, and there is little in the way of peer pressure of a positive sort in the mineral species marketplace. One could make pleas to different constituencies, for example:

1. A plea to species collectors to dampen the demand for such itty-bitty meaningless chips of doubtful integrity. In fact, were species collectors to inquire as to the true significance of such samples, they might not acquire them at all. However, the quest for profit coupled with other, more benign motivations of collectors, such as the "urge for completeness," might render such a plea ineffective.

2. A plea to mineral scientists not to provide samples to those dealers and collectors well-known for fragmentation activity. This is, unfortunately, not an action which benefits the hobby of mineral collecting. Given the low probability of success of other actions, however, this type of request might be the most effective.

One point should be seriously considered by the collector, and that is the relative worthlessness (scientifically and economically) of such tiny fragments. Without associations, and recognizable features, and without enough material for testing, verification, and careful curation of the remainder, such tiny fragments have little significance, and thus little real value, especially in the long term. For those interested in the investment aspects of mineral collecting, such concerns should be significant ones. Perhaps never before in the history of mineral collecting have so many held so much on so much faith, and little or no proof of even correct identity. These matters deserve careful consideration by the prudent collector.

Concerned collectors, curators and scientists can discourage the destructive fragmentation of rare minerals by employing the power of the mineral marketplace. Patronize only those dealers who avoid commerce in fragmented rare species. Refuse to do business with dealers who impede the advancement of science by such activity, and couple this with articulate, vocal and persistent criticism. Market forces created this problem, and market forces can solve it!

notes from the EDITOR

THE RECORD COPYRIGHT

The *Mineralogical Record*, as most people know, is a copyrighted publication. This means that it is illegal for unauthorized copies to be made of anything which appears in the magazine. Actually the situation isn't quite that simple, and we receive inquiries regularly to clarify various points of procedure. Here are some useful facts to know:

• Advertisements

An advertiser may retain the rights to his own ad, particularly if he provided it as camera-ready on his own. If we have designed his ad for him at no charge, as we often do, he may still use the ad anywhere else he wishes but may not have exclusive rights to particular illustrations we have provided.

• Illustrations (drawings)

Many drawings seen in the magazine are taken from old publications having expired copyrights. These are in the public domain, and anyone may use them. However, some drawings are recently created and may *not* be used by the public. Generally speaking, unless the art has been purchased outright by someone else, the artist retains ownership of the reproduction rights. I have seen artwork stolen from the pages of the *Mineralogical Record* and used on fliers, ads in other magazines, show posters, and even personal stationery. This exposes the user to legal action, and "ignorance of the law is no excuse."

• Illustrations (photos)

The *Mineralogical Record* does not pay for the use of photos, but even if we did we would have only "one-time" rights. Photographers in all cases retain the copyright on their photos.

• Article and Column Text

All text published in the magazine (except material reprinted from elsewhere) is thereafter owned exclusively by the *Mineralogical Record*. However, because we do not pay for articles, we are usually cooperative with authors who ask for permission to have their writings reprinted by another publisher.

Upon occasion some of our readers may wish to make photocopies of certain articles and columns. If they wish to make only a single copy, for private use, we hereby grant permission (no written permission required). If they wish to make multiple copies for distribution (e.g., in classrooms or for club bulletins) written permission is required and generally granted, provided that the copies will be given away free or sold at cost or less. All permissions granted in writing must be for *specific* items . . . blanket permission is almost never given.

Extracts for republication require no permission provided the extracts are short. It is common practice among professional journals to allow articles to be quoted or line figures (not photos) used in subsequent articles without written permission, but the source must be stated. If the extracts are long, or many illustrations are to be reproduced, written permission should be obtained.

Articles and columns to be reprinted should *not* be re-edited, reorganized, translated or abridged without the author's approval of the new version.

• Our Attitude on Reprinting

The purpose of the *Mineralogical Record* is to distribute information in a helpful, useful way, so we are not, in principle, opposed to reprinting. The additional circulation is good publicity which may help to introduce us to new readers. And we are always interested in helping our authors and photographers achieve wider exposure. Republication for commercial purposes is also negotiable but may involve royalties to authors and/or fees for the use of photos and art (payable to the photographers and artists). It is conceivable that under some circumstances republication would be detrimental to the image of the *Mineralogical Record*, or would be against our best interests in some other way. Permission might then be refused, but that is a relatively remote possibility. Our main wish is to facilitate the spread of information; to that end we strive to cooperate fully on matters of copyright.

(My thanks to Janet Cares for suggesting this note.)

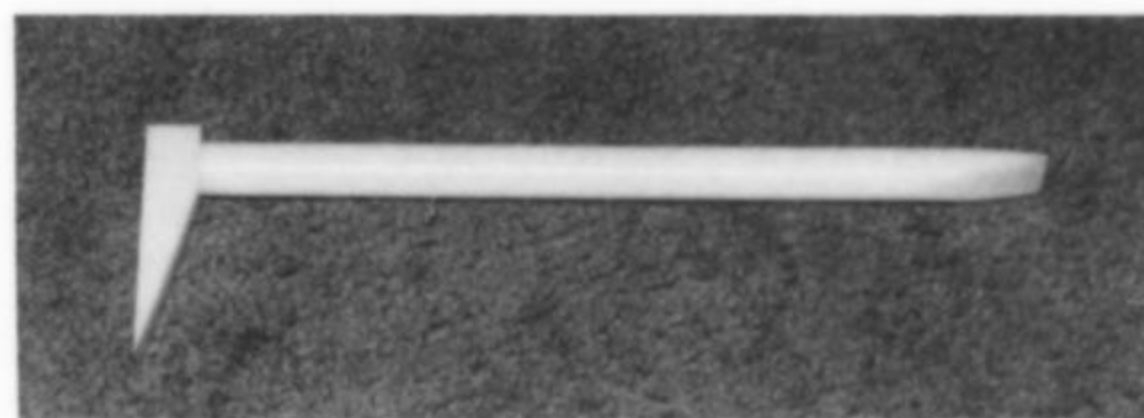
PHOTO COMPETITION

The honor of being the photographer for the 1987 Tucson Show Poster is up for grabs, with a \$250 award thrown in as a sweetener! The only catch is that entries for consideration must be *large format*, 4x5-inch color transparencies (not 35 mm). Furthermore, the mineral pictured, a single specimen or a group, must be *quartz* (any variety or varieties, any size).

Specimens pictured should, of course, be of high quality since the photo will be representing the show. Ownership of specimens must be established and owner permission granted for publication. The winner will be notified by October 20. It is customary for the photographer to be available at the show to autograph posters.

Entries should be mailed to Susan Angelon (2206 E. Jasmine Drive, Tucson, AZ 85706) so as to arrive no later than October 1, 1986. All entries will be returned; decision of the judges is final; right to choose none of the entries is reserved.

Incidentally, it will *not* be acceptable to have a 35 mm slide blown up to a 4x5 transparency; the photo must be shot originally on 4x5 film, using a large format camera. Color bars and a gray scale would be nice, but are not required.



NON-SCRATCHING TOOL

Larry Venezia (115 Coleridge St., East Boston, MA 02128) has come up with a collecting tool strong enough to pry apart rocks and soft enough to avoid scratching crystals. It's an 18-inch prybar made of solid nylon, which he has been field testing for two years. Though generally thought of as being rather flexible, nylon is quite rigid and hard when sufficiently thick. It can be hammered into crevices and vigorously levered with virtually no prospect of breaking or even bending much. It's also a lot lighter than tool steel . . . an important consideration in remote areas where specimens must be backpacked out. The price: \$20 postpaid (wholesale lots available).

MINERALS IN CHINA

What will surely go down in history as the rarest important mineral book of modern times is the English edition of *Minerals in China*. This is a beautiful, full-color, hardcover book of 165 pages,

published by Shanghai Scientific and Technical Publishers in an edition limited (they say) to *only 200 copies*.

The original edition, of course, was in Chinese, but some intelligent and kindly soul thought it might be nice to run off a few copies in English, too. The Geology Museum in Peking traded many copies to their foreign correspondents for library materials, and sold the remainder to visitors . . . and to the Mineralogical Record. The edition is now sold out in China, but we have in stock 40 copies of the English edition and 7 copies of the *Chinese* edition. The price is \$67 each, postpaid (foreign orders add \$1), while they last. Some copies have minor mildew along the edge and are sold as is. Order from Mineralogical Record Book Dept., P.O. Box 1656, Carson City, NV 89702.

GONIOMETER PINS

Our bronze lapel pin depicting the Mineralogical Record's goniometer (see Vol. 16, no. 3, p. 169) sold briskly. So briskly, in fact, that we ran out of them in six months. We have recently placed another order with Josten's, however, and expect to have them back in stock in July. Readers can obtain one by sending \$12 plus \$2 postage and packaging (\$3 foreign) to the Mineralogical Record Book Department, P.O. Box 1656, Carson City, NV 89702.

GREEK NEWSLETTER

A bimonthly, bilingual newsletter is being initiated in Greece. The title is *Oryktologika Nea - News on Minerals*. Apparently this will be the first newsletter of its kind in Greece. It will be published in English and Greek, and will deal with Greek mineral localities and collectors. Collectors wishing to advertise for trading partners in *Oryktologika* may do so for \$1 per insertion. A one-year airmail subscription is \$25. Send to: D. G. Minatidis, Editor, 70 Queen Sophia Avenue, Piraeus 185-32, Greece.

GOLDSCHMIDT'S ATLAS

Most collectors know about Victor Goldschmidt's monumental collection of crystal drawings, *Atlas der Krystallformen* (1913-1923). The work consists of nine volumes of crystal drawings and nine volumes of supporting text on localities, references, crystal forms, etc. The bound set in the Mineralogical Record Library occupies 61 cm (2 feet) of shelf space and is probably worth around \$2500 on the rare book market. Obviously, few collectors can afford the luxury of owning an original set.

Finally a high-quality reprint of this invaluable work is being published by the Rochester Academy of Science. The first two volumes (text and figures combined, softbound) are now available from the Mineralogical Record Book Dept. (P.O. Box 1656, Carson City, NV 89702) at \$70 for the two plus \$4 postage (\$1 more for foreign orders). We have in hand 60 copies of each volume, from a total softcover press run of 300 copies.

The RAS is publishing these by the bootstrap method, that is, sales of each volume must be more or less complete in order to finance subsequent publication of the next volume in the set. Purchasers of the first two volumes should be prepared to purchase each of the seven succeeding volumes as it becomes available, at the same price of \$35 per volume. That will mean a total price of \$315 for the nine combined volumes of text and figures. Consider it as a sort of subscription!

SAM WELLER CUSTOMERS

What could be worse for a mail-order dealer than to have all his customer records go up in smoke? That is unfortunately what happened on February 16 when the premises of *Sam Weller Minerals*

were destroyed by fire. Rose and Sam will continue the business from a small house next door, and are even now rapidly replacing lost inventory, but they need to hear from previous customers so they can rebuild their mailing list. Their new listing of specimens should be ready to send out right about now. Write to: Sam Weller, "Chy-an-Stennack," Lower Boscaswell, Pendeen, West Cornwall, England.

CALL FOR PAPERS

The eighth Friends of Mineralogy-Mineralogical Society of America-Tucson Gem and Mineral Society symposium will be held on February 15, 1987, at the Tucson Community Center. The symposium topic will be quartz, the featured mineral for the Tucson Gem and Mineral Show, and papers covering any aspect of quartz mineralogy including structure, physical and chemical properties, paragenesis and localities are invited.

Presentations will be limited to 20 minutes with an additional 5 minutes for questions. An audience of knowledgeable amateurs as well as professional mineralogists is expected so the emphasis of papers should be topical rather than specialized.

Papers by students will be considered for the "Best Student Paper" award from Friends of Mineralogy. The award is intended to cover part or all of the travel expenses of the student attending the symposium.

Abstracts should be submitted to FM-MSA-TGMS 8th Symposium, Attn: Henry Truebe, St. Joe American Corp., 2002 North Forbes Blvd., Tucson, AZ 85745. Deadline is September 31, 1986.

CHECK YOUR GARAGE

If you have old copies of *Rock & Gem* magazine, you can trade them in for a credit toward your next subscription renewal for the *Mineralogical Record*. We're trying to complete our set, and will give a **\$2 credit per copy** for issues we still need. (We'll also accept them as outright donations, of course.)

Check the following list of what we need and send a letter listing what you have available. I'll write back telling you which ones to send (this will avoid duplicates). Send to Editor, Mineralogical Record, 4631 Paseo Tubutama, Tucson, AZ 85715.

- 1971 All issues
- 1972 Jan., Feb., Mar., Apr., May, Aug., Sept., Oct.
- 1973 Jan., Mar., May, July, Aug., Sept., Oct.
- 1974 Jan., Mar., Aug., Sept., Oct., Nov., Dec.
- 1975 Jan., Feb., Apr., Aug., Sept., Nov.
- 1976 Jan., Mar., Apr., May, Oct., Dec.
- 1977 All issues
- 1978 Mar., Apr., May, Nov.
- 1979 Feb., Apr.
- 1981 July, Aug.
- 1982 May, July, Aug., Oct.
- 1984 June, Aug.

Incidentally, it turns out our set of *Gems & Minerals* isn't complete after all. The bindery found several seriously damaged copies needing replacement. So we make the same offer for the following copies of *Gems & Minerals*:

- 1964 Feb., Apr., May, Sept., Oct., Dec.
- 1965 May, Sept., Oct.
- 1969 Feb., Mar.
- 1985 Jan., Apr.

MINERAL MUSEUMS OF EUROPE BOOK

Uli Burchard, author of the recently published *Mineral Museums of Europe*, will be available to autograph copies at the *Mineralogical Record's* table at the Denver Show, Sept. 12-14. ☒

Great Pockets:

the National Belle mine

Arthur E. Smith, Jr.
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Houston, Texas 77036

Tom Rosemeyer
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Ouray, Colorado 81427

T*he National Belle mine became famous in 1883 when the first of several ore-lined caverns was discovered. A number of species including enargite, sphalerite, barite, quartz, pyrargyrite and tetrahedrite can still be collected on the dumps.*

INTRODUCTION

South from Ouray, Colorado, U.S. Route 550 hugs the Uncompahgre River Canyon and twists and climbs into the heart of the San Juan Mountains. The only respite from the twisting and climbing is a quick, straight stretch across Ironton Park before a final tortuous assault on the mountains. Here the mountain sides are spotted with abandoned mines, the faded hopes of the past. There is the Silver King, Silver Bell, Cora Belle, American Girl, Yankee Girl, Belfast, Guston, Congress, Genessee, Vandervilt, Joker and many more. Some made money for their finders, owners and investors but many did not.

Evidence of mineralization is prominent. Shades of red and yellow iron oxide paint the looming Red Mountain peaks. In the warm sun the acrid smell of disintegrating sulfides lingers on aban-

doned mine dumps. As the road climbs again it twists and passes by the Treasury Tunnel trestle of the famous Idarado mine which is now closed. At approximately 3414 meters (11,100 feet) and about 16.1 km (10 miles) from Ouray is the climax, Red Mountain Pass. South from here it is all down hill to Silverton. Backtracking about 600 meters (2000 feet) north of the pass on the Ouray side, a dirt road leads east a short distance to the National Belle mine and what was the final location of the town of Red Mountain, which had several different settings on both sides of the pass.

HISTORY

In 1874 the inaccessibility of the San Juan Mountains made the "Ho, for the San Juans" rush more of a bust than a boom. In spite



Figure 1. The National Belle mine (Lakes, 1905).

of this problem, settlement and mining slowly progressed. It was not until 1881 that the mineral deposits of the Red Mountain area were discovered. The National Belle mine was discovered in 1882 by Richard Journey. Little is known about him except that he drowned in 1904 when the *Discovery*, out of Nome, sank; he had been returning from a prospecting and mining venture in Alaska.

The town of Red Mountain, located in the gulch below the mine, was built in 2 meters of snow during the winter of 1882-1883. The National Belle Silver Mining Company was organized and extensive development was started in the summer of 1883. Initial development consisted of drifts driven into the sides of the siliceous knob which led to the discovery of "cave deposits" in the oxidized zone.

In the summer of 1883, Ernest Ingersoll, a 31-year-old explorer, naturalist and writer, traveled about southern and central Colorado in three private narrow-gauge railroad cars. With him were his wife and three friends. Their adventures were published in a book entitled *The Crest of the Continent*. A side trip was made to Ouray and, while he indicates that they went on toward Red Mountain, he gives no description of the town or district. Excitement about the area was high at the time of his trip. After a brief history of the discovery of the district he relates the following:

Upon the heels of this discovery there was a great rush of miners and speculators toward the scarlet heights, and several large settlements—principally Ironton and Red Mountain Town—sprang up on the rough and forested hillside. Claim stakes dotted the mountain as thick as the poles in a hop-field, and astonishing success attended nearly every digging. Among them all the first lode opened, the Yankee Girl, held supremacy, as is so often the case; but a few months later a neighboring property, the National Belle, leaped far to the front at a single bound.

This occurred by the accident of a workman breaking through the tunnel wall into a cavity. Hollow echoes came back from the blows of his pick, and stones thrown were heard to roll a long distance. Taking a candle, one of the men descended and found himself in an immense natural chamber, the flickering rays of the light showing him the vaulted roof far above, seamed with bright streaks of galena and interspersed with masses of soft carbonates, chlorides and pure white talc. On different sides of this remarkable chamber were small openings leading to other rooms or chambers, showing the same wonderful rich formation. Returning from this brief reconnaissance a party began a regular exploration. They crept through the narrow opening into an immense natural tunnel running above and across the route of their working drift for a hundred feet or more, in which they clambered over great boulders of pure galena, and mounds of soft gray carbonates, while the walls and roof showed themselves a solid mass of chloride and carbonate ores of silver. Returning to the starting point they passed through another narrow tunnel of solid and glittering galena for a distance of forty feet, and found indications of other large passages and chambers beyond. "It would seem," cries the local editor in his account of this romantic disclosure, "as though Nature had gathered her choice treasures from her inexhaustible storehouse, and wrought these tunnels, natural stopping places and chambers, studded with glittering crystals and bright mineral to dazzle the eyes of man in after ages, and lure him on to other treasures hidden deeper in the bowels of the earth. . . . The news of the discovery spread like wildfire, and crowds came to see the sight, and to many of them it was one never to be forgotten."

This was only the first of these surprises, for many cavities have since been divulged, great and small, in each of which



Figure 2. Illustration inspired by the National Belle mine discovery (Ingersoll, 1890).

crude wealth had been locked up since the world was made. The character of the ores, the occurrence of these cavities, and the extremely short distances beneath the turf at which rich ore is struck, have combined to cause much discussion among geologists as to the true history of the district.

The *Red Mountain Pilot*, published in Silverton on July 19, 1883, was even more exuberant but gave even less factual information on the discovery of the caves.

Wonderful! Miraculous! Astonishing!
The wonder of Aladdin's Lamp!
A million dollars in sight!
Greatest discovery in the world!
The Comstock bonanza discounted!
The National Belle mine is boss!

A Cave was opened in the National Belle on Wednesday last, 200 feet in length and 100 feet in width with other caves running in all directions and all a mass of richest ore. Reliable parties claim that there is a million dollars in sight. A responsible capitalist from the East offered, for a one fourth interest, cash down, more than the owners asked for the mine last week. This is not exaggerated, it is simply a mountain of mineral.

Although both of these accounts exaggerate the value of the ore, the caves were a spectacular sight. However, it was the ore values that concerned the miners and the caves' contents were soon mined, sacked and shipped by mule or wagon to the nearest rail line and then to the smelter.

A toll road across the San Juans from Ouray to Silverton, passing through Red Mountain, was completed in 1883 but it was not until 1888 that the railroad from Silverton reached the mine and provided cheap transportation. The mine was sold to the American Belle Mines Ltd. of London in 1891. They changed the name to the American Belle mine and continued mining operations.

When much of the town of Red Mountain was destroyed by a fire in August of 1892 the period of peak production was past. In May of 1897 production ceased and the mine closed, due to the low price of silver and the lack of high-grade ore. Lessees worked the mine for several years after that and the U.S. Bureau of Mines lists the National Belle as shipping ore for several years between 1906 and 1916.



Figure 3. Location of the National Belle mine, Red Mountain, Ouray County, San Juan Mountains, Colorado.

GEOLOGY

The National Belle breccia pipe is located within the ring fault zone on the northwest side of the Silverton Caldera. The Silverton Caldera, a collapse structure, formed during middle Tertiary eruptions of large amounts of rhyolitic and latitic ash flows. Middle Tertiary Silverton volcanics, made up of layered volcanics, intrusive stocks, dikes and breccia bodies, comprise the rocks in the mine area. The layered volcanics are andesite, rhyodacite and rhyolite. The intrusive rocks range from rhyolite to quartz latites and andesites.

The National Belle pipe is a conspicuous siliceous knob approximately 120 meters in diameter and 60 meters high. The pipe is composed of an inner core of silicified and kaolinized breccia encircled by an envelope of silicified cavernous breccia. The outermost zone of the pipe is composed of pyritized illitic rock which grades into propylitized country rock.

In addition to the drifts into the sides of the knob, the mine was

developed by a shaft which was sunk to a depth of 150 meters (490 feet). Development of the mine was from four levels. Much of the ore in the mine occurred in caves or vugs. Level 1 was in oxidized ore, level 2 was in mixed oxidized and sulfide ore, and levels 3 and 4 were in the sulfide zone. No paying ore was found below level 3. Level 4 consisted primarily of crumbly pyrite with small, isolated patches of ore minerals.

PRODUCTION

Complete production statistics for the mine were not available but estimates indicate a total production of several hundred thousand dollars to one million dollars. In 1883 the mine produced 980 tons of silver-lead ore worth \$70,000, far short of the millions ballyhooed by the newspapers. This was from the oxidized ore which consisted of cerussite, probable anglesite and unaltered cores of galena from the caves. In 1890 the total production value was \$87,000, of which \$4,500 was from gold, \$29,091 was from silver, and \$54,000 was from copper. It is evident that this production was from the deeper sulfide ores which consisted primarily of enargite, tetrahedrite and pyrite with some galena (Ransome, 1901).

THE CAVES

There are no good, complete descriptions of any of the National Belle caves. Professor Arthur Lakes recalls an earlier visit to the mine and caves while writing for *Mines and Minerals* after a visit to the National Belle mine in 1900:

In the little open valley in which the village of Red Mountain is located stands a prominent, isolated castle rock about 300 feet in height, the top part of which is composed largely of quartz in a columnar form and intersected by numerous minor veins of quartz which would appear to meet in a general center. The columnar sides of this mass are honey-combed with little caves and caverns. On the largest of the fissures intersecting the hill on the northeast side the original National Belle mine was opened in a fissure full of cavities and containing large bodies of kaolin and decomposed ore of various kinds. When I first visited the mine some years ago they were digging out the decomposed ore with pick and shovel, and mainly with a shovel. Not a shot had as yet been let off in the mine. I crawled in with the foreman through various cavern-like openings, and we had to be very careful not to shake down masses of decomposed ore and kaolin from the roof and sides, whilst it was difficult to keep our footing from the slippery and decomposed ore lying at our feet. They were at the time shoveling out brown mud and drying it on a big iron pan and shipping this rich dried mud in sacks. This was the beginning of the celebrated mine called the National Belle.

A more scientific description can be obtained from the following composite with Schwarz (1889), Ransome (1901) and Burbank (1941). T. E. Schwarz, a mining engineer, was the mine superintendent at the National Belle and nearby Yankee Girl mine in the 1880s. Ransome, a geologist with the U.S. Geological Survey, visited the National Belle mine in 1897 just after it was shut down. He also had the unpublished notes of S. F. Emmons, who visited the mine in 1886. Burbank, also a geologist with the U.S. Geological Survey, studied and wrote on the San Juans from the 1930s into the 1970s. Only the words in brackets have been added by the writers.

These knolls [knobs] present a rough mass of quartz cut up by cross fractures, and showing small vugs and cavities on the exposed cliff faces. The ore-bearing caves which ramify throughout the mass, generally come to the surface along the cliff-base where they are partially or wholly covered by [rock]



Figure 4. National Belle mine site showing prominent knob and headframe, September 1983; Tom Rosemeyer photo.

slide. In size the caves vary up to chambers of 50 feet [15.4 m] in diameter which are connected by irregular rounded passages, branching out toward the surface but diminishing and coming together in depth. In some cases the formation of the caves along fracture or cleavage planes is evident but in others all traces of such planes are quite obliterated. The cave-walls are a porous sandy quartz, the sand from the disintegration of which forms part of the cave filling.

Down to an altitude of about 10,930 feet [3363 m, roughly the base of the knoll] in the mine workings, the sulphides within the caves and those that originally impregnated the walls have been removed by leaching. On the bottoms of some caves and lining the sides of others there are remnants of the less soluble ore and gang filling. These consist of porous spongy masses of minute quartz crystals, clay minerals, alunite, barite, partly oxidized galena and remnants of enargite. Much of the material is encrusted with ore stained with limonite. The ores are mainly carbonates of lead and iron together with iron oxide, lead sulfates, and arsenates. Kaolin [actually dickite but several other clay minerals are also present] occurs in considerable quantities, and zincblende is common. The latter occurs in botryoidal masses consisting of nearly concentric fibrous layers, and is usually found detached from the cave-walls. They are as large as 12 inches [30 cm] in diameter and emit a phosphorescent light on scratching with steel. Galena also occurs but generally as the core of an oxidized mass. The ores of adjoining or connecting caves are sometimes greatly different in [ore] grade.

The unoxidized ore [is] chiefly enargite in the deeper workings of the mine. Some caves are lined with interlocking quartz crystals between which the sulphides were deposited. Some of the enargite occurred lining caves, forming beautiful radiating clusters of orthorhombic prisms several centimeters in length. These prisms are usually covered with a thin moss-like film of malachite and often with minute crystals of quartz. It [enargite] carrying at its best as much as 30 ounces of silver and 40 percent copper was associated with both copper and iron pyrites, some tetrahedrite and galena in diminishing amount down to nearly the No. 4 level.

There is no doubt that these caves supplied some excellent crystallized mineral specimens. The five papers written in the late 1800s on National Belle minerals show the impact of this mine on mineralogy. Unfortunately few good specimens have survived. However, the mine dumps still provide good micromineral specimens for the serious collector. This material is mostly from the sulfide zone. The following mineral descriptions are a compilation from the literature, combined with what has been observed on the dumps.

MINERALS

Alunite $KAl_3(SO_4)_2(OH)_6$

Alunite is reported but analysis shows it to be closer to natroalunite (Hurlburt, 1894; Eckel, 1961).

Anglesite $PbSO_4$

Lead sulfate is reported from the oxidized ores in caves by Schwarz (1889) but anglesite has not been definitely identified.

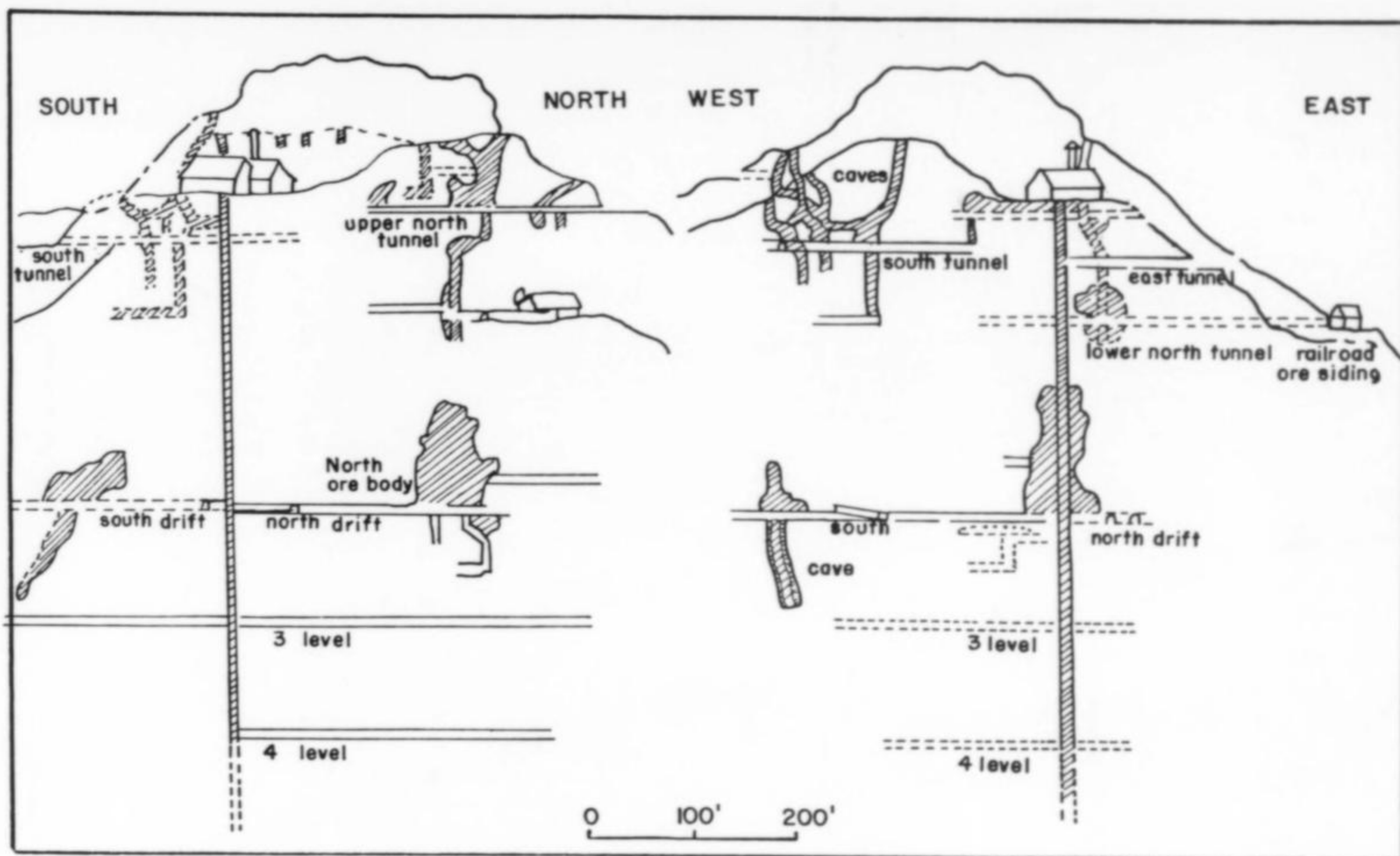


Figure 5. Sections, National Belle mine showing ore bodies and natural caverns (Ransome, 1901).

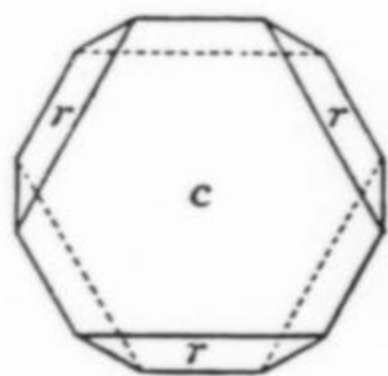


Figure 6. Alunite, National Belle mine (Hurlburt, 1894).

Barite $BaSO_4$

Barite was common in the oxidized ores. It is rare in the sulfide zone. It can be collected on the dumps as white, opaque to transparent, tabular microcrystals in vugs with quartz, enargite and pyrite crystals.

Cerussite $PbCO_3$

Cerussite is the primary ore of the oxidized zone. It occurs surrounding residual galena masses in the caves.

Chalcopyrite $CuFeS_2$

Chalcopyrite occurs with other sulfides in the deeper portions of the mine (Burbank *et al.*, 1972).

Colusite $Cu_3(As,Sn,V,Fe)S_4$

Colusite occurs with other sulfides in the unoxidized ore (Burbank, 1941). It is bronze-colored with a black streak. None has been collected in recent years.

Covellite CuS

Covellite is reported from mines in the Red Mountain district. Thin, iridescent blue coatings on some enargite may be covellite.

Diaspore $AlO(OH)$

Diaspore was reported from the National Belle mine by Ransome (1901).

Dickite $Al_2Si_2O_5(OH)_4$

Dickite was originally identified as kaolinite (Hills, 1884; Dick, 1888). In 1930 Ross and Kerr described it as a new mineral using the National Belle mine material as the basis for their description. They describe it as an incoherent glistening powder made up of six-sided plates averaging about 0.08 mm in diameter. Many of the crystals are in the form of curved or straight piles of plates that reach a maximum height of 0.3 mm. Dickite is common on the mine dumps in vugs with quartz crystals and sulfides.

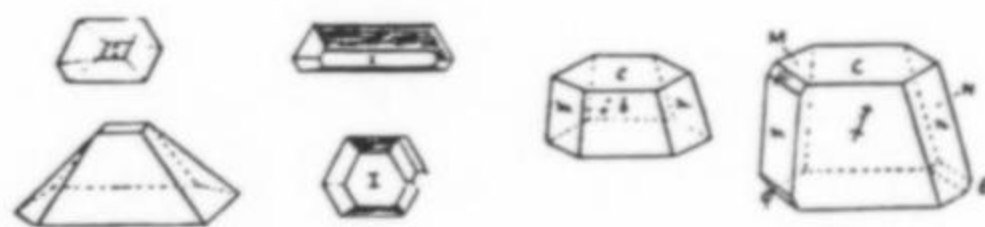


Figure 7. Dickite, National Belle mine (Dick, 1888; Hills, 1884).

Enargite Cu_3AsS_4

Enargite is the most abundant ore mineral in the sulfide zone. It occurs as masses, groups of radiating crystals, and single crystals up to several centimeters long (Ransome, 1901). Some crystals may have a coating of velvet malachite. Crystals from the dump, generally under 5 mm long, occur with quartz and pyrite crystals and may be coated with dickite. Enargite occurs in a variety of crystal forms.

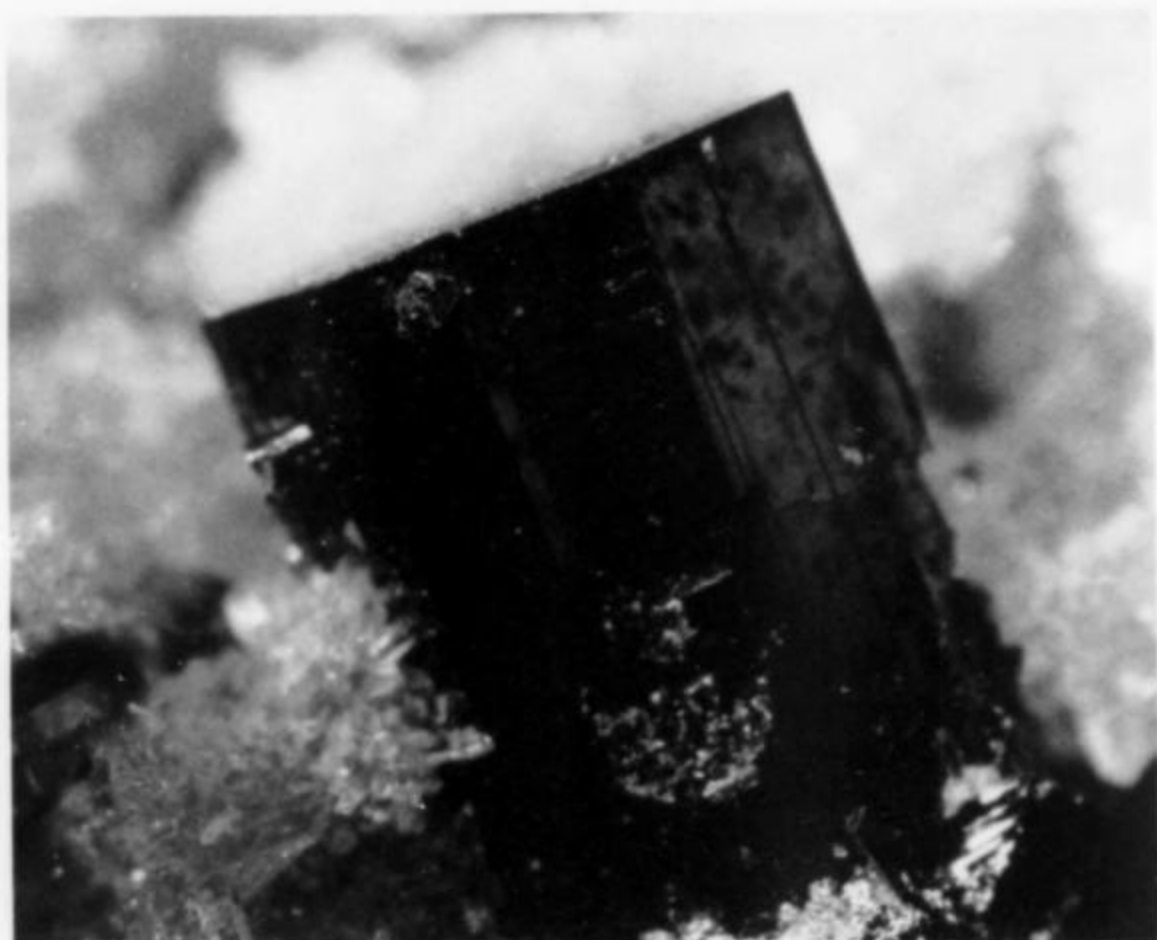


Figure 8. Enargite crystal, 2.6 mm tall, with quartz crystals; National Belle mine, 1978; Dan Behnke photo.

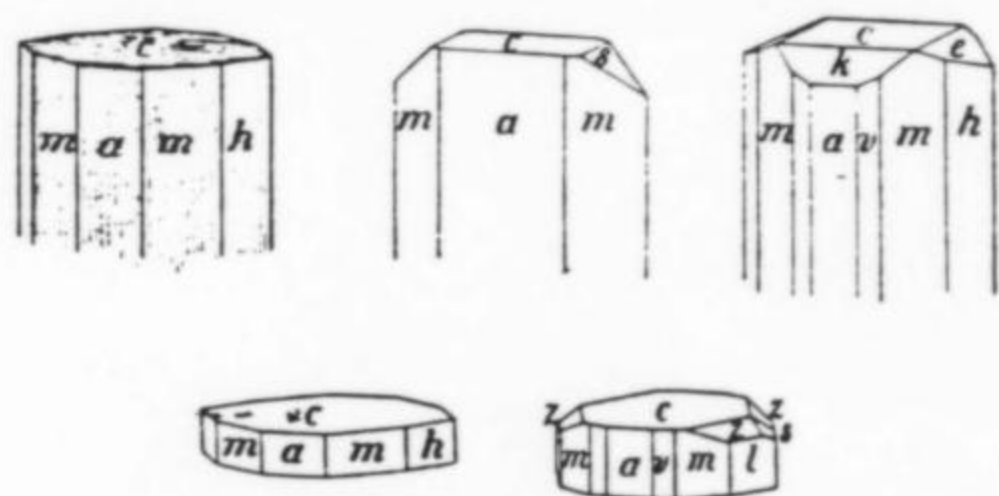


Figure 9. Enargite, National Belle mine (Pirsson, 1894).



Figure 10. Enargite crystal group. The largest crystal is 0.8 mm long; National Belle mine, 1978; Dan Behnke photo.



Figure 11. Sphalerite crystal, 1.9 mm across, brownish red; National Belle mine, 1972; Dan Behnke photo.



Figure 12. White barite crystal 4.2 mm tall occurring with enargite and unidentified white crystals along its edge; National Belle mine, 1978; Dan Behnke photo.

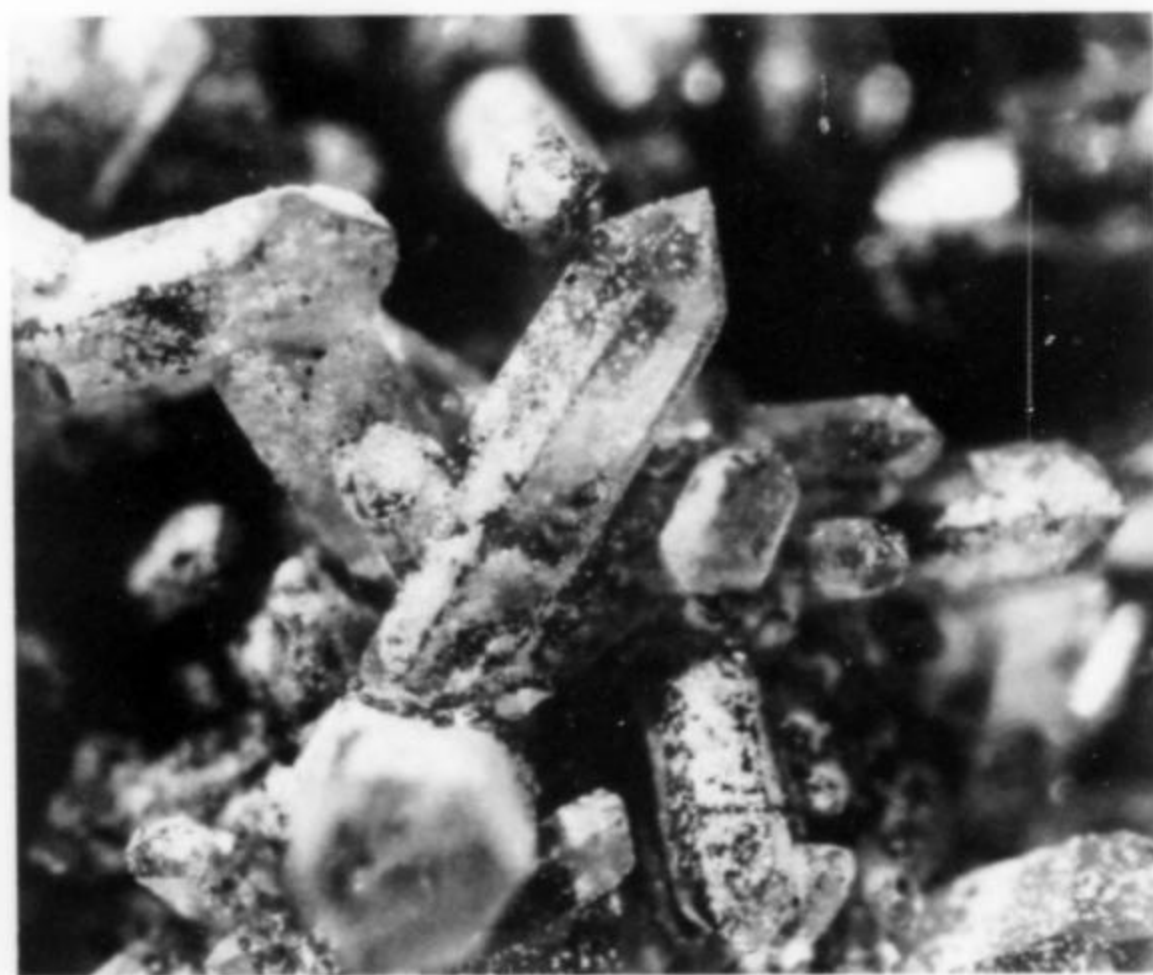


Figure 13. Dickite coating clear quartz crystals. The largest quartz crystal is 1.2 mm long. National Belle mine, 1978; Dan Behnke photo.

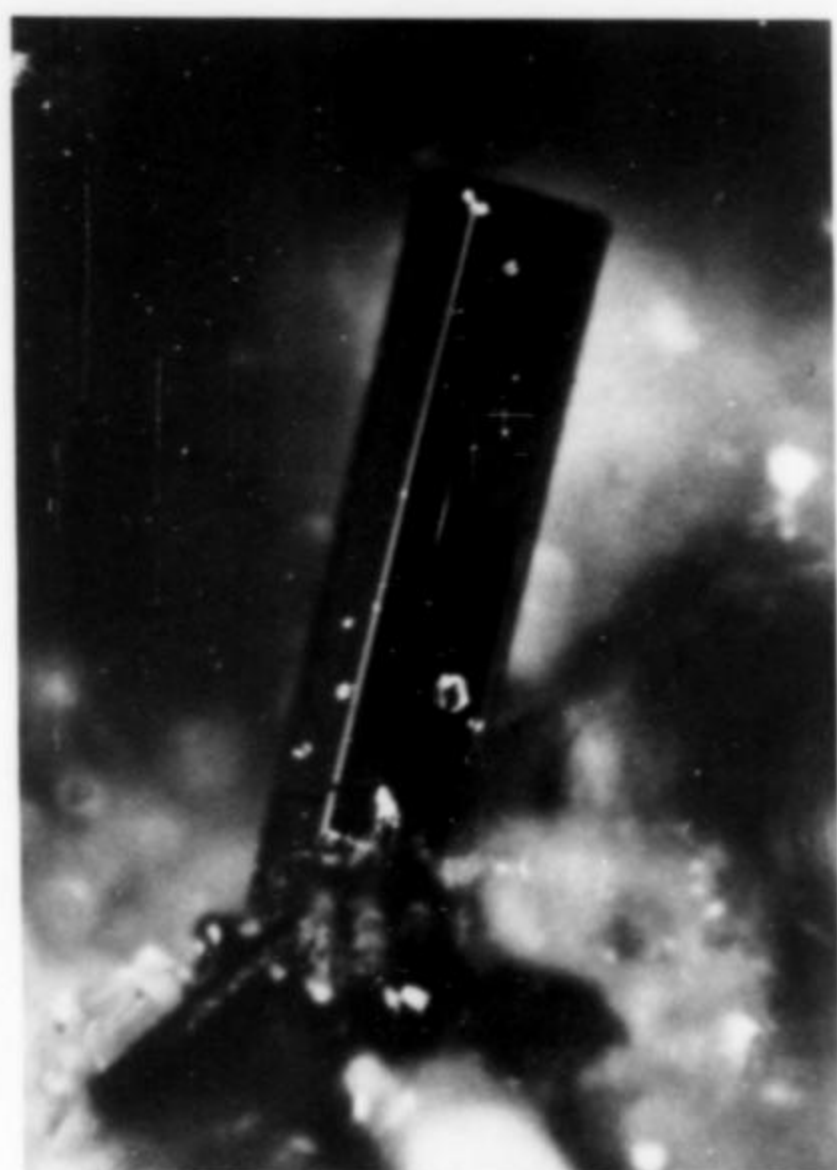


Figure 14. Pyrargyrite crystal, 0.6 mm long in vug of enargite; National Belle mine, 1978; Dan Behnke photo.

Galena PbS

In the oxidized zone galena occurs as cores of unaltered material surrounded by cerussite, probable anglesite, and spongy masses of quartz. In the sulfide zone it becomes less common with depth and is rare in the exposed dump material.

Gold Au

Free gold has been reported from the sulfide ore but it was not common.

Kaolinite $Al_2Si_2O_5(OH)_4$

The material described as kaolinite is dickite (Ross and Kerr, 1930). Studies on the clay minerals of the Longfellow mine just across the pass in San Juan County show a zonation from the center of the deposit outward. Besides dickite, Luedke and Hosterman (1971) identified high alumina montmorillonite, pyrophyllite, illite, and chlorite. A similar group of minerals and zonation may also occur at the National Belle mine.

Malachite $Cu_2(CO_3)(OH)$

Malachite occurs as stains and velvety coatings on enargite in the upper levels of the mine (Ransome, 1901).

Natroalunite $NaAl_3(SO_4)_2(OH)_6$

Natroalunite is present as aggregates of minute crystals which resemble dickite. It fills seams and pockets and occurs on quartz, pyrite and enargite. The largest crystals are 0.13 mm in diameter and 0.01 mm thick (Hurlburt, 1894).

Pitticite hydrous ferric arsenate-sulfate

Pitticite from the walls of a cave above the water level was identified by Charles Milton (Burbank *et al.*, 1972).

Pyrargyrite Ag_3SbS_3

Tiny, shiny black, elongated, prismatic crystals of pyrargyrite occur in small vugs. They are very rare in the dump material.

Pyrite FeS_2

Pyrite is abundant in the sulfide zone and on the dump where it seems to permeate much of the rock. It occurs as masses and

numerous 0.1 to 3 mm octahedral crystals. Small, modified cubic crystals occur in some vugs but are much less common.

Quartz SiO_2

Clear bright quartz crystals up to 1 cm long and 3 mm wide line vugs of the sulfide zone. They were also abundant in the caves of the oxidized zone where they were mixed with dickite and other clay minerals. Crystals are abundant in the dump material and occur with dickite, pyrite, enargite and other minerals.

Scorodite $Fe^{+3}AsO_4 \cdot 2H_2O$

Scorodite has not been reported from the National Belle mine though it might have been one of the arsenate minerals in the oxidized zone. It occurs in the Charter Oak mine across the gulch as green botryoidal encrustations on enargite and on wall rocks (Penfield, 1893).

Sphalerite $(Zn,Fe)S$

Sphalerite occurs as large botryoidal masses on the floor of the caves in the oxidized zone (Schwarz, 1890). Small reddish brown, opaque crystals and more rarely yellow, frosted, translucent crystals occur with quartz and enargite in cavities in the mine dump material.

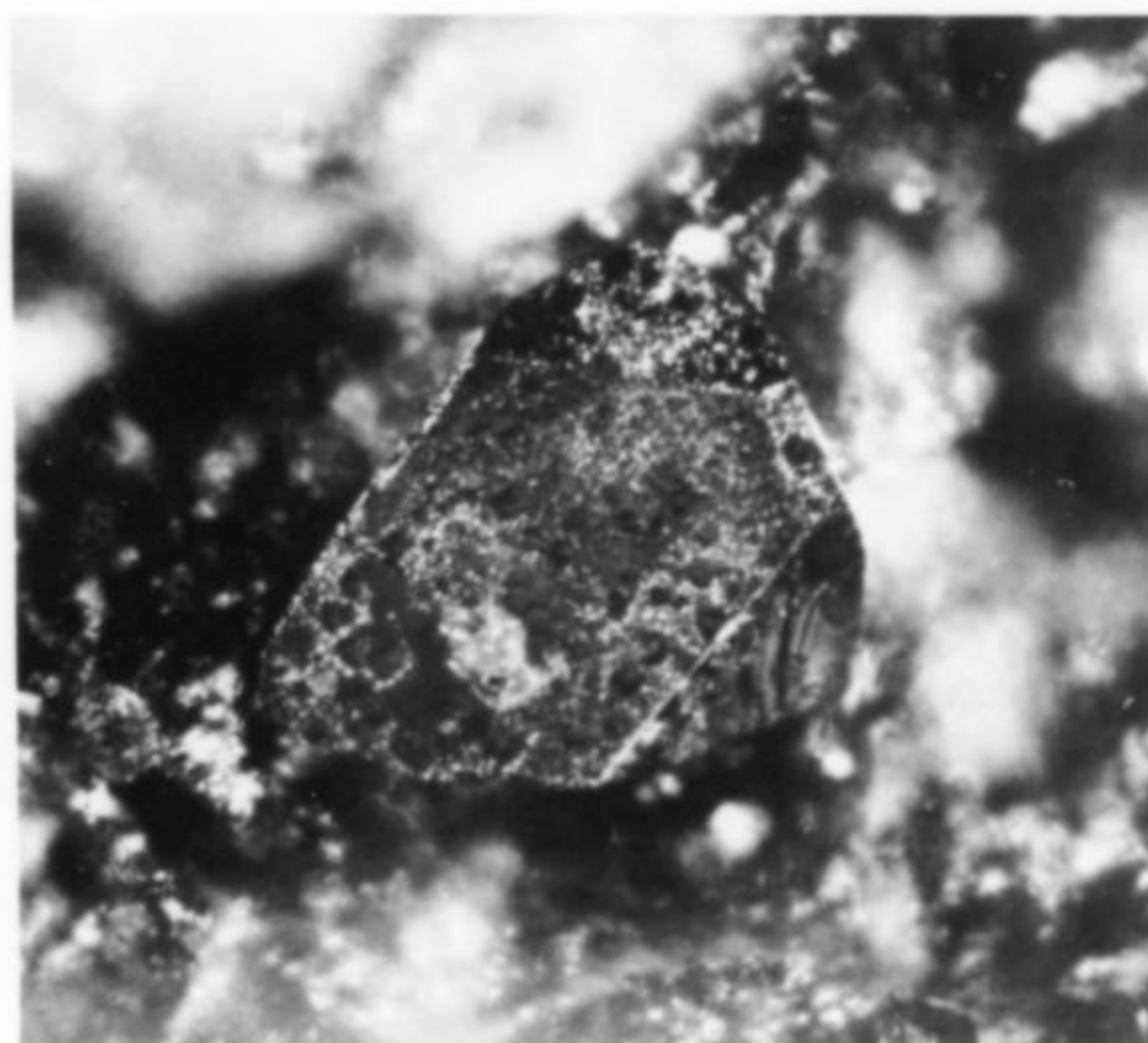


Figure 15. Tetrahedrite crystal, 0.33 mm tall; National Belle mine, 1982; Dan Behnke photo.

Tetrahedrite - Tennantite $(Cu,Fe)_{12}As_4S_{13} - (Cu,Fe)_{12}Sb_4S_{13}$

Both tetrahedrite and tennantite have been reported from the National Belle mine (Hillebrand and Kelley, 1957; Burbank *et al.*, 1972). Black tetrahedral crystals, many with truncated corners and usually under 1 mm, occur in vugs. It also occurs as masses that show numerous small faces or are composed of intergrown crystals. Enargite and quartz crystals are the most common associated minerals.

COLLECTING

The National Belle is easily accessible by a short hike or drive east from U.S. 550 on a poor dirt road. The mine area has not been posted in recent years. Examination of some of the shallow caves in the knob by a miner from Ouray in 1966 revealed nothing of specimen interest. The deeper levels of the mine are flooded but good micromineral specimens from the sulfide zone may be collected from the exposed mine dumps.

ACKNOWLEDGMENTS

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Phosphate Minerals from the Tip Top Mine, Black Hills, South Dakota

•
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First staked in the 1880s as a tin prospect, the Tip Top pegmatite has since yielded over 50 phosphate minerals including eight new species, and many specimens which are the finest known for their species.

INTRODUCTION

The Tip Top is one of a unique group of pegmatites noted for their phosphate minerals. The pegmatite has yielded approximately 80 mineral species, at least 54 of which are phosphates (Table 1). Eight new species have been identified from this locality: jahnsite, robertsite, segelerite, fransoletite, tinsleyite, tiptopite, ehrlite and a new phosphate which is presently under investigation. In addition to these new species, the authors and their colleagues have found 23 other minerals previously unreported from Tip Top (these are noted in Table 1). The total number of phosphate species from the Tip Top pegmatite is exceeded only by the Palermo No. 1 quarry in New Hampshire (64) (Segeler *et al.*, 1981) and possibly the Hagen-dorf pegmatites in Bavaria. Tip Top has also produced some of the finest crystals of pegmatite phosphates and probably holds the "best of species" title for many. These include lustrous, deep pink hureaulite crystals to 3 cm, colorless to yellowish 1-cm whitlockite rhombs, brown jahnsite crystals to 1 cm, 8-mm leucophosphate crystals and red montgomeryite crystals to 5 mm.

LOCATION

The Tip Top mine is located 200 meters southwest of the center of Section 8, R6E, T4S, 8.5 km southwest of Custer, Custer County, South Dakota. It is identified on the Fourmile quadrangle

(U.S.G.S. 7.5-minute topographic map) and on the U.S. Forest Service map of the Black Hills. The mine is located in a very scenic portion of the southern Black Hills, an area known as Pleasant Valley, which, as all visitors soon realize, lives up to its name.

HISTORY

According to Fisher (1942), the Tip Top was originally staked in the 1880s as the Tip Top lode by William Nevin of Custer, South Dakota, who filed the claim as a tin prospect; the mine is still owned by the Nevin family of Custer. In the late 1800s and very early 1900s the Tip Top was mined sporadically and the workings consisted of only a few small pits. The property was idle for several years but Nevin restaked it in 1925. He opened a small pit for development work from which a load of triphylite was mined, but no market could be found for triphylite at that time. From about 1934 to 1936 Nevin mined the pegmatite for feldspar and later leased the property to Pierce Byers, who mined it through the 1940s. In the 1950s and 1960s the Tip Top was leased and mined intermittently on a rather large scale by the International Mineral and Chemical Company for feldspar and minor amounts of beryl, montebrasite, spodumene and columbite-tantalite. The mine was inoperative during the 1970s.

In 1981 the property was leased by a couple of local miners who

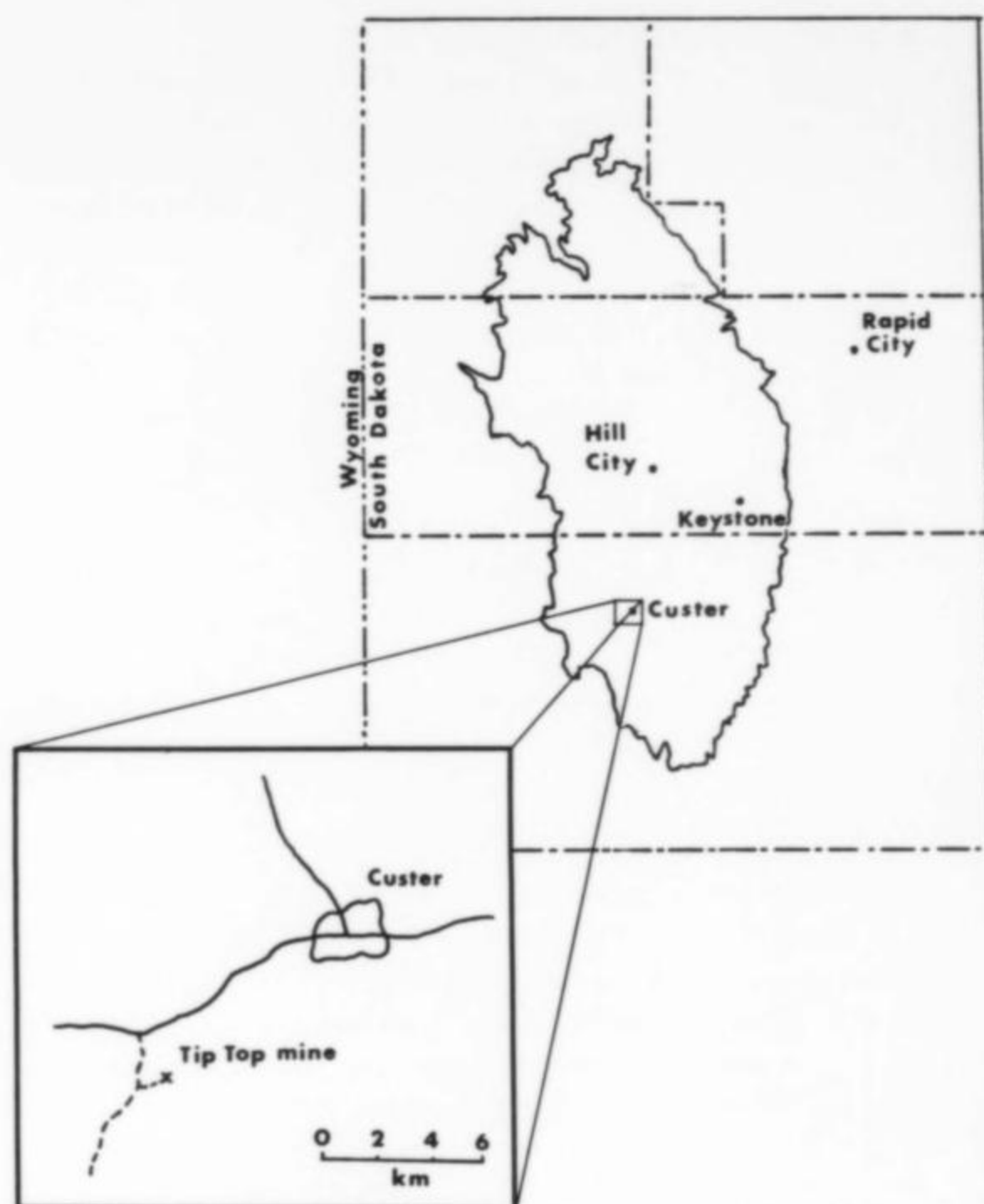


Figure 1. Location of the Tip Top mine within the Black Hills.

Figure 2. A portion of the Tip Top pegmatite showing recent working in the bottom south side. Front-end loader for scale. Photo taken by T. J. Campbell in 1982.



reopened it and started excavations in the bottom of the south side. They mined the pegmatite until late 1982 and produced beryl and feldspar along with small amounts of scrap mica and columbite-tantalite. Presently, the mine is again idle and filling with water. The pit dimensions are 60 x 120 meters and 42 meters deep.

GEOLOGY

The Tip Top pegmatite is located in the southern part of the uplifted Precambrian core of the Black Hills, which consists of low-grade to middle-grade metamorphosed eugeosynclinal sediments. The general trend of the rocks is northwest, but the detailed structure of the region has been complicated by several periods of folding (Kleinkopf and Redden, 1975; Norton, 1975). The Tip Top, like all Black Hills pegmatites, is Hudsonian (1.7 billion years) in age and, according to the classification of Ginsburg *et al.* (1979), it is a rare-element pegmatite (moderately enriched in Be, Li, Ta-Nb, Rb and Cs) of intermediate depth of formation (3.5 to 7 km).

The pegmatite itself, as described by Roberts *et al.* (1982), is a large, lenticular, zoned body which trends roughly N45°W and dips 40°SW. It is emplaced discordantly into a gray, fine-grained quartz-mica schist. The bedding and foliation strike N25°W and dip 35°SW. Locally, the foliation is parallel to the pegmatite contact. Alteration of the schist near the contact consists of a gray, fine-grained granulite. Tourmaline crystals to 5 cm can be found locally in the schist near the contact.

The pegmatite can be divided into five units consisting of four zones and one fracture filling unit. Perthite-quartz-biotite pegmatite makes up the wall zone which includes a thin border zone. The

outer-intermediate zone consists of perthite-quartz-muscovite pegmatite and has a gradational outer boundary with the wall zone. This zone contains accessory beryl and fluorapatite. The inner-intermediate zone is a perthite-quartz-triptychite pegmatite having a gradational contact with the outer-intermediate zone. Accessory minerals of importance in this zone are beryl and columbite-tantalite. This is also the zone which hosts most of the secondary phosphate minerals. The core of the pegmatite is chiefly composed of quartz with minor amounts of montebrasite and spodumene (quartz-montebrasite-spodumene pegmatite) and its outer boundary is gradational with the inner-intermediate zone. Quartz-albite-muscovite pegmatite occurs as a fracture filling unit in the wall zone.

MINERALOGY

The descriptive mineralogy of the various primary and secondary phosphate minerals found at the Tip Top pegmatite is here subdivided into primary phosphates which formed as primary phases during pegmatite consolidation, and secondary hydrothermal alteration products that formed subsequent to the crystallization of the pegmatite or during a very late stage of pegmatite consolidation.

For the sake of brevity, associations of the various phosphate minerals are given in Tables 3 and 4; note also that all of the mineral species discussed herein have been confirmed by at least X-ray diffraction methods, and frequently by other methods as well. Miller indices of the various crystallographic forms for some of the minerals were determined by visual estimation and/or visual comparison.

Table 1. Minerals of the Tip Top Pegmatite.

<i>Native Elements</i>	<i>Phosphates</i>	
bismuth	alluaudite	ludlamite
	autunite	messelite
<i>Sulfides</i>	barbosalite	mitridatite
arsenopyrite	beraunite	montebrasite
galena	bermanite	montgomeryite
marcasite	carbonate-	parascholzite
pyrite	hydroxylapatite	phosphoferrite
sphalerite	collinsite	*robertsite
	crandallite	rockbridgeite
<i>Oxides and</i>	diadochite	roscherite
<i>Oxyhydroxides</i>	dufrenite	scholzite
cassiterite	*ehrlite	*segelerite
columbite-tantalite	englishite	stewartite
goethite	eosphorite-	strengite
pyrolusite	childrenite	strunzite
todorokite	fairfieldite	switzerite
uraninite	ferrisicklerite	tavorite
	fluorapatite	*tinsleyite
<i>Silicates</i>	*fransoletite	*tiptopite
albite	frondelite-	triphyllite
almandine	rockbridgeite	vivianite
beryl	gordonite	whiteite
biotite	heterosite	whitlockite
elbaite	hureaulite	xanthoxenite
microcline-perthite	hurlbutite	*Mineral X
muscovite	hydroxylapatite	<i>Unverified</i>
quartz	hydroxyl-herderite	<i>Phosphates</i>
schorl	*jahnsite	cacoxenite
spessartine	kryzhanovskite	calcioferrite
zircon	laueite	ernstite
	leucophosphite	kidwellite
<i>Carbonates</i>	lithiophosphate	sarcopside
siderite		

Minerals previously unreported from the Tip Top mine are shown in bold type.

* Minerals for which Tip Top is the type locality.

Primary Phosphates

Primary phosphates at Tip Top are largely anhydrous and crystallized throughout the various stages of pegmatite consolidation (Campbell, 1984). Some of the minerals included in this section are actually high-temperature metasomatic alteration minerals or products of alkali leaching, but were formed during late stages of pegmatite consolidation. Granular masses exhibiting little or no crystal form (i.e., alluaudite and heterosite) are typical of metasomatic minerals. Only half a dozen species are represented in this group but they constitute the largest amount of phosphates by volume.

Alluaudite $(\text{Na,Ca})_4\text{Fe}_4^{+2}(\text{Mn,Fe}^{+2},\text{Fe}^{+3},\text{Mg})_8(\text{PO}_4)_{12}$

Alluaudite is not common at the Tip Top mine; it was first reported from this locality by Roberts and Rapp (1965). It occurs as a massive, compact, granular, dull, greenish black, subtranslucent to opaque, metasomatic alteration product of triphylite.

Ferrisicklerite $\text{Li}(\text{Fe}^{+3},\text{Mn}^{+2})\text{PO}_4$

Ferrisicklerite is the iron end-member of the ferrisicklerite-sicklerite series; it occurs only in massive form as an alteration product of triphylite and is rather abundant. It is generally found as a brown, subtranslucent to almost opaque rind several millimeters to a few centimeters thick around partially or totally altered pods of triphylite.

Fluorapatite $\text{Ca}_5(\text{PO}_4)_3\text{F}$

Fluorapatite is a ubiquitous primary phosphate mineral and can be found from the wall zone to the intermediate zone. The degree of crystallinity ranges from predominantly anhedral to occasionally euhedral. Blue-gray to green anhedral masses up to a few kilograms are not uncommon in the intermediate zone. Some euhedral blue-gray crystals to 1 cm were found in one cavity. Blue, euhedral, 2-mm to 5-mm crystals were found associated with beryl. Recent workings in the intermediate zone showed an abundance of fluorapatite partially to totally dissolved and replaced by carbonate-hydroxylapatite and minor whitlockite.

Heterosite $(\text{Fe}^{+3},\text{Mn}^{+3})\text{PO}_4$

Heterosite is the iron end-member of the heterosite-purpurite series and is moderately abundant. It is found as an outer rind on altered triphylite pods as a massive alteration product of triphylite. On fresh fracture surfaces heterosite is deep rose to reddish purple with a satiny luster and weathers to brown or brown-black with an earthy luster.

Montebrasite $(\text{Li,Na})\text{AlPO}_4(\text{OH,F})$

Montebrasite is a moderately abundant primary phosphate mineral found in the outer portion of the core zone and in the inner-intermediate zone. It occurs as large, white to grayish, subtranslucent, cleavable masses up to several kilograms. Montebrasite is commonly associated with quartz and minor amounts of columbite-tantalite and altered spodumene. No secondary alteration products of montebrasite have been discovered to date.

Triphylite $\text{Li}(\text{Fe,Mn})^{+2}\text{PO}_4$

Triphylite is the most abundant primary phosphate found at the Tip Top mine. The mineral occurs as giant subhedral crystals or nodules of which the largest recorded from this locality was 4 by 4.5 by 5 meters. This nodule was encountered in the recent working shown in Figure 2. Chemically, triphylite from the Tip Top is Fe-rich as shown by the chemical analysis in Table 2.

Most of the triphylite at the Tip Top has an outer rind of heterosite and ferrisicklerite with some alluaudite, and most shows some degree of alteration to secondary phosphates. Large masses of unaltered triphylite are rare. Triphylite is also the host for bismuth and many of the sulfides, with pyrite and sphalerite being the most abundant.

Secondary Phosphates

This group of phosphates represents the greatest diversity, scarcity and complexity of species at the Tip Top mine and includes chemical components derived from triphylite as well as from beryl, muscovite, albite, microcline-perthite and to a lesser extent sphalerite and uraninite. Triphylite supplied phosphate anions, transition elements and lithium; the silicates provided alkali metals. Cations and anions that were released subsequently produced alkali metal phosphates, and also reduced, oxidized and mixed-valence-state transition metal phosphates, some of them alkali-bearing.

Autunite $\text{Ca}(\text{UO}_2)_2(\text{PO}_4)_2 \cdot 10-12\text{H}_2\text{O}$

Meta-autunite $\text{Ca}(\text{UO}_2)_2(\text{PO}_4)_2 \cdot 2-6\text{H}_2\text{O}$

Autunite and meta-autunite are not abundant at Tip Top. The two minerals are found in close proximity to uraninite and are rarely associated with other phosphates. Autunite and meta-autunite occur as thin, tabular, bright yellow-green, vitreous, crystals up to 1 mm.

Barbosalite $\text{Fe}^{+2}\text{Fe}_2^{+3}(\text{PO}_4)_2(\text{OH})_2$

Barbosalite is not common and generally occurs as dark blue-green to almost black, very thin veinlets and powdery crusts associated with tavorite in triphylite. Very rarely it is found as lustrous, black, short prismatic crystals less than 0.25 mm

associated with hureaulite, and less commonly with jahnsite and rockbridgeite.

Beraunite $\text{Fe}^{+2}\text{Fe}^{+3}(\text{PO}_4)_4(\text{OH})_5 \cdot 4\text{H}_2\text{O}$

Beraunite is rare and is found as dark green radial sprays of translucent, vitreous to dull crystals less than 0.5 mm in size. These crystals are elongated and striated parallel to [010]. Tip Top beraunite is very similar to the material reported by Segeler *et al.* (1981) from the Palermo No. 1 quarry in New Hampshire. The oxidized variety, which is very rare, is bright orange. Beraunite has only been observed in association with leucophosphite and a jahnsite-group mineral.

Bermanite $\text{Mn}^{+2}\text{Mn}^{+3}(\text{PO}_4)_2(\text{OH})_2 \cdot 4\text{H}_2\text{O}$

Bermanite is common and was first reported from this locality by Roberts *et al.* (1974). Bermanite occurs as vitreous to sub-resinous, brown-red, tabular, pseudo-orthorhombic crystals up to 1.5 mm and as rosette-like aggregates which form druses on earlier formed secondary phosphates such as robertsite and messelite. Dominant forms are pinacoids. A partial microprobe analysis of bermanite is given in Table 2.

The Tip Top pegmatite, along with the Palermo No. 1 pegmatite in New Hampshire (Segeler *et al.*, 1981), is one of the few localities where triphylite is the primary phosphate hosting bermanite; triplite or lithiophilite is the primary phase at most other localities.

Carbonate-hydroxylapatite $\text{Ca}_5(\text{PO}_4)_3(\text{CO}_3)(\text{OH},\text{F})$

Carbonate-hydroxylapatite is abundant and occurs in a variety of forms. The most common form is very pale green to yellowish, botryoidal to stalactitic or finger-like growths coating fractures in perthite and quartz. These crusts have a radial-fibrous structure, are often banded, can be up to 5 mm thick and cover large surfaces.

This apatite species forms a base upon which other phosphates are sometimes deposited. Carbonate-hydroxylapatite is also found as colorless to white, transparent to translucent, silky, acicular, hexagonal crystals less than 1 mm in size, in cavities in quartz and triphylite.

Collinsite $\text{Ca}_2(\text{Mg},\text{Fe})(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$

Collinsite is rather uncommon and was reported from the mine by Roberts *et al.* (1974). It occurs as white to pale pinkish brown radiating bundles of silky, translucent, tabular to bladed crystals less than 4 mm in size. This is probably one of the best occurrences for fine collinsite crystals, especially when compared to the type material described by Wolfe (1940) from Francois Lake, British Columbia, where it is found as crusts with a radial fibrous structure.

Crandallite $\text{CaAl}_3(\text{PO}_4)_2(\text{OH})_5 \cdot \text{H}_2\text{O}$

Crandallite occurs as minute white balls associated with ehrleite (Grice and Robinson, 1984).

Diadochite $\text{Fe}_2^{+3}(\text{PO}_4)(\text{SO}_4)(\text{OH}) \cdot 5\text{H}_2\text{O}$

Diadochite occurs as brown to red-brown botryoidal crusts that are sometimes hollow. It is usually found encrusting earlier formed minerals such as mitridatite, rockbridgeite and jahnsite, rarely roscherite and montgomeryite.

Dufrenite $\text{Fe}^{+2}\text{Fe}_4^{+3}(\text{PO}_4)_3(\text{OH})_5 \cdot 2\text{H}_2\text{O}$

Dufrenite occurs as botryoidal, dark green, concentrically banded crusts displaying a radial fibrous structure. It can easily be confused with some forms of mitridatite and rockbridgeite. Dufrenite can be distinguished from the other two by X-ray diffraction techniques.

Table 2. Chemical analyses of some Tip Top phosphates.

Mineral Species	Li ₂ O	Na ₂ O	K ₂ O	CaO	MgO	BeO	ZnO	FeO	MnO	Fe ₂ O ₃	Mn ₂ O ₃	Al ₂ O ₃	P ₂ O ₅	H ₂ O	Total
Bermanite ³ §§	—	—	—	—	1.3	—	—	—	13.3	0.7	32.8	0.3	30.5	21.1*	100.0
Englishite ⁴	—	0.9	3.6	15.8	0.1	—	—	2.5	—	—	—	21.8	41.0	14.9**	100.6
Fransoletite ⁵	—	—	—	28.8	—	9.3	—	—	0.3	—	—	—	47.3	14.3*	100.0
Hureaulite ³	—	—	—	0.8	0.8	—	0.8	0.2	46.2	—	—	—	40.1	—	—
Hureaulite ¹	—	—	—	—	—	—	—	0.91	47.09	—	—	—	—	—	—
Messelite ¹	—	—	—	—	—	—	—	13.62	6.54	—	—	—	—	—	—
Montgomeryite ⁶	—	—	—	19.1	3.5	—	—	—	0.5	—	—	17.1	36.6	23.2*	100.0
Montgomeryite ¹	—	—	—	17.2	3.62	—	—	0.92	0.52	—	—	19.21	—	—	—
Rockbridgeite ¹ §	—	—	—	—	—	—	—	5.73	0.91	56.22	—	—	—	—	—
Rockbridgeite ¹ §	—	—	—	—	—	—	—	3.49	3.33	55.36	—	—	—	—	—
Frondelite ¹ §	—	—	—	—	—	—	—	1.37	5.19	56.87	—	—	—	—	—
Roscherite ²	—	—	—	10.3	3.52	9.97	0.95	7.03	11.72	—	—	3.66	40.40	12.62*	100.0
Switzerite ³	—	—	—	0.2	0.6	—	—	—	45.3	—	—	—	32.7	—	—
Tavorite ¹	8.47	—	—	—	—	—	—	—	—	45.58	—	0.09	—	—	—
Tinsleyite ⁷	—	—	12.4	—	—	—	—	—	—	5.2	1.1	26.6	42.2	12.5*	100.0
Tiptopite ⁸	4.5	6.1	9.9	4.3	—	15.1	—	—	0.2	—	—	0.3	55.1	3.4	98.9
Triphylite ²	9.24	—	—	0.3	1.36	—	—	32.17	10.07	—	—	—	—	—	—

* — H₂O by difference.

** — H₂O from theoretical value.

¹ — atomic absorption analysis by TJC.

² — atomic absorption analysis by Donald R. Campbell, General Tire and Rubber Co., Akron, Ohio.

³ — microprobe analysis by PJD.

⁴ — Dunn *et al.* (1984a).

⁵ — Peacor *et al.* (1983).

⁶ — Dunn *et al.* (1983).

⁷ — Dunn *et al.* (1984b).

⁸ — Grice *et al.* (1985).

§§ — Mn was proportioned between Mn⁺³ and Mn⁺² based on the known valence requirements for bermanite.

§ — Fe was proportioned between Fe⁺³ and Fe⁺² based on the known valence requirements for rockbridgeite.



Figure 3. Collinsite showing typical bladed habit. Crystals are 1.5 mm in length. Julius Weber specimen and photo.

Ehrleite $\text{Ca}_4\text{Be}_3\text{Zn}_2(\text{PO}_4)_6 \cdot 9\text{H}_2\text{O}$

Ehrleite (Grice and Robinson, 1984) occurs as colorless to white, triclinic crystals up to 2 mm, associated with parascholzite, roscherite and crandallite. Only two specimens of this mineral are known.

Englishite $\text{Na}_2\text{K}_3\text{Ca}_{10}\text{Al}_{15}(\text{PO}_4)_{21}(\text{OH})_7 \cdot 26\text{H}_2\text{O}$

The discovery of englishite at the Tip Top mine represents the first pegmatite occurrence for this mineral and the second world occurrence (Dunn *et al.*, 1983). Previously it had been known only from the variscite nodules found in a sedimentary deposit near Fairfield, Utah (Larsen and Shannon, 1930). Englishite is a relatively abundant mineral in the most recently found phosphate assem-

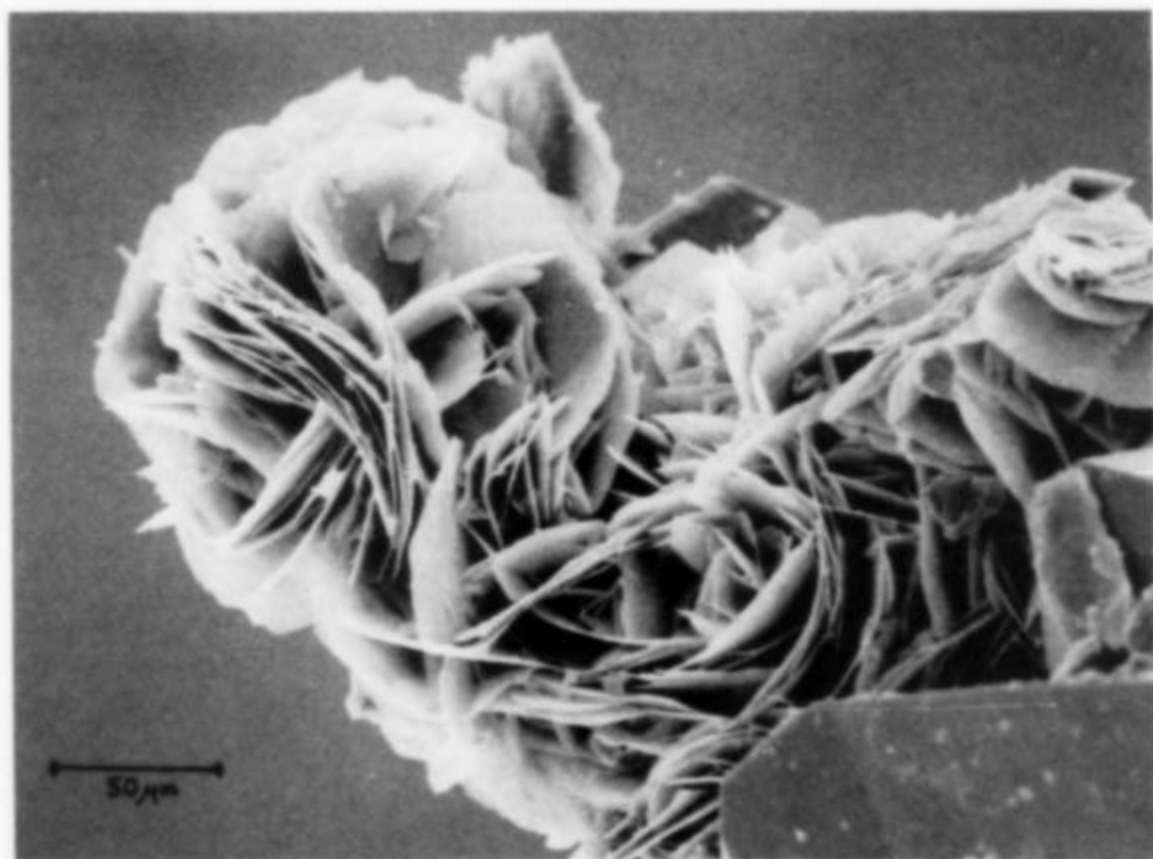


Figure 4. SEM photomicrograph of englishite depicting the typical aggregation of platy crystals. Large rosette is 0.5 mm across.

blages at Tip Top, where it occurs with other alkali-metal phosphates. It is found as pearly white spherules and hemispherules up to 1 mm that are composed of platy, micaceous crystals and as white massive seams and scaly crusts which, with roscherite, fill fractures in beryl. Englishite has been redefined chemically by Dunn *et al.* (1984b) and their analysis (Table 2) resulted in the formula given above.

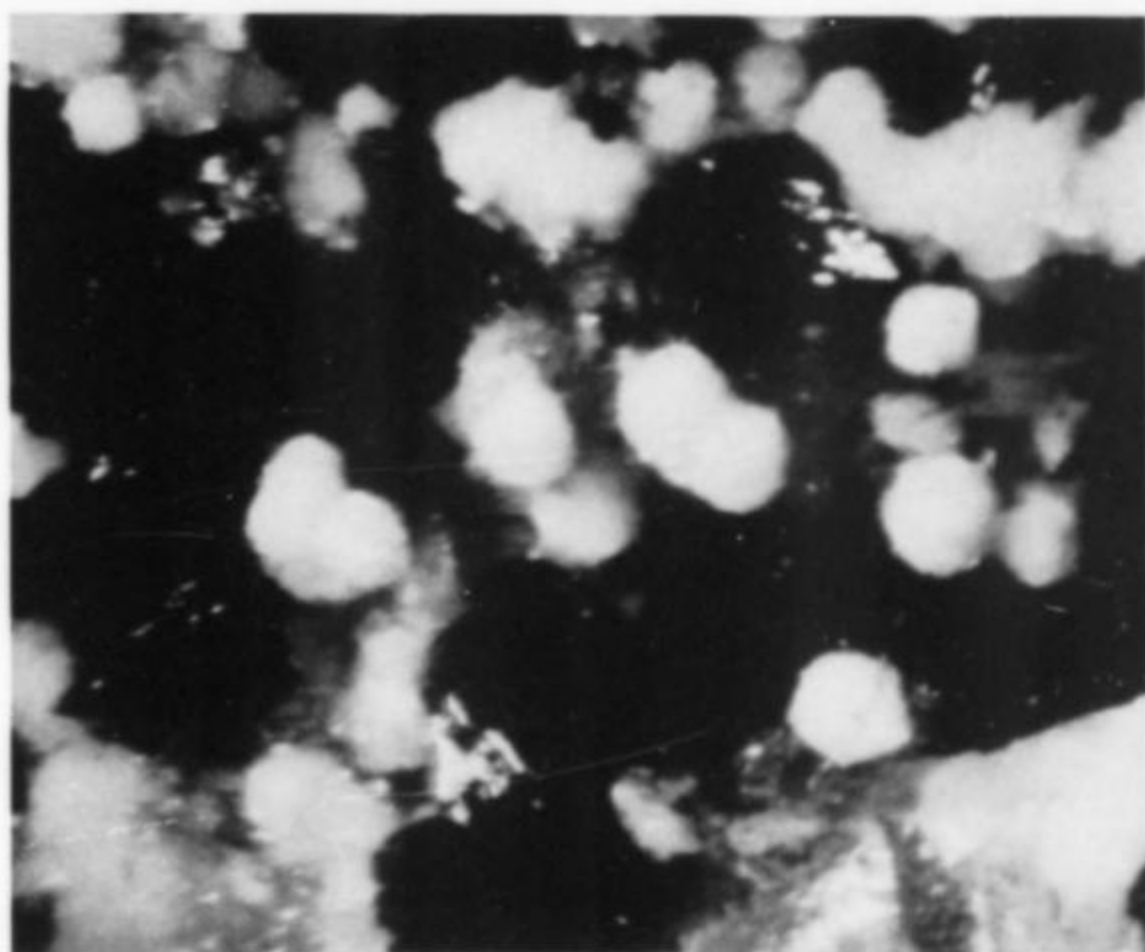


Figure 5. White, spherical aggregates of englishite associated with roscherite (dark olive-green) and eosphorite-childrenite (pale brown). Field of view is 6 mm. T. J. Campbell specimen and photo.



Figure 6. Bladed eosphorite-childrenite crystals with englishite. Cluster is 0.75 mm tall. T. J. Campbell specimen, Dan Behnke photo.

Eosphorite-Childrenite $(\text{Mn,Fe})^{+2}\text{Al}(\text{PO}_4)(\text{OH})_2 \cdot \text{H}_2\text{O}$

Eosphorite-childrenite is uncommon and occurs as pale brown, fan-like aggregates of transparent, platy to bladed, 0.5-mm crystals. Single crystals are similar in morphology to the material from Black Mountain and Red Hill, Rumford, Maine, as described by Hurlbut (1950). Crystals are flattened parallel to {010} with well developed {100} and poorly developed {010} faces; the {110} form is either not present or very poorly developed. The prisms are terminated by small {121} faces. Optical studies indicate that most of the material is an intermediate member of the series.

Fairfieldite $\text{Ca}_2(\text{Mn,Fe})^{+2}(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$

Fairfieldite is the manganese end-member of the fairfieldite-messelite solid solution series and is moderately abundant. The mineral occurs as bundles and lamellar aggregates of white, translucent, 0.5 to 5-mm, platy crystals which have a vitreous to pearly luster. It is also found as stacked, parallel aggregates of transparent to translucent crystals of similar size and luster. Less commonly,

fairfieldite is found as 0.3 to 3-mm thick fracture-fillings in beryl, commonly associated with dark olive-green roscherite. It also occurs with whitlockite and red montgomeryite; in these occurrences, the mineral is almost pure end-member fairfieldite. Pseudomorphs of todorokite after fairfieldite are common.

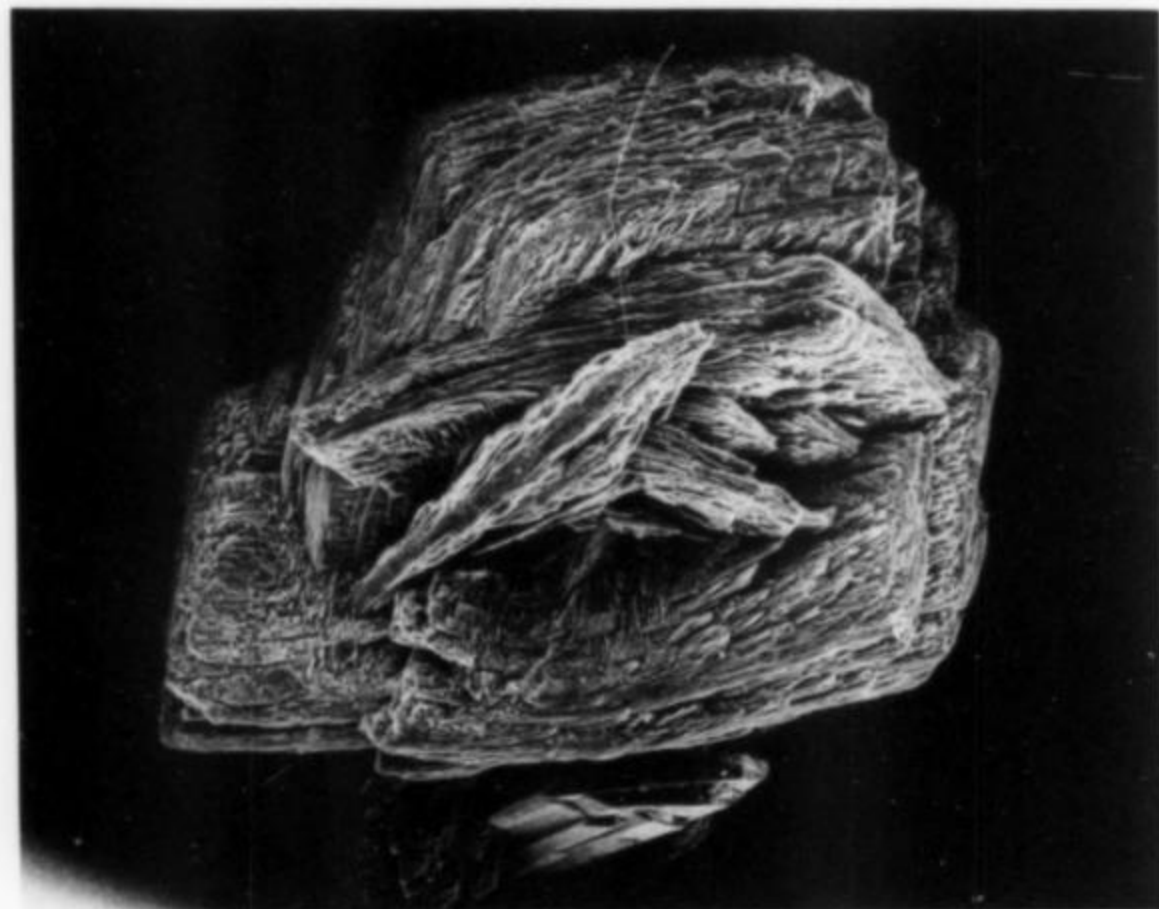


Figure 7. SEM photograph of fransoletite showing how aggregates of subparallel individuals create a large curved crystal. Aggregate is 0.5 mm across.

Fransoletite $H_2Ca_3Be_2(PO_4)_4 \cdot 4H_2O$

Fransoletite (Peacor *et al.*, 1983) is a new mineral from the Tip Top and occurs as translucent, milky to white, arrowhead-shaped aggregates up to 3 mm. These aggregates are composed of hundreds of microcrystals which are slightly offset relative to each other. Dominant forms of the composite crystals are {100}, {011} and {010}. Crystals are tabular on {010} and elongate parallel to [101]. Fransoletite is intimately associated with other Be-bearing minerals such as tiptopite, roscherite and hurlbutite. A chemical analysis of fransoletite, from Peacor *et al.* (1983), is given in Table 2.

Gordonite $MgAl_2(PO_4)_2(OH)_2 \cdot 8H_2O$

Gordonite has been found on only one specimen and occurs as colorless, transparent, vitreous, 1.5-mm prismatic crystals in subparallel groups, associated with robertsite and collinsite.

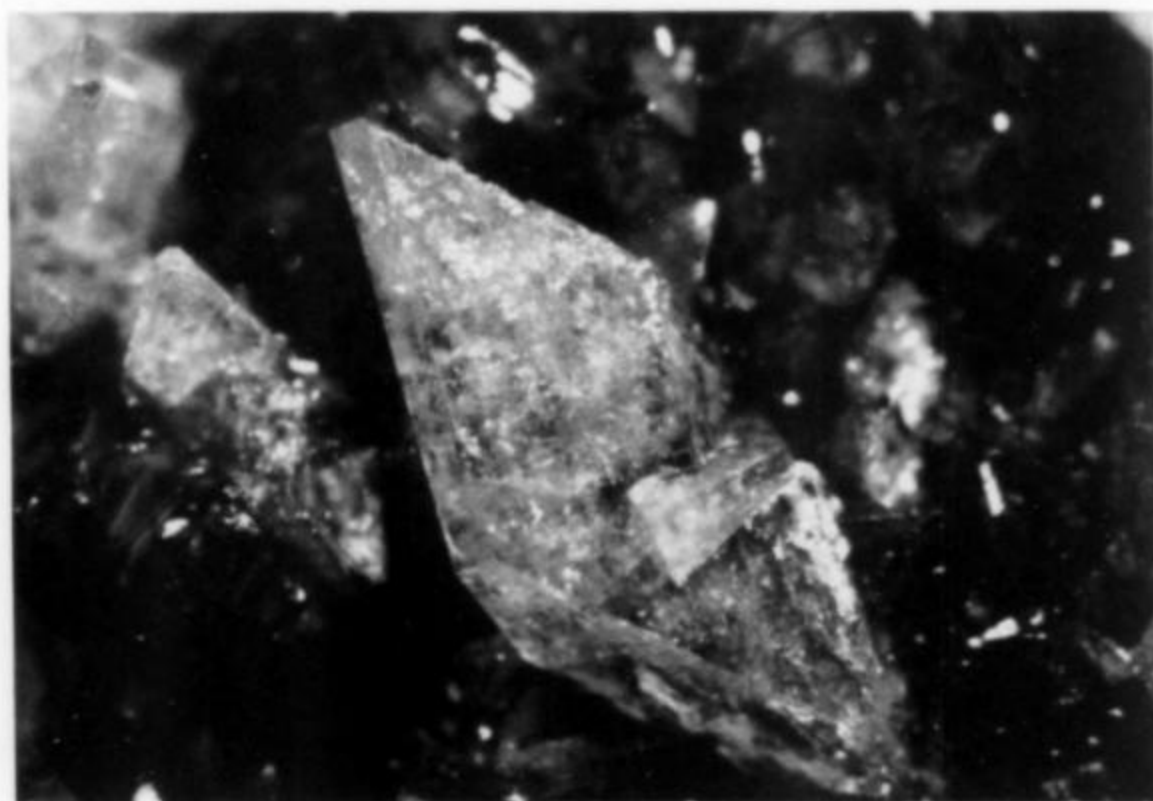


Figure 8. Hureaulite crystals to 3.2 mm on jahnsite. T. J. Campbell specimen, Dan Behnke photo.

Hureaulite $H_2(Mn,Fe)_5(PO_4)_4 \cdot 4H_2O$

Hureaulite is one of the most common secondary phosphates found at Tip Top. It occurs as vitreous, transparent to translucent, prismatic crystals, rarely tabular or equant. Dominant forms are {110}, {101} and {100}. It is found in a variety of colors ranging from very pale pink to deep orange-pink to red-brown. Crystals occur singly and in groups. Size ranges from a fraction of a millimeter up to 3.2 cm, some of which may be the largest and finest crystals in the world. Hureaulite is commonly associated with rockbridgeite and leucophosphite. Partial chemical analyses of Tip Top hureaulite are given in Table 2.

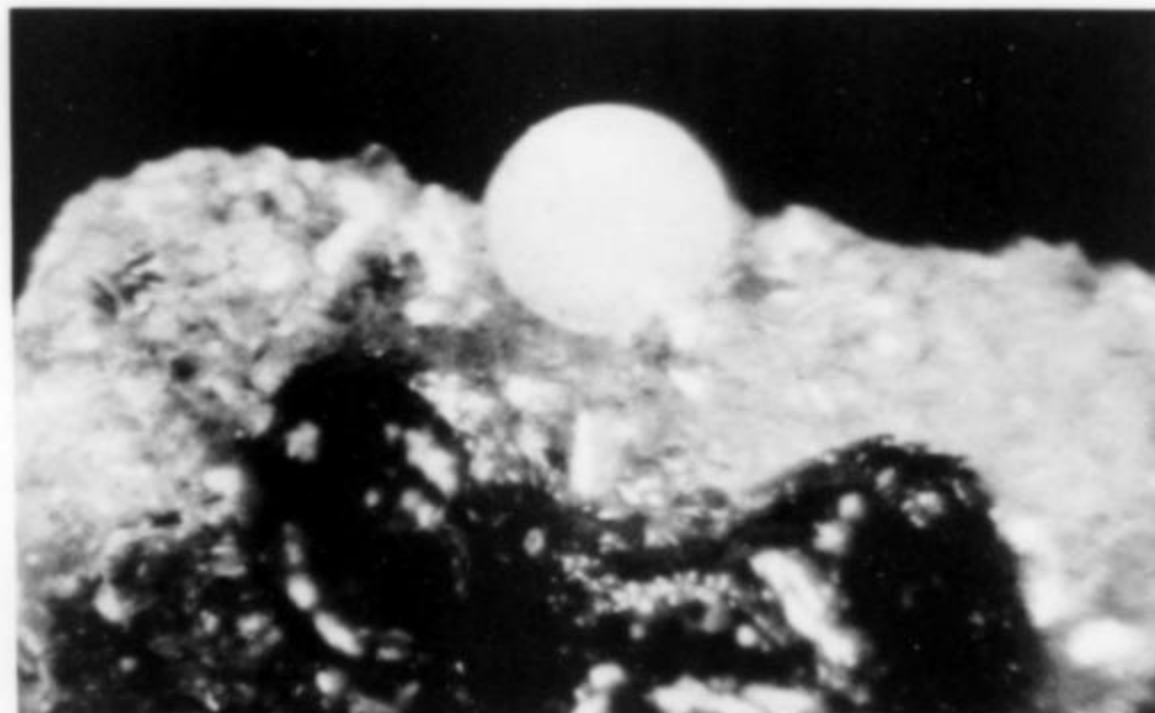


Figure 9. Hurlbutite spherule, 0.7 mm, on whitlockite with roscherite (red). T. J. Campbell specimen and photo.

Hurlbutite $CaBe_2(PO_4)_2$

Hurlbutite occurs as pale pink to bluish white, pearly spherules to 1 mm with a radial fibrous structure and is often associated with tiptopite, englishite and montgomeryite. These spherules are unlike hurlbutite from the type locality at the Smith Mine, Chandler's Mill, New Hampshire, where it occurs as stout prismatic crystals over 2.5 cm in length (Mrose, 1952).

Hydroxylapatite $Ca_5(PO_4)_3OH$

Hydroxylapatite is not common and is the least abundant apatite group mineral at this locality. It is similar in occurrence to carbonate-hydroxylapatite found in quartz or triphylite. Hydroxylapatite is found as colorless to white, hexagonal, bipyramidal crystals up to 0.5 mm. Members of the apatite group can easily be distinguished by their indices of refraction.

Hydroxyl-herderite $CaBePO_4(OH)$

Hydroxyl-herderite is chemically related to hurlbutite; the two minerals are very similar in appearance at the Tip Top mine but can be distinguished by X-ray diffraction methods. Hydroxyl-herderite most commonly occurs as colorless to white spherules associated with parascholzite and ehrlite (Grice and Robinson, 1984).

Jahnsite $CaMn^{+2}(Mg,Fe^{+2})_2Fe^{+3}(PO_4)_4(OH)_2 \cdot 8H_2O$

Jahnsite is abundant at the Tip Top mine and its varieties can be differentiated basically by color and occurrence.

The first type is the "brown" jahnsite examined by Moore (1974) in the first description of the species. It occurs as light to dark brown, deep red-brown and red-orange, vitreous, short to long prismatic crystals up to 1 cm in length which are vertically striated parallel to [010] and are often tabular parallel to {100}. The dominant forms include {001}, {100}, {201}, {201} and {111} with striae on {201} and {100} (Moore, 1974). Jahnsite is found as single crystals, clusters and subparallel groups of crystals. Brown jahnsite is generally found in triphylite pods which are moderately

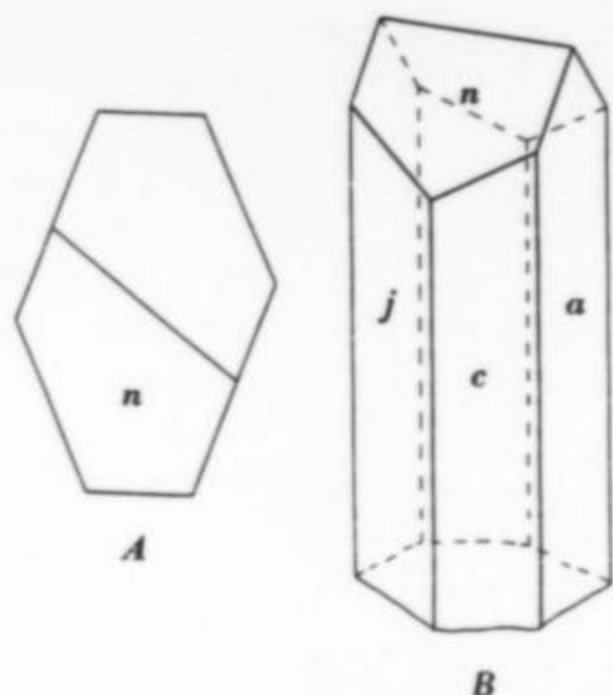


Figure 10. Crystal showing the typical forms of jahnsite and whiteite. The forms $c\{001\}$, $a\{100\}$, $j\{201\}$ and $n\{\bar{1}11\}$ are indicated. A. Plan view. B. Clinographic projection (b-axis polar) (Moore, 1978).

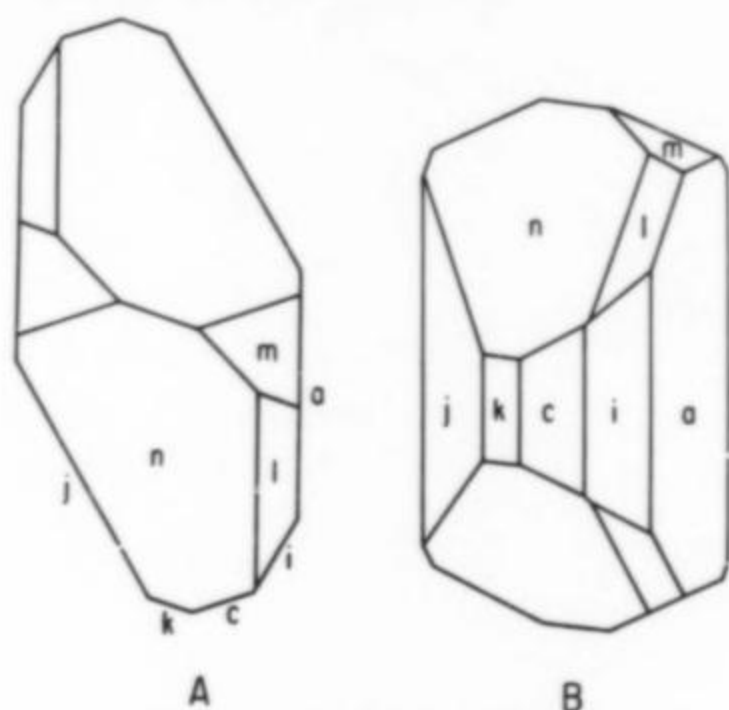


Figure 11. Crystal drawing of jahnsite showing the forms $c\{001\}$, $a\{100\}$, $i\{201\}$, $j\{201\}$, $k\{\bar{1}01\}$, $l\{011\}$, $m\{110\}$ and $n\{\bar{1}11\}$. A. Plan view. B. Clinographic projection (Moore, 1974).

altered. It is associated with rockbridgeite, hureaulite, leucophosphate and sometimes tavorite.

The second type of Tip Top jahnsite is the "green-yellow" variety. It is morphologically identical to the brown jahnsite but differs somewhat chemically in that it is lower in FeO and slightly higher in MgO and MnO. Color ranges from yellow-green to green-yellow to pale green. The green-yellow variety is largely restricted to more highly altered portions of triphylite nodules in or near the heterosite or ferrisicklerite rind. Common associated minerals include leucophosphate and messelite.

Kryzhanovskite $MnFe_2^+(PO_4)_2(OH)_2 \cdot H_2O$

Kryzhanovskite is rare and was more commonly encountered in earlier mine workings along with ludlamite, phosphoferrite and siderite. Kryzhanovskite occurs as red-brown, vitreous to dull, translucent crystals which replace pseudo-octahedral phosphoferrite crystals.

Laueite $MnFe_2^+(PO_4)_2(OH)_2 \cdot 8H_2O$

Laueite is a polymorph of strunzite and stewartite and occurs sparingly. It is found as vitreous, yellow-orange, transparent to translucent, wedge-shaped crystals up to 1.5 mm in length which are longitudinally striated. Dominant forms are $\{100\}$, $\{010\}$, $\{110\}$, $\{1\bar{1}0\}$, $\{0\bar{1}1\}$ and $\{011\}$. Laueite from the Tip Top is almost indistinguishable in habit from the material first described by Strunz (1954) from the Hagendorf pegmatite in Bavaria.

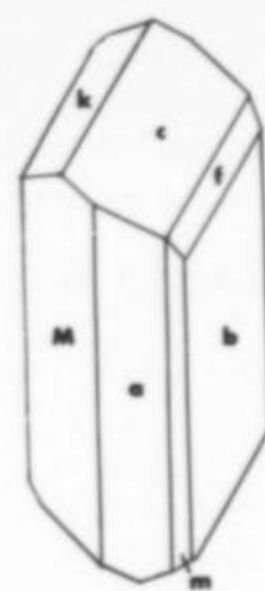


Figure 12. Crystal drawing of laueite showing the forms $a\{100\}$, $b\{010\}$, $m\{110\}$, $M\{1\bar{1}0\}$, $k\{0\bar{1}1\}$ and $f\{011\}$. Slightly modified from Strunz (1954).

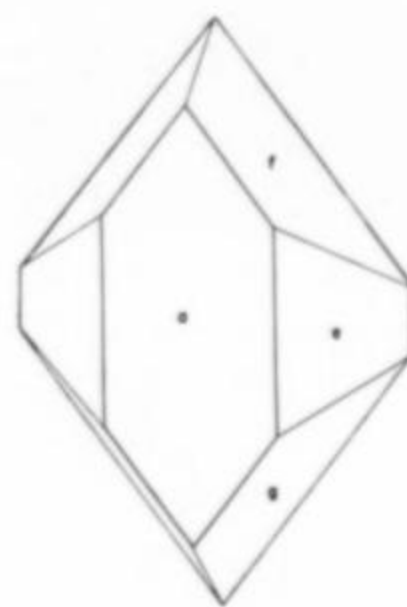


Figure 13. Crystal drawing depicting typical forms for leucophosphate and tinsleyite; $a\{100\}$, $e\{210\}$, $f\{111\}$, $g\{1\bar{1}1\}$ (Lindberg, 1957).

Leucophosphate $KFe_2^+(PO_4)_2(OH) \cdot 2H_2O$

Leucophosphate forms a series with the new mineral tinsleyite, the aluminum analogue (Dunn *et al.*, 1984a). Leucophosphate is one of the most abundant and earliest formed secondary phosphates found at Tip Top. It occurs as vitreous, pale brown to green-brown and brownish purple, short to long prismatic, diamond-shaped crystals up to 1 cm, though 1 mm crystals are most common. Dominant forms are $\{100\}$, $\{210\}$, $\{111\}$ and $\{1\bar{1}1\}$. These forms are identical to those found on crystals from the Sapucaia pegmatite, Brazil, as described by Lindberg (1957). The $\{100\}$ and $\{210\}$ forms may be slightly striated. Some crystals show large $\{100\}$ and $\{210\}$ faces and poorly developed $\{111\}$ and $\{1\bar{1}1\}$ faces. A parting surface along $\{100\}$ is evident in some crystals. This parting was mistakenly reported as a cleavage by Lindberg (1957). Crystals of leucophosphate occur singly and sometimes doubly terminated, and as parallel to subparallel groups of crystals. Leucophosphate crystals in stacked parallel growth on rockbridgeite needles are not uncommon. Leucophosphate also occurs as green to green-brown, granular masses up to several cm^3 as a replacement of triphylite. Leucophosphate can be differentiated from tinsleyite optically and by microprobe analysis.

Lithiophosphate Li_3PO_4

Lithiophosphate occurs as blocky, translucent, colorless to turbid, dull vitreous, 0.5-mm crystals in clusters up to 2.5 mm associated with leucophosphate, hureaulite, switzerite, jahnsite and rockbridgeite. This is the first reported occurrence of this mineral at the mine.

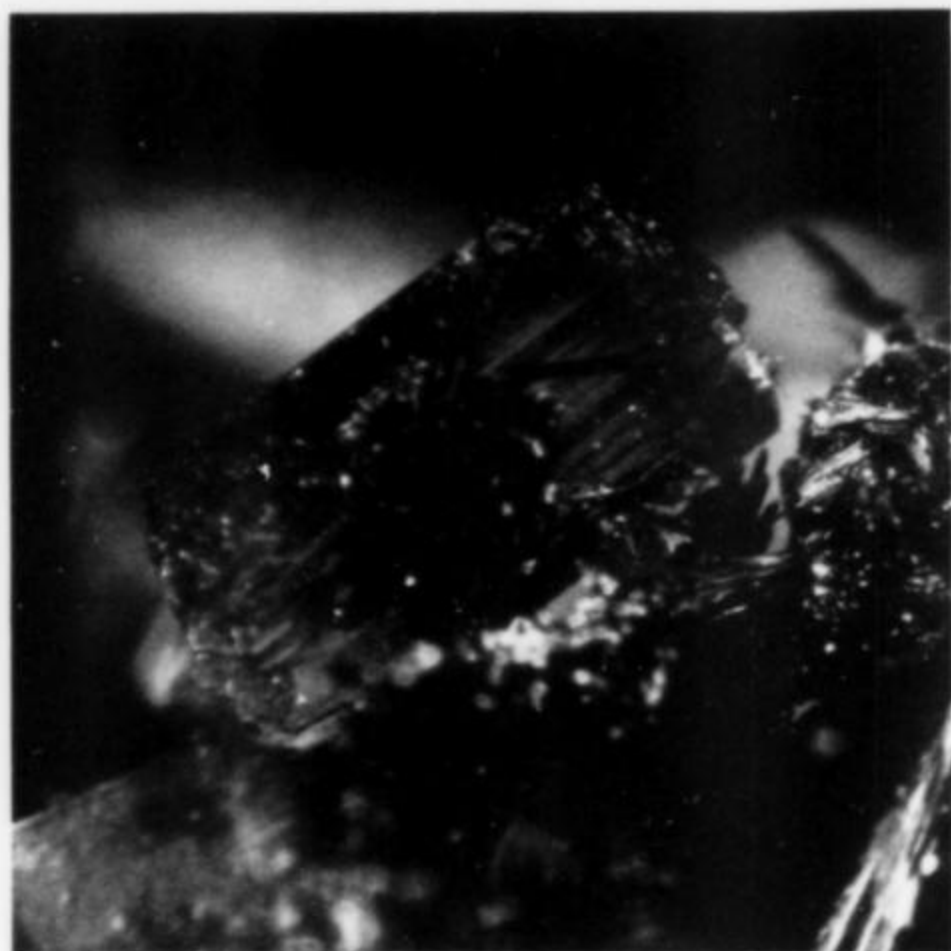


Figure 14. A typical brown jahnsite crystal on rockbridgeite. Jahnsite crystal is 0.7 mm in length. T. J. Campbell specimen and photo.



Figure 15. Prismatic laueite crystal 0.7 mm long. T. J. Campbell specimen, Dan Behnke photo.



Figure 16. Jahnsite crystal group 0.8 mm tall. T. J. Campbell specimen, Dan Behnke photo.

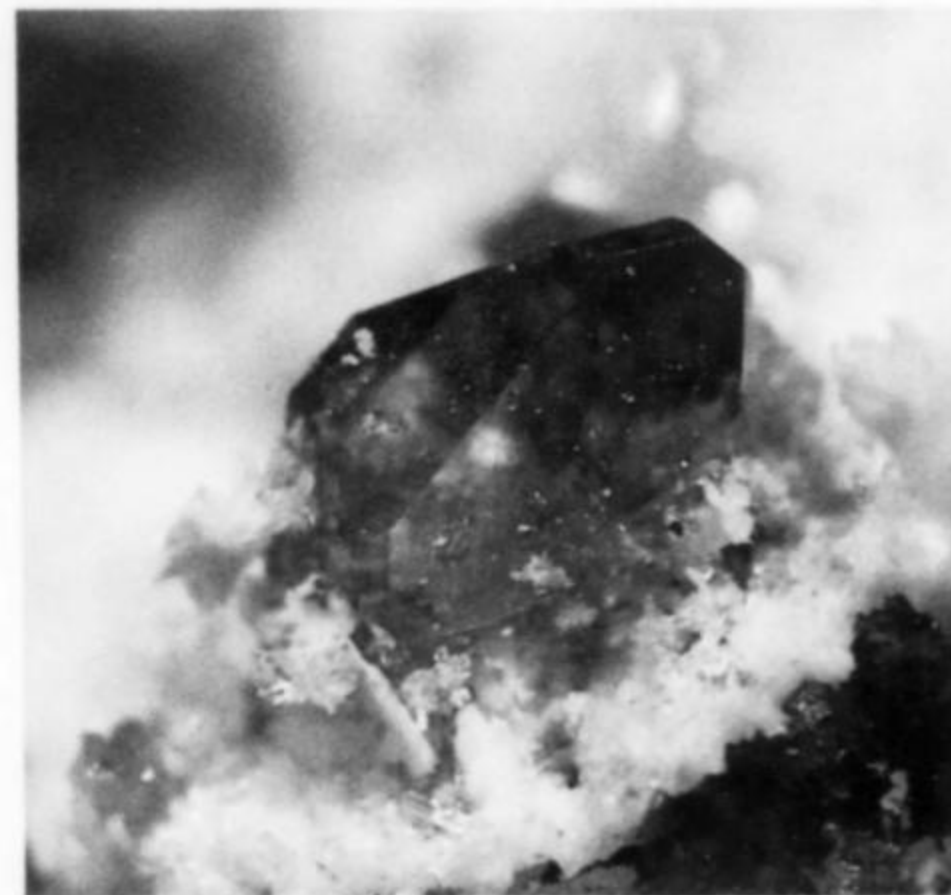


Figure 17. Leucophosphite crystal 1.1 mm in length. T. J. Campbell specimen, Dan Behnke photo.

Ludlamite $\text{Fe}_3^{+2}(\text{PO}_4)_2 \cdot 4\text{H}_2\text{O}$

Ludlamite is uncommon now, but was more abundant in the earlier workings. It occurs as vitreous transparent to translucent, pale green to bright green, thick, tabular crystals up to 3 mm. Ludlamite is also found as granular masses up to 2 cm³ replacing triphylite and is generally associated with siderite.

Messelite $\text{Ca}_2(\text{Fe,Mn})^{+2}(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$

Messelite, the iron analog of fairfieldite, is moderately abundant at the Tip Top mine. Although messelite is morphologically identical to fairfieldite and has many of the same physical properties such as color, luster, hardness and cleavage, it can be distinguished from fairfieldite chemically and optically. The occurrences of fairfieldite and messelite are distinct; messelite is largely restricted to highly altered pods of triphylite and is very commonly associated with green jahnsite whereas fairfieldite is chiefly found with montgomeryite and roscherite in fractures in beryl and quartz. A partial chemical analysis of messelite is given in Table 2.

Mitridatite $\text{Ca}_3\text{Fe}_4^{+3}(\text{PO}_4)_4(\text{OH})_6 \cdot 3\text{H}_2\text{O}$

Mitridatite, the iron analog of robertsite, is a very abundant secondary phosphate at this mine. It usually occurs as green stains and crusts coating broad surfaces of pegmatite exposures which have been subjected to weathering.

Other than occurring as a weathering product, mitridatite occurs as deep red to bronze-red, platy crystals up to 2 mm which commonly form spherical aggregates and thin crusts in altered triphylite. Mitridatite can also be found lining fractures in beryl and quartz. Prismatic, barrel-shaped crystals of mitridatite, very similar to robertsite, have been found, in some cases as a nucleus on which a thin continuous rind of robertsite has been deposited. Mitridatite can be distinguished from robertsite by its green streak.

Montgomeryite $\text{Ca}_4\text{MgAl}_4(\text{PO}_4)_6(\text{OH})_4 \cdot 12\text{H}_2\text{O}$

Montgomeryite as described by Dunn *et al.* (1983), is found as bladed, prismatic crystals, striated parallel to [001]. The morphology is similar to that of the type material described by Larsen



Figure 18. Spherical aggregate of mitridatite 0.56 mm in diameter. T. J. Campbell specimen, Dan Behnke photo.

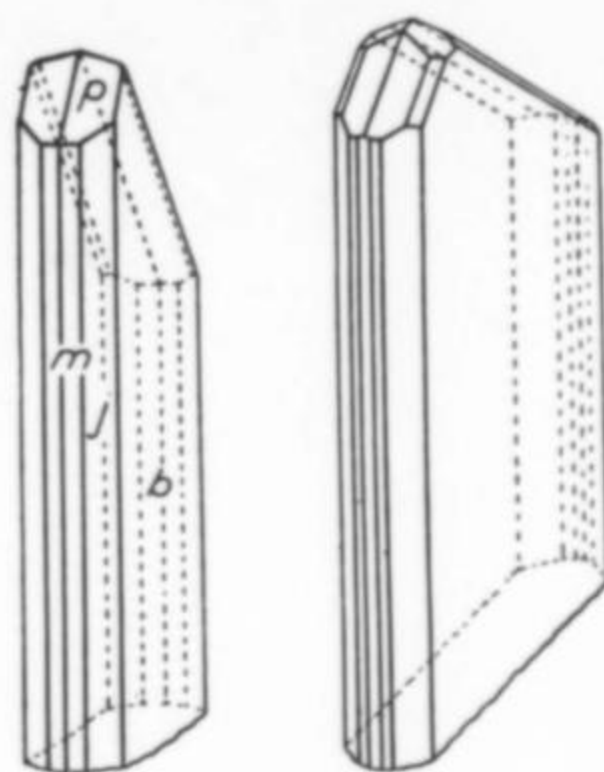


Figure 19. Crystals of Fairfield, Utah, montgomeryite which show forms that are very similar to those from Tip Top (Larsen, 1940).

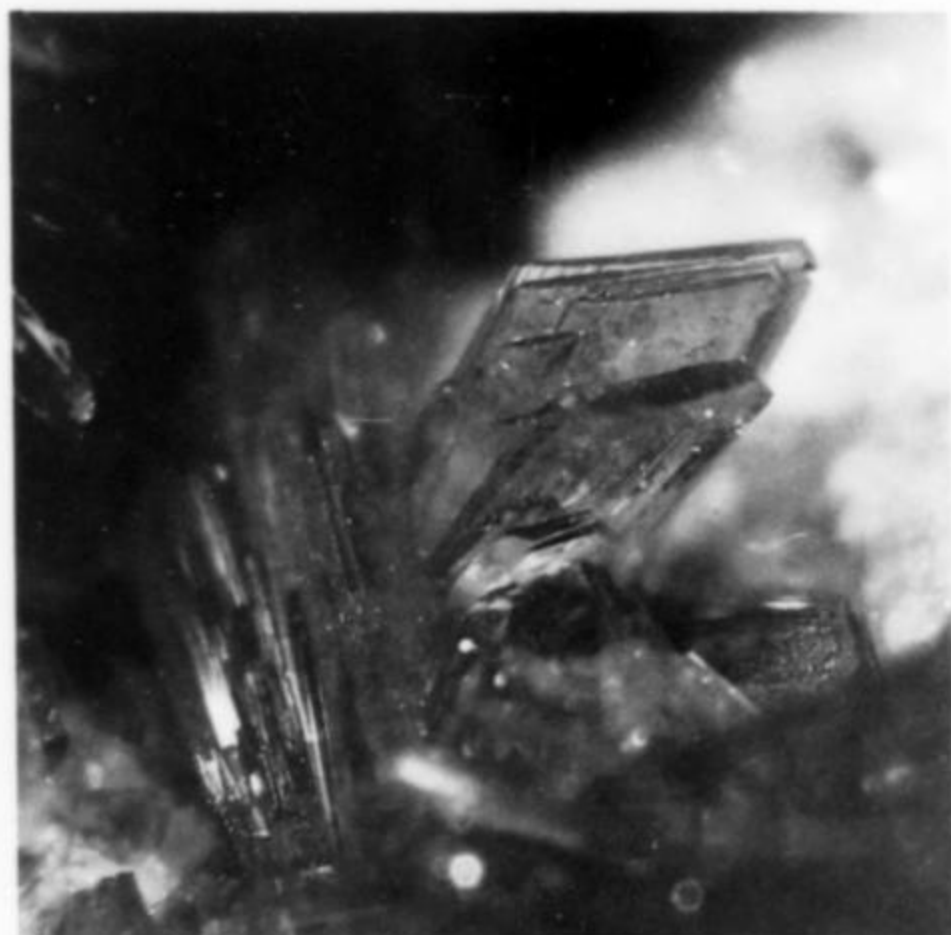


Figure 20. Red montgomeryite crystals with englishite. Largest crystal is about 1 mm tall. T. J. Campbell specimen, Dan Behnke photo.

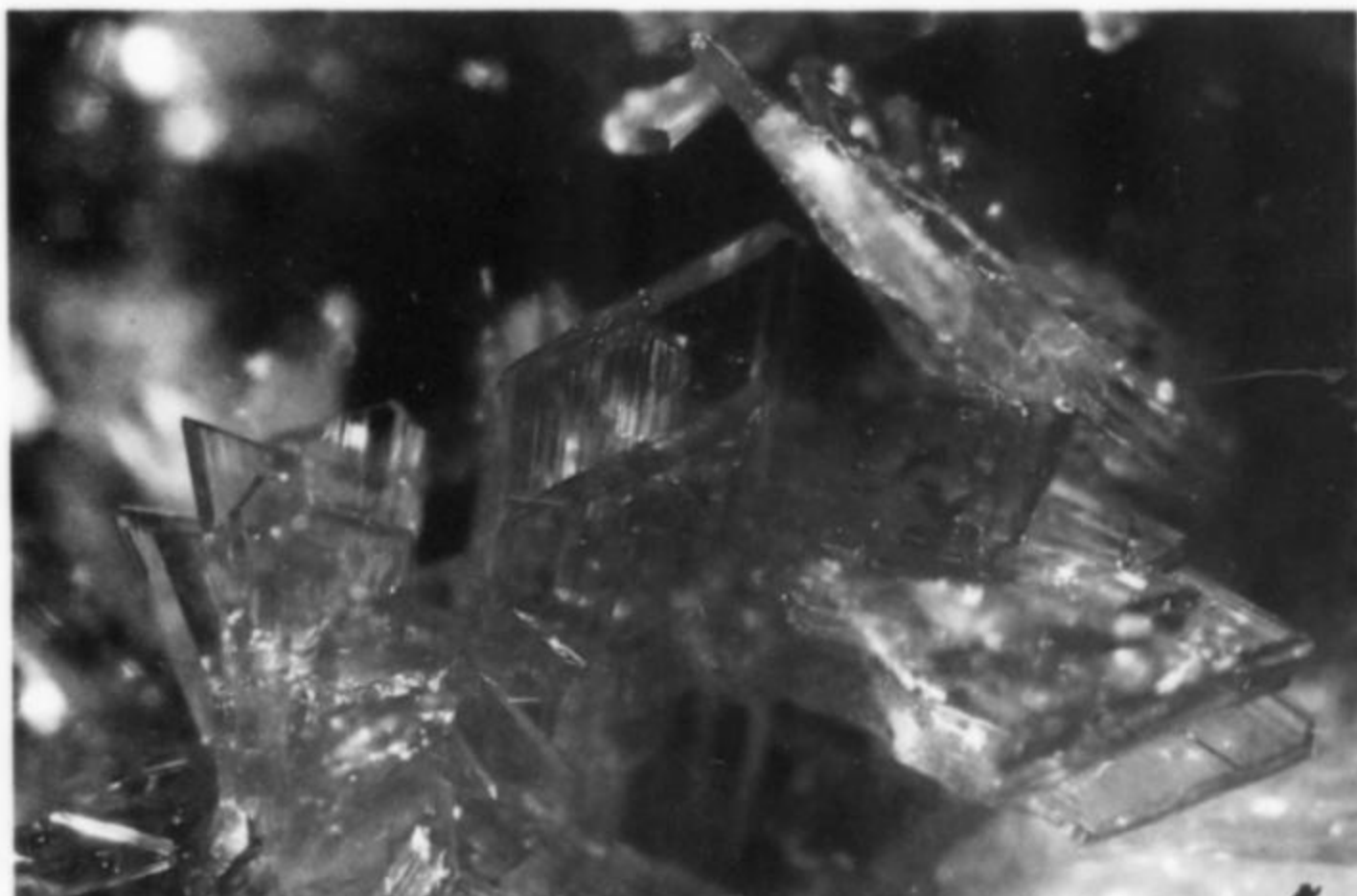


Figure 21. Red montgomeryite crystal group. Largest crystal is 1.25 mm in length. T. J. Campbell specimen, Dan Behnke photo.

(1940) from the variscite nodules found near Fairfield, Utah, except that crystals from the Tip Top mine are markedly elongate parallel to [001]. Montgomeryite commonly occurs as radial clusters and divergent fascicles up to 6 mm in diameter and as single crystals to 5 mm. Montgomeryite from this locality is found in a variety of colors ranging from colorless to shades of yellow and pink to bright red, with pink to red colors dominant. Montgomeryite crystals are often zoned and on many specimens the crystals are bright red at the base and gradually grade to pink or colorless, rarely to yellow or green-yellow. Certain colors of montgomeryite, along with roschertite, help characterize the different phosphate suites as discussed below. Chemical analyses of montgomeryite are given in Table 2.

Parascholzite $\text{CaZn}_2(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$

Parascholzite was described by Sturman *et al.* (1981) from the Hagedorf pegmatite, Bavaria, and is dimorphous with scholzite. Only two specimens containing parascholzite have been found to date at the Tip Top mine. The mineral occurs as colorless to white,

vitreous crystals elongated parallel to [001] and measuring up to 1 mm. Dominant forms are the prism and pinacoid.

Phosphoferrite $(\text{Fe,Mn})_3(\text{PO}_4)_2 \cdot 3\text{H}_2\text{O}$

Phosphoferrite is the iron end-member of the phosphoferrite-reddingite series and is not abundant. Like ludlamite and kryzhanovskite, it was more abundant in the earlier workings. Phosphoferrite occurs as colorless to pale green, vitreous to resinous, translucent, 0.5-mm, pseudo-octahedral crystals and as small granular masses associated with ludlamite and siderite.

Robertsite $\text{Ca}_3\text{Mn}_4^{+3}(\text{PO}_4)_4(\text{OH})_6 \cdot 3\text{H}_2\text{O}$

Robertsite was first described from Tip Top by Moore (1974). The mineral is moderately abundant at this locality and occurs in a variety of forms. It is most commonly found as brown-red, cleavable, massive seams up to 3 mm thick in triphylite and quartz, rarely in perthite or beryl. In open spaces, robertsite is found as vitreous, translucent to almost opaque, deep red to black, tabular,

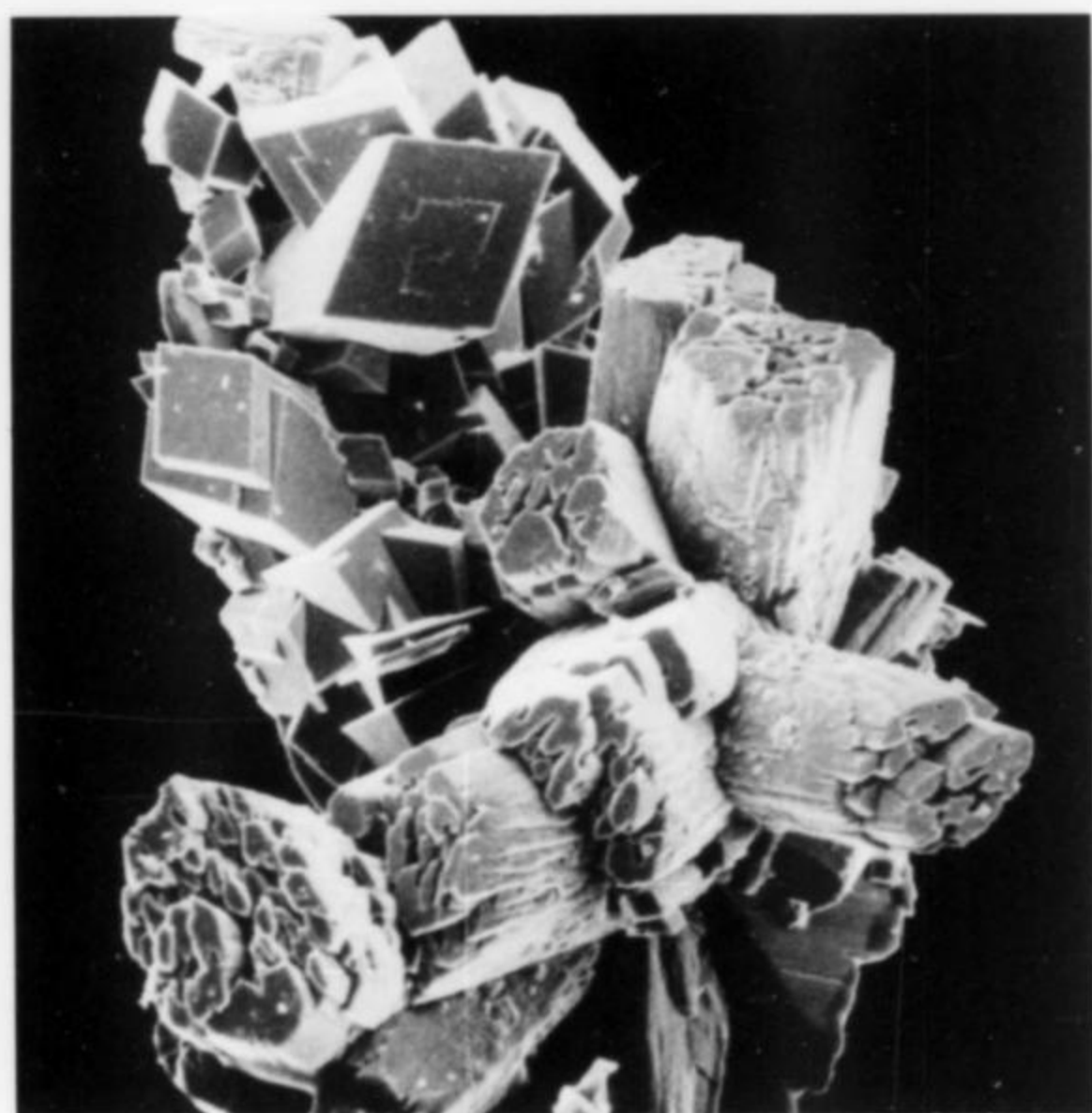


Figure 22. SEM photomicrograph of robertsite crystals in parallel growth intergrown with whitlockite. Cluster is 1.5 mm tall.

prismatic crystals to 1 mm; as black, spherical or botryoidal aggregates to 7 mm with a radial-fibrous structure and less commonly as thin platy pseudo-hexagonal crystals. Common forms are {001}, {100} and {031} as determined by Moore (1974). Crystals are sometimes twinned by rotation normal to {100}. Robertsite is present in most of the phosphate assemblages at Tip Top.

Rockbridgeite-Frondelite Series $(\text{Fe,Mn})^{+2}\text{Fe}_4^{+3}(\text{PO}_4)_3(\text{OH})_5$

Members of the rockbridgeite-frondelite series are the most abundant of all secondary phosphates at this pegmatite with rockbridgeite by far the dominant species in the series. Rockbridgeite is chiefly found as a massive replacement of triphylite. It sometimes occurs as fracture fillings in quartz and rarely in beryl or perthite. The appearance is characteristic: compact fibrous to bladed, deep green to black masses when unoxidized; or brown, bladed fibers when highly oxidized. Fracture-fillings of rockbridgeite may exceed 1 cm in thickness. In open spaces, rockbridgeite commonly occurs as black, botryoidal aggregates that are radial-fibrous in cross-section and as sprays of more isolated, greenish black, prismatic crystals up to 1 mm, rarely up to 1 cm.

Some intermediate members of this series found at Tip Top are termed frondelite-rockbridgeite; they have compositions and optical properties between frondelite and rockbridgeite. Some of this material is more Mn-rich and is thus frondelite. Frondelite-rockbridgeite and frondelite are associated with tinsleyite, robertsite and tavorite in highly oxidized pods of triphylite. Partial chemical analyses of minerals in the rockbridgeite-frondelite series are given in Table 2.

Roscherite $\text{Ca}(\text{Al}^{+3}, \text{Fe}^{+2}\text{Mn,Mg})_3\text{Be}_2(\text{PO}_4)_3(\text{OH})_3 \cdot 2\text{H}_2\text{O}$ (?)

Roscherite occurs as 0.5 to 3 mm spherical aggregates and loose clusters composed of vitreous to resinous, tabular, pseudo-hexagonal, 0.3-mm crystals. Most crystals are composed of the forms {100}, {010} and {111}, listed in order of dominance. They appear to be identical to forms found by Lindberg (1958) on crystals from the Sapucaia pegmatite, Brazil. Additional forms, however, have been observed. Roscherite also occurs as 0.5 to 4 mm thick granular

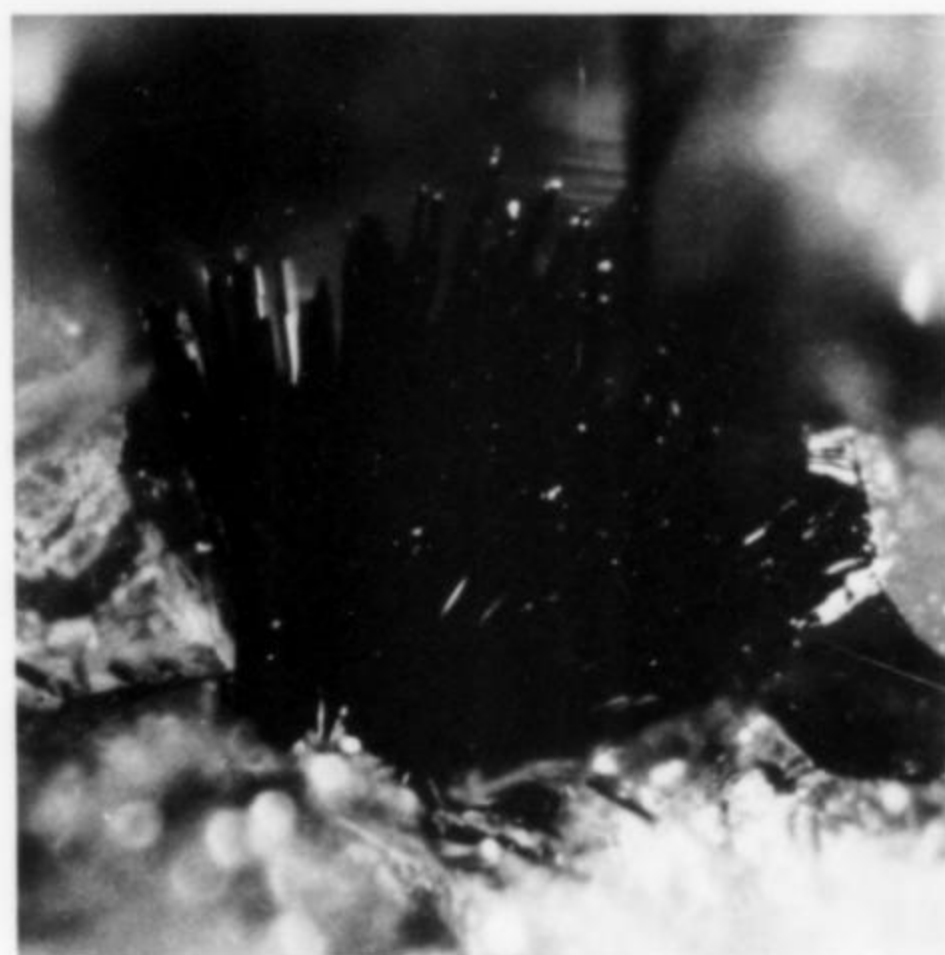


Figure 23. A 0.9-mm rockbridgeite spray. T. J. Campbell specimen, Dan Behnke photo.

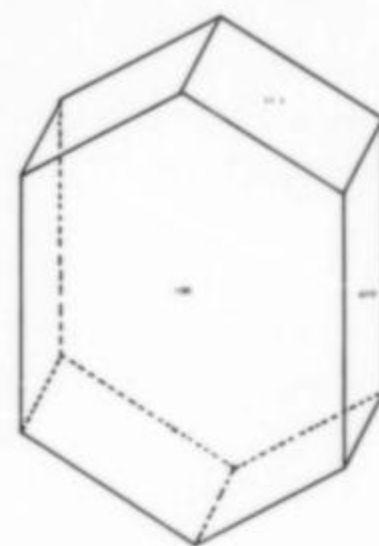


Figure 24. Crystal of roscherite showing the forms $a\{100\}$, $b\{010\}$ and $K\{111\}$ (Lindberg, 1958) (Sapucaia pegmatite).

massive fracture fillings in beryl. Roscherite is found in a variety of colors such as orange, lavender, red-orange, red-brown and dark olive-green to almost black. Some of these colors have been used as an aid to help characterize the various phosphate suites as discussed below. Chemical data for Tip Top mine roscherite are listed in Table 2.

Scholzite $\text{CaZn}_2(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$

Scholzite was just recently found at Tip Top and only one specimen is known. It occurs in two different morphologies on this specimen: colorless, slender prismatic, 3-mm crystals striated parallel to [001] and white, translucent, blocky crystals to 0.5 mm; both associated with mitridatite.

Segelerite $\text{CaMgFe}^{+3}(\text{PO}_4)_2(\text{OH}) \cdot 4\text{H}_2\text{O}$

Segelerite is the iron analog of overite and is chemically related to jahnsite. It was first described by Moore (1974) from Tip Top and is also known from Milgun Station, Western Australia. The mineral is rare and occurs as vitreous, pale green, orthorhombic prisms up to 1 mm. Crystals have a nearly square cross-section, are striated parallel to [001] and are terminated by the pyramid. Common forms determined by Moore (1974) are {100}, {010}, {110} and {121}. Segelerite is restricted to altered triphylite and is in close association with robertsite, collinsite and jahnsite. Segelerite can be



Figure 25. Aggregates of dark olive-green roscherite associated with englishite (white) and eosphorite-childrenite. Field of view is 4 mm. T. J. Campbell specimen, Dan Behnke photo.

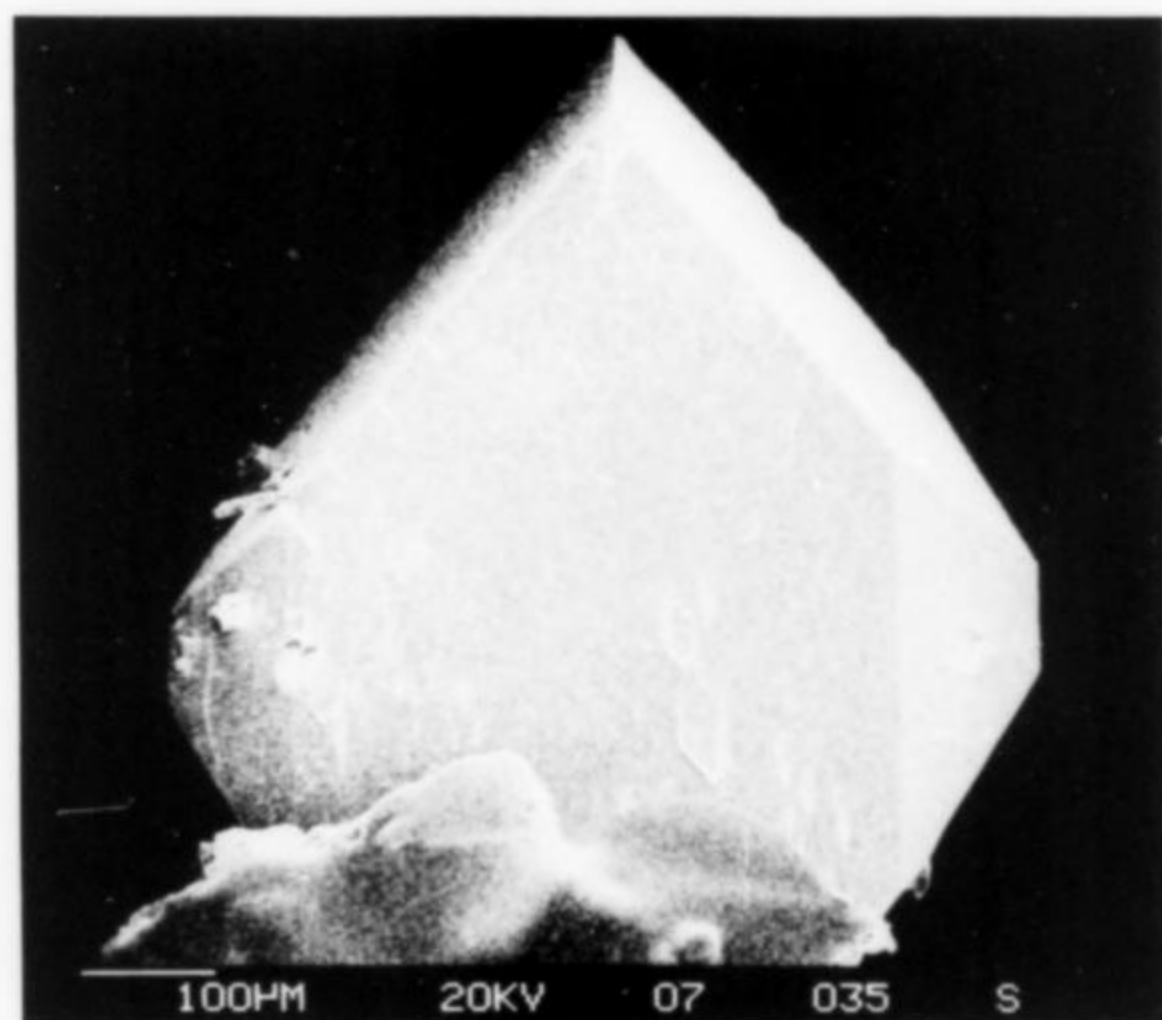


Figure 26. SEM photomicrograph of tinsleyite. The large face is {100}.

visually confused with jahnsite but can be distinguished by chemical or X-ray diffraction methods.

Stewartite $Mn^{+2}Fe_2^{+3}(PO_4)_2(OH)_2 \cdot 8H_2O$

Stewartite is a polymorph of laueite and strunzite and is rare at this mine. It occurs as yellow to yellow-brown, vitreous, bladed, triclinic crystals which do not exceed 0.3 mm and can be easily confused with laueite. Stewartite can be distinguished from laueite by its steeper termination and lack of striations.

Strengite $FePO_4 \cdot 2H_2O$

Strengite has been found on one specimen as white granular aggregates (J. Van Velthuizen, personal communication, 1984).

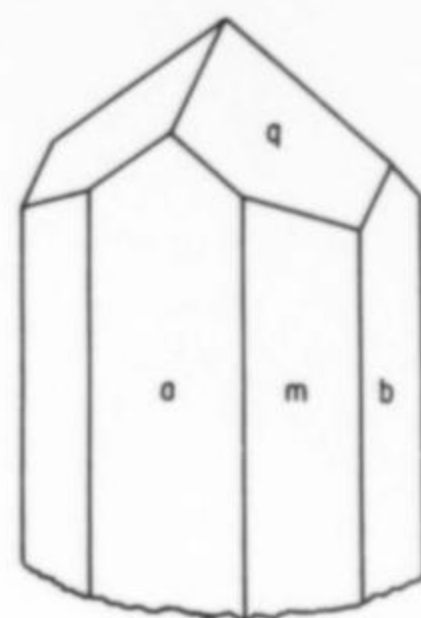


Figure 27. Segelerite crystal showing the forms $a\{100\}$, $b\{010\}$, $m\{110\}$ and $q\{121\}$. (Moore, 1974).

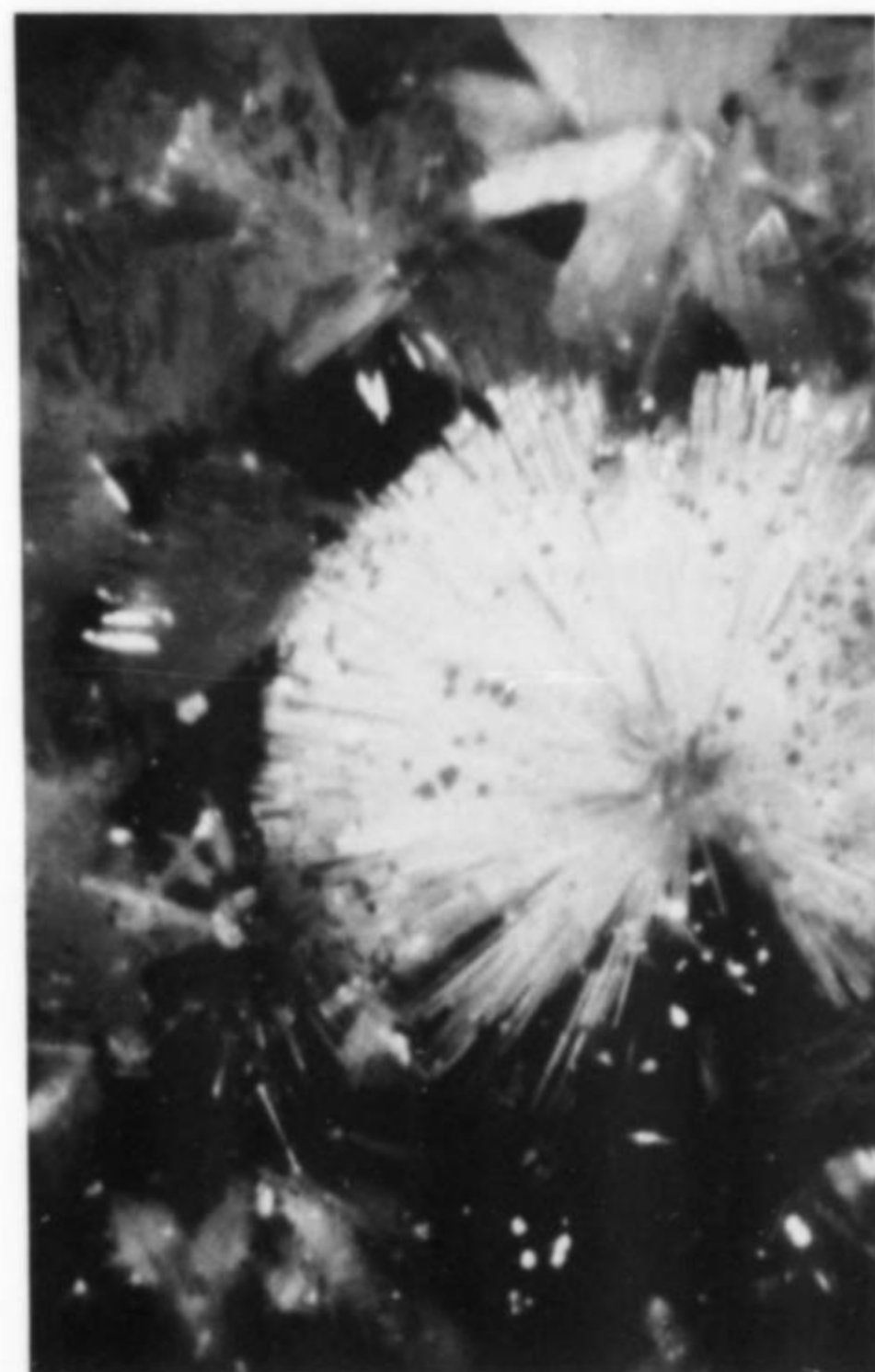


Figure 28. Radial spray of tiptopite associated with roscherite (red) and montgomeryite (pale brown). Spray is 2 mm across. T. J. Campbell specimen and photo.

Strunzite $Mn^{+2}Fe_2^{+3}(PO_4)_2(OH)_2 \cdot 8H_2O$

Strunzite is not presently abundant; it was more abundant in earlier mine workings. It is found as divergent tufts and felted aggregates composed of straw-yellow, silky, lath-like crystals flattened on {010}.

Switzerite $(Mn,Fe)_3(PO_4)_2 \cdot 4H_2O$

Switzerite is not abundant and occurs as silky, white to pale pink or yellowish interlocking aggregates up to 5 mm composed of thin,



Figure 29. Rounded aggregates of red roscherite with tiptopite, hurlbutite and todorokite (black). Field of view is 4 mm. T. J. Campbell specimen and photo.

Figure 30. A group of tinsleyite-coated leucophosphate crystals. Largest crystal is 0.9 mm tall. T. J. Campbell specimen, Dan Behnke photo.

bladed crystals, striated and elongated parallel to [001]. Switzerite is, without exception, associated with hureaulite which generally displays a certain degree of dissolution. Microprobe data for switzerite is given in Table 2.

Tavorite $\text{LiFe}^{+3}\text{PO}_4(\text{OH})$

Tavorite is isostructural with montebasite. It is moderately abundant in small amounts and occurs as bright yellow-green, fine-grained crystalline aggregates or earthy blebs disseminated in triphylite. Tavorite has been found rarely in masses up to 7 cm³ replacing triphylite. Although crystals are not common, this is probably the world's best locality for tavorite crystals; they are transparent to translucent, equant to short prismatic, with a vitreous to dull luster and can be found up to 2 mm in maximum dimension, though 0.3 mm crystals are most common. A chemical analysis for tavorite is given in Table 2.

Tinsleyite $\text{KAl}_2(\text{PO}_4)_2(\text{OH}) \cdot 2\text{H}_2\text{O}$

Tinsleyite (Dunn *et al.*, 1984a) is a new mineral from the Tip Top mine and is the aluminum analog of leucophosphate. It commonly occurs as magenta-red, thin ($\ll 0.1$ mm), morphologically continuous layers on earlier formed leucophosphate; parting surfaces occur between the two. Tinsleyite is rarely found as magenta-red short prismatic crystals morphologically identical to leucophosphate. A chemical analysis of tinsleyite is given in Table 2. It should also be noted that not all magenta colored material is tinsleyite. Some of the magenta colored specimens analyzed were essentially pure leucophosphate with only traces of Al (Dunn *et al.*, 1984a). Therefore, the dark red-purple color is not a diagnostic property. Optical examination (refractive indices) and microprobe analysis are the most reliable means for differentiating between tinsleyite and leucophosphate.

Tiptopite $(\text{Li}, \text{K}, \text{Na}, \text{Ca}, \square)_3\text{Be}_6(\text{PO}_4)_6(\text{OH})_4$

Tiptopite (Grice *et al.*, 1985) was the first new mineral encountered in recent work at the Tip Top. It occurs as radial sprays and divergent fascicles composed of vitreous, colorless, acicular, prismatic crystals terminated by the pinacoid. Crystals are sometimes cavernous or hollow.

Tiptopite was moderately abundant at the mine in 1982 but subsequent mining activity removed most, if not all, of the area where it was found. Tiptopite occurred predominantly along fracture surfaces in beryl, occasionally in quartz and rarely in perthite. Its occurrence is discussed in detail by Grice *et al.* (1985).

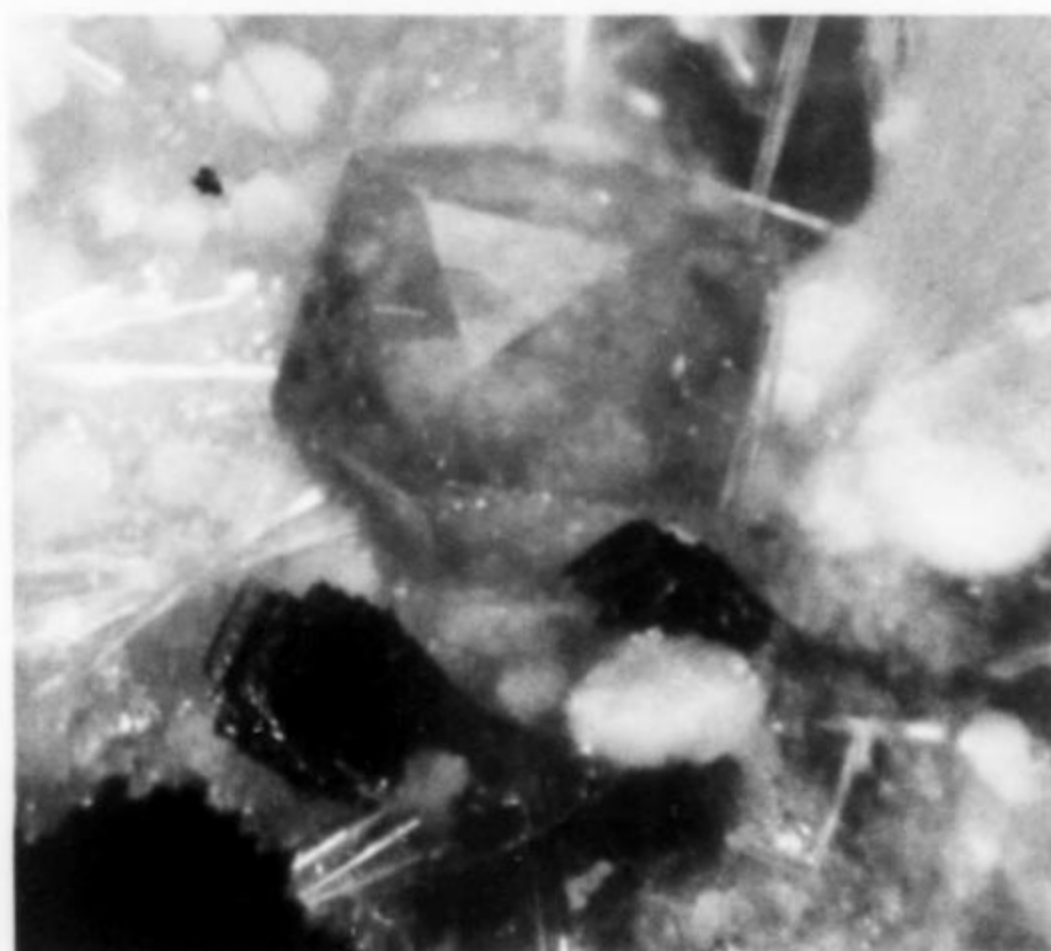
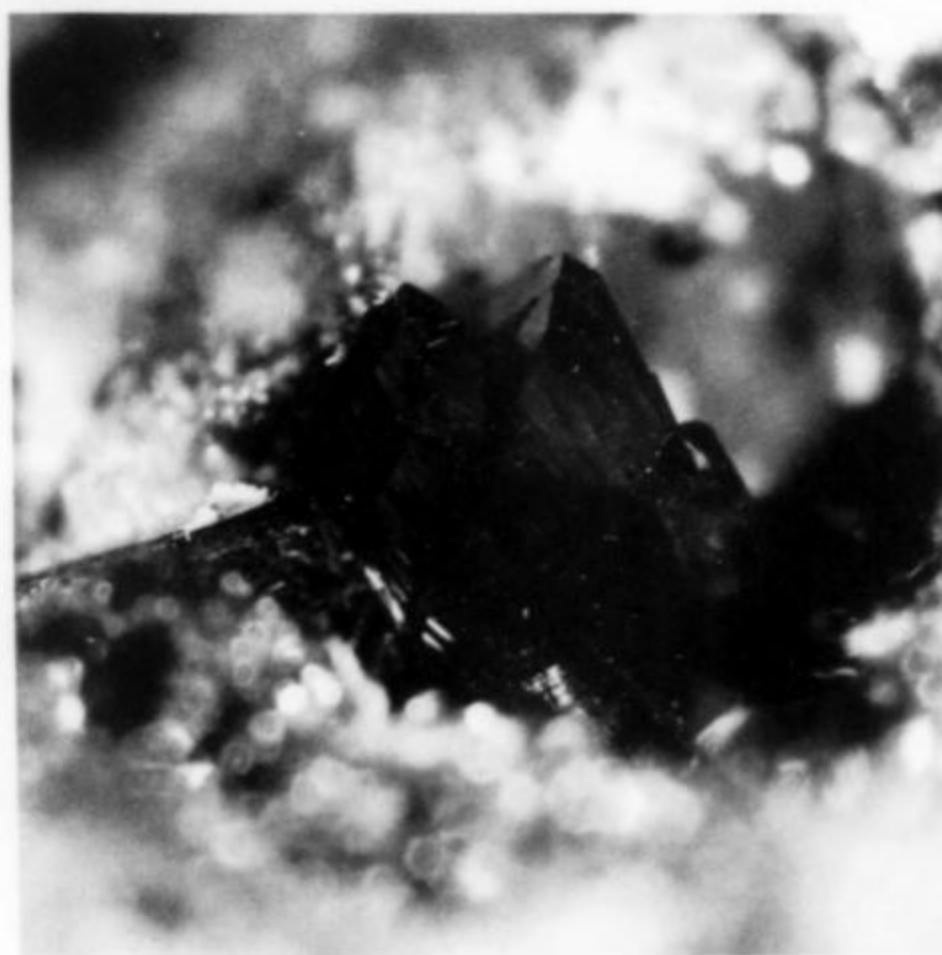


Figure 31. Photomicrograph of a 0.5-mm crystal of the undescribed phosphate with roscherite (dark olive-green), englischite (white spherules), and tiptopite (colorless needles). T. J. Campbell specimen and photo.



Figure 32. SEM photomicrograph of a tiptopite spray showing the cavernous nature on a few of the crystals.

Vivianite $\text{Fe}_3^{+2}(\text{PO}_4)_2 \cdot 8\text{H}_2\text{O}$

Vivianite is one of the more abundant and earliest formed secondary phosphates found at the mine. It occurs as pale to dark blue, cleavable, massive veinlets or small pod-like masses replacing triphylite and as blue, 1 to 3-mm, tabular, bladed or prismatic crystals often flattened on {010} and somewhat elongate parallel to [001]. Several crystals exceeding 4 cm have been collected. Vivianite has a vitreous luster, a fibrous fracture and a perfect {010} cleavage which is easily produced. Thin lamellae are flexible. Fresh, unaltered vivianite is generally pale green, the color darkening slowly upon exposure. Some of the larger vivianite crystals show inclusions of 0.2 mm sprays of rockbridgeite. Vivianite also alters to a yellow, granular, unidentified mineral.

Whiteite $\text{Ca}(\text{Fe,Mn})^{+2}\text{Mg}_2\text{Al}_2(\text{PO}_4)_4 \cdot 8\text{H}_2\text{O}$

Whiteite is the aluminum analog of jahnsite (Moore and Ito, 1978). Whiteite is not abundant here and has a restricted mode of occurrence. It is restricted to roscherite-bearing assemblages that occur as fracture-fillings in beryl and the adjacent quartz, rarely in fractures cutting perthite and triphylite. Whiteite occurs as yellowish, vitreous to resinous, small ($\ll 4$ mm), prismatic crystals and subparallel groups. Some crystals display a morphology almost identical to Tip Top jahnsite and to jahnsite from the Sapucaia pegmatite in Brazil. Dominant forms include {001}, {100}, {201} and {011}. It is not uncommon to find partially dissolved whiteite crystals displaying a pink core. Some intermediate compositions between jahnsite and whiteite have also been found.

Whitlockite $\text{HCa}_9\text{Mg}(\text{PO}_4)_7$

Whitlockite is rather abundant, especially in recent workings. Crystals are simple rhombs, rarely modified, are transparent to translucent and have a vitreous luster. Visual inspection shows the modified rhombs to be identical to those from the Palermo No. 1 quarry, New Hampshire, as described by Frondel (1941). Crystals range in size from 0.2 to 5 mm. A crystal measuring 1 cm on edge was measured by one of us (TJC). Color varies from colorless, white to gray, yellowish and pink. Inclusions of less than 0.05 mm robertsite crystals sometimes give whitlockite a pale red color. Whitlockite also occurs as granular massive fracture-fillings up to 5 mm in thickness in quartz, beryl and sometimes perthite and triphylite.

Xanthoxenite $\text{Ca}_4\text{Fe}_2^{+3}(\text{PO}_4)_4(\text{OH})_2 \cdot 3\text{H}_2\text{O}$

Xanthoxenite was reported to occur at this locality by Moore (1973).

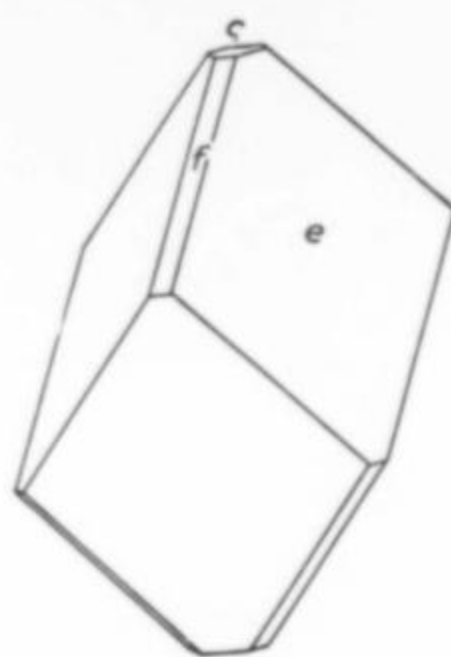


Figure 33. Crystal of whitlockite showing the forms $a\{11\bar{2}0\}$, $e\{01\bar{1}2\}$, $f\{10\bar{1}4\}$ and $p\{10\bar{1}1\}$ (Frondel, 1941).

New Phosphate

A new phosphate mineral recently discovered at the Tip Top pegmatite is presently under investigation by the authors. This mineral is very rare and its occurrence is restricted to dark, olive-green roscherite-bearing assemblages. It occurs as glassy, colorless to pale pink or pale yellow-green, 1-mm, isometric crystals.

PARAGENESIS

Secondary phosphate minerals at the Tip Top pegmatite were formed by percolating meteoric waters at depth and/or by an aqueous residual fluid remaining after core consolidation. The origin of secondary phosphates at this locality is discussed in detail by Campbell (1984). Various stages of phosphate mineral formation at Tip Top are similar to those outlined by Segeler *et al.* (1981) for the Palermo No. 1 pegmatite, and are as follows:

- 1) Crystallization of primary phosphates ($\sim 700^\circ\text{--}600^\circ\text{C}$).
- 2) Metasomatism and initial oxidative reactions ($600^\circ\text{--}300^\circ\text{C}$).
- 3) Hydrothermal attack and remobilization ($300^\circ\text{--}25^\circ\text{C}$).

The first step involved the crystallization of primary phosphate minerals such as triphylite and fluorapatite. These phosphates formed directly from the pegmatite melt or vapor with triphylite largely being restricted to the inner-intermediate zone near the core. Triphylite and other primary phosphates are largely anhydrous and contain no bonded water molecules (H_2O ligands). Bonded water molecules are not stable at temperatures above about 250°C .

During step two, triphylite suffered extensive retrograde oxidative and metasomatic alteration. This involved Li-leaching, oxidation of some Fe and Mn and, to a lesser extent, the addition of Na and Ca. These processes may be the result of the action of hydrothermal fluids which may have carried excess Ca, Mg and Na. Again, these minerals are essentially anhydrous. Granular to cleavable massive minerals are characteristic of metasomatic minerals at the Tip Top pegmatite. Examples include heterosite, ferrisicklerite, alluaudite and some apatite.

The third step involved the hydrothermal alteration of triphylite, muscovite, microcline-perthite, beryl, albite and quartz. This alteration was facilitated by fractures that formed during and after pegmatite consolidation and served as conduits for these hydrothermal fluids. Because temperatures at this stage are below about 250°C , H_2O ligands are stable and become incorporated into the crystal structures of some of the many secondary phosphate species produced (Moore, 1973). Eh-pH conditions probably spanned a rather wide range and the solution chemistry changed temporally as well as spatially in the pegmatite.

Table 3. Mineral associations of secondary phosphates that represent products resulting from hydrothermal attack of silicate minerals and triphylite. Not all of the minerals have been included.

	Carbonate-apatite	Englishite	Eosphorite/Childrenite	Fairfieldite	Fransoletite	Hurlbutite	Hydroxyl-herderite	Mitridatite	Montgomeryite	Robertsite	Roscherite	Tiptopite	Whiteite	Whitlockite	New Phosphate	Ehrleite	Parascholzite
Carbonate-apatite	•	R	R	C	C	C	R	NC	VC	C	NC	C	NC	VC	R	?	*
Englishite	R	•	C	NC	C	C	C	NC	C	C	C	C	NC	C	C	*	*
Eosphorite/Childrenite	R	C	•	NC	*	NC	C	R	NC	R	C	C	R	C	C	*	*
Fairfieldite	C	NC	NC	•	*	NC	NC	NC	VC	C	C	VC	NC	C	R	*	*
Fransoletite	C	C	*	*	•	C	NC	R	VC	C	VC	VC	NC	NC	*	*	*
Hurlbutite	C	C	NC	NC	C	•	*	R	VC	C	VC	VC	R	VC	*	*	*
Hydroxyl-herderite	R	C	C	NC	NC	*	•	R	NC	R	C	C	*	R	NC	C	C
Mitridatite	NC	NC	R	NC	R	R	R	•	NC	C	NC	NC	C	C	R	NC	NC
Montgomeryite	VC	C	NC	VC	VC	VC	NC	NC	•	C	VC	VC	C	VC	C	*	*
Robertsite	C	C	R	C	C	C	R	C	C	•	C	NC	NC	VC	R	*	*
Roscherite	NC	C	C	C	VC	VC	C	NC	VC	C	•	VC	C	C	C	C	C
Tiptopite	C	C	C	VC	VC	VC	C	NC	VC	NC	VC	•	C	C	VC	*	*
Whiteite	NC	NC	R	NC	NC	R	*	C	C	NC	C	C	•	C	R	*	*
Whitlockite	VC	C	C	C	NC	VC	R	C	VC	VC	C	C	C	•	NC	*	*
New Phosphate	R	C	C	R	*	*	NC	R	C	R	C	VC	R	NC	•	*	*
Ehrleite	?	*	*	*	*	*	C	NC	*	*	C	*	*	*	*	•	C
Parascholzite	*	*	*	*	*	*	C	NC	*	*	C	*	*	*	*	C	•

Legend: VC = very common, C = common, NC = not common, R = rare, ? = association uncertain, * = not associated.

The secondary phosphates encountered at the Tip Top pegmatite have two distinct hydrothermal origins which help categorize them into the following two groups:

Group 1: products of direct solution, oxidation, and recrystallization of primary phosphates

Group 2: products resulting from hydrothermal attack on silicate minerals and triphylite. These phosphates usually crystallize along fracture surfaces in the pegmatite.

Group 1 consists of phosphate Suites I and II; Group 2 consists of phosphate Suites III, IV and V.

The paragenetic sequences in Suite I through V are rather complex as manifested by the variety of cations that were present in the system. The following is a discussion of the mineral parageneses of the various suites accompanied by paragenetic diagrams. Because of the complexity and variance of the parageneses, only generalized paragenetic trends are presented here.

Group 1

Group 1 phosphates largely incorporate metals of the first transition series such as Fe and Mn with minor amounts of K, Ca, Mg and Al. The action of hydrothermal solutions on triphylite liberated the phosphate anion along with Fe, Mn and Li. Ca and Mg were brought in by a late hydrothermal fluid that may have represented a residuum from a higher temperature metasomatic fluid. Calcium may also have been released from the dissolution of fluorapatite. Potassium and Al were released by the hydrothermal alteration of muscovite, which often coexists with triphylite. Secondary phosphates formed from these solutions are largely restricted to solution cavities in triphylite and within fractures that transect the adjacent quartz.

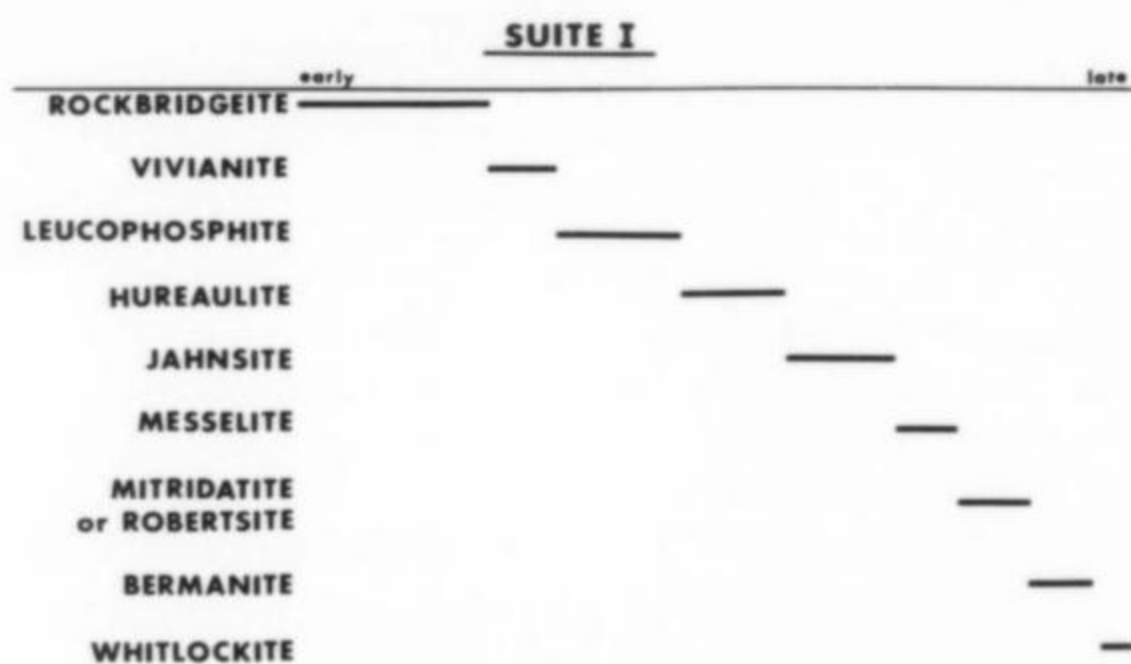


Figure 34. Paragenetic sequence of Suite I minerals.

Phosphate Suite I

Minerals of the Suite I assemblage are dominated by transition metals with subordinate amounts of alkaline earth metals and K as the persistent alkali metal. Minerals represented here in order of decreasing abundance are rockbridgeite, leucophosphite, hureaulite, jahnsite, mitridatite, robertsite, vivianite, messelite, bermanite and whitlockite. Carbonate-hydroxylapatite or hydroxylapatite, collinsite, quartz, tavorite and laueite may also be present. These minerals are found in fractures or solution cavities in triphylite and quartz. Some minerals, such as robertsite, mitridatite and jahnsite, occur as at least two generations, complicating the paragenetic sequences. The following, therefore, is a generalized sequence.

Rockbridgeite and vivianite, without exception, were the earliest minerals formed and were followed closely by leucophosphite. Hureaulite formed after leucophosphite and brown or green

Table 4. Mineral associations of secondary phosphates that chiefly represent products resulting of direct solution, oxidation and recrystallization of triphylite. Not all of the minerals have been included.

	Barbosalite	Beraunite	Bermanite	Collinsite	Dufrenite	Hureaulite	Jahnsite	Kidwellite	Kryzhanovskite	Laeite	Leucophosphate	Ludlamite	Messelite	Mitridatite	Montgomeryite	Phosphoferrite	Robertsite	Rockbridgeite	Segelerite	Switzerite	Tavorite	Tinsleyite	Vivianite	Whitlockite
Barbosalite	*	*	*	*	*	C	C	*	*	*	R	*	*	*	*	*	*	C	*	*	C	*	R	*
Beraunite	*	*	*	R	*	*	C	R	*	*	C	*	R	R	*	*	R	*	*	*	*	*	*	*
Bermanite	*	*	*	*	*	R	NC	*	*	R	R	*	R	NC	*	*	VC	NC	*	*	NC	C	R	R
Collinsite	*	R	*	*	*	*	NC	R	*	R	NC	*	?	NC	*	*	C	C	C	*	*	R	*	R
Dufrenite	*	*	*	*	*	*	R	*	*	*	NC	*	*	*	*	*	*	C	*	*	*	*	*	*
Hureaulite	C	*	R	*	*	*	VC	*	*	*	VC	*	NC	NC	*	*	C	VC	*	VC	C	*	C	C
Jahnsite	C	C	NC	NC	R	VC	*	C	*	NC	VC	*	VC	C	*	*	VC	VC	C	C	C	C	C	C
Kidwellite	*	R	*	R	*	*	C	*	*	*	C	*	*	R	*	*	C	C	*	*	*	*	*	*
Kryzhanovskite	*	*	*	*	*	*	*	*	*	NC	*	C	*	*	*	VC	*	C	*	*	*	*	NC	*
Laeite	*	*	R	R	*	*	NC	*	NC	*	C	NC	*	NC	*	NC	NC	C	R	*	*	C	*	*
Leucophosphate	R	C	R	NC	NC	VC	VC	C	*	C	*	*	C	C	R	*	C	VC	C	C	C	VC	C	C
Ludlamite	*	*	*	*	*	*	*	*	C	NC	*	*	*	*	*	C	*	C	*	*	*	*	NC	*
Messelite	*	R	R	?	*	NC	VC	*	*	*	C	*	*	VC	*	*	C	VC	?	*	*	*	R	R
Mitridatite	*	R	NC	NC	*	NC	C	R	*	NC	C	*	VC	*	R	*	C	C	NC	*	*	NC	NC	NC
Montgomeryite	*	*	*	*	*	*	*	*	*	*	R	*	*	R	*	*	R	R	*	*	*	*	*	R
Phosphoferrite	*	*	*	*	*	*	*	*	VC	NC	*	C	*	*	*	*	*	C	*	*	*	*	NC	*
Robertsite	*	R	VC	C	*	C	VC	C	*	NC	C	*	C	C	R	*	*	C	NC	R	R	VC	NC	C
Rockbridgeite	C	*	NC	C	C	VC	VC	C	C	C	VC	C	VC	C	R	C	C	*	C	NC	C	C	VC	NC
Segelerite	*	*	*	C	*	*	C	*	*	R	C	*	?	NC	*	*	NC	C	*	*	R	R	*	*
Switzerite	*	*	*	*	*	VC	C	*	*	*	C	*	*	*	*	*	R	NC	*	*	*	*	R	*
Tavorite	C	*	NC	*	*	C	C	*	*	*	C	*	*	*	*	*	R	C	R	*	*	NC	R	*
Tinsleyite	*	*	C	R	*	*	C	*	*	C	VC	*	*	NC	*	*	VC	C	R	*	NC	*	*	R
Vivianite	R	*	R	*	*	C	C	*	NC	*	C	NC	R	NC	*	NC	NC	VC	*	R	R	*	*	R
Whitlockite	*	*	R	R	*	C	C	*	*	*	C	*	R	NC	R	*	C	NC	*	*	*	R	R	*

Legend: VC = very common, C = common, NC = not common, R = rare, ? = association uncertain, * = not associated.

jahnsite formed subsequently. White interlocking plates of messelite formed toward the end of, and after, jahnsite deposition. Spherical aggregates and plates of mitridatite or robertsite, the species depending on the amount of Fe or Mn left in the solution, crystallized after messelite. Red-brown, platy aggregates of bermanite are among the last minerals formed, followed by whitlockite. This particular sequence has also been observed by Moore (1973 and 1974).

Phosphate Suite II

Minerals of Suite II are restricted to highly altered and oxidized pods of triphylite in which the muscovite has been altered. This assemblage contains frondelite-rockbridgeite, leucophosphate, robertsite, carbonate-hydroxylapatite, tinsleyite and laeite in approximate order of decreasing abundance.

Brown, slightly oxidized, fibrous masses and isolated sprays of frondelite-rockbridgeite and yellow-green blebs of tavorite are among the earliest phosphates formed; frondelite-rockbridgeite is dominant. Green, prismatic crystals of leucophosphate formed after frondelite-rockbridgeite and tavorite. Tinsleyite formed after leucophosphate and commonly occurs as morphologically continuous overgrowths on leucophosphate. Deep magenta-colored single crystals of tinsleyite, which are morphologically identical to leucophosphate, are also present. Black spherical aggregates of robertsite formed near the end of and subsequent to leucophosphate and tinsleyite deposition. Yellow prismatic crystals of a jahnsite-group mineral or crystals of yellow-green jahnsite formed during and after robertsite. Segelerite, where present, formed after the jahnsite-

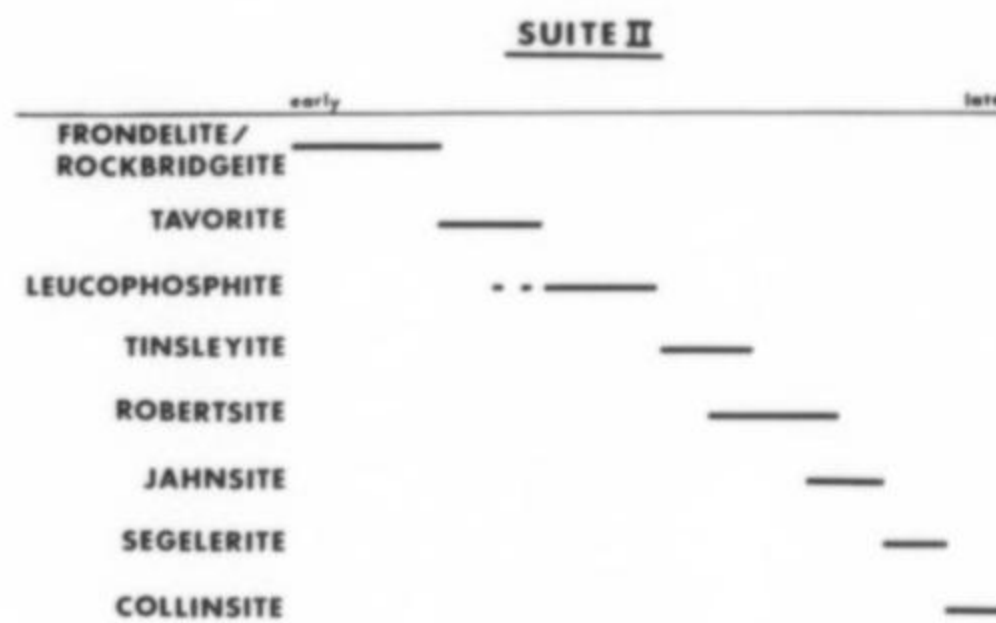


Figure 35. Paragenetic sequence of Suite II minerals.

group minerals and was followed by bladed, rosettes of collinsite; robertsite, and possibly some mitridatite may be intimately associated with these two minerals as well. Other minerals paragenetically younger than the previously described phosphates include white, radial sprays of carbonate-hydroxylapatite and orange, prismatic crystals of laeite.

Group 2

Reactions resulting in the formation of Group 2 phosphates involve the attack of phosphate-rich hydrothermal fluids on primary silicate minerals such as perthite, beryl and albite. The solutions probably came in contact with triphylite which supplied the phos-

phate anion along with Fe, Mn and Li. Subsequent attack on the aforementioned silicates by solutions supplied K, Al, Be, Ca and Na, and upon crystallization the phosphates proceeded to fill fractures in these primary silicate minerals.

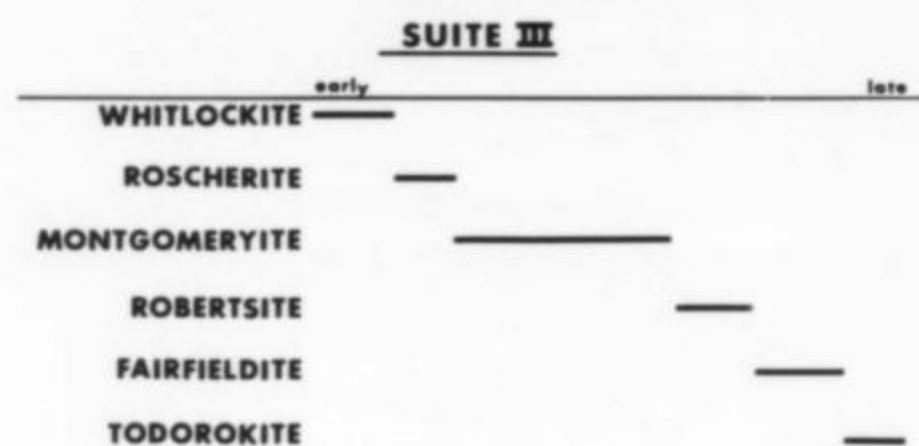


Figure 36. Paragenetic sequence of Suite III minerals.

Phosphate Suite III

Suite III is characterized by whitlockite, carbonate-hydroxylapatite, red to pink montgomeryite, pyrolusite, robertsite, lavender roscherite, fairfieldite and todorokite, in order of decreasing abundance. These minerals are found lining fractures in quartz, perthite, beryl and albite, rarely in triphylite.

In most cases, the fractures are coated with a crust of carbonate-hydroxylapatite and/or a microcrystalline druse of colorless to pink whitlockite. Some specimens show continued whitlockite crystallization through the entire sequence. Tabular crystals and small clusters of lavender roscherite were then deposited, followed by fascicles of red to pink montgomeryite. Deep red to almost black, prismatic, sometimes twinned, crystals of robertsite formed during and after montgomeryite. Plates of fairfieldite crystallized after robertsite and black, submetallic, 0.5-mm, radial tufts of pyrolusite formed subsequent to fairfieldite. Gray-black, submetallic, spherical aggregates of todorokite were the last mineral to form. This sequence is slightly modified from the one outlined in Dunn *et al.* (1983).

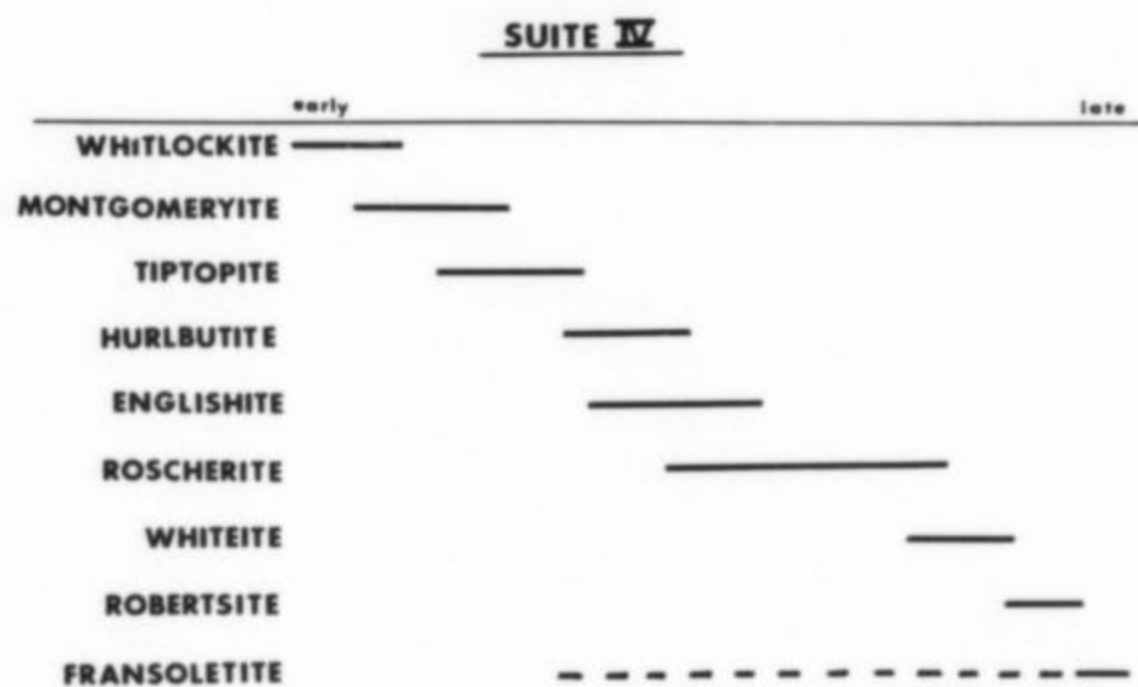


Figure 37. Generalized paragenetic sequence of Suite IV minerals.

Phosphate Suite IV

Suite IV is similar to Suite III but is dominated by alkali metal-bearing and Be-bearing phosphates with subordinate amounts of alkaline earth and transition metal phosphates. This suite of phosphates represents a chemically very diverse assemblage of minerals. These phosphates also prove to be exceedingly colorful and very esthetic. The colors of montgomeryite and roscherite characterize this suite; montgomeryite in shades of yellow and pink with roscherite which varies from orange-red to red-brown. In addition to montgomeryite and roscherite, this assemblage includes

whitlockite, tiptopite, englishite, carbonate-hydroxylapatite, hurlbutite, robertsite, whiteite, todorokite, fairfieldite and fransoletite, in order of decreasing abundance. Minerals of this assemblage occur as fracture fillings in beryl and quartz. Many of the fractures are lined with a druse of 0.2-mm whitlockite crystals while others are lined with white to pale green, botryoidal crusts of carbonate-hydroxylapatite. In general, whitlockite is the dominant early phase. Pink to red clusters of montgomeryite formed contemporaneously with whitlockite and are frequently covered by subsequent whitlockite crystallization (Dunn *et al.*, 1983). This portion of the sequence may be related to Suite III mineralization. Yellow to yellow-brown fascicles of montgomeryite formed toward the end of whitlockite crystallization. Tiptopite formed just after the crystallization of yellow montgomeryite; then pale pink to bluish white spherules of hurlbutite and pearly white hemispherules of englishite, which appear to be contemporaneous, were formed. Spherulitic aggregates of orange-red to red-brown roscherite then formed during and after hurlbutite and englishite formation. Deep red to red-black clusters and single crystals of robertsite or robertsite-mitridatite crystallized just before or during roscherite formation, and whiteite formed subsequent to roscherite. Fairfieldite, which is not abundant in this assemblage, crystallized after whiteite. Fransoletite, generally present in the absence of fairfieldite, was the last phosphate to form, although a few specimens show it as an earlier formed phase. Gray-black, submetallic, spherical aggregates and crusts of todorokite formed last.

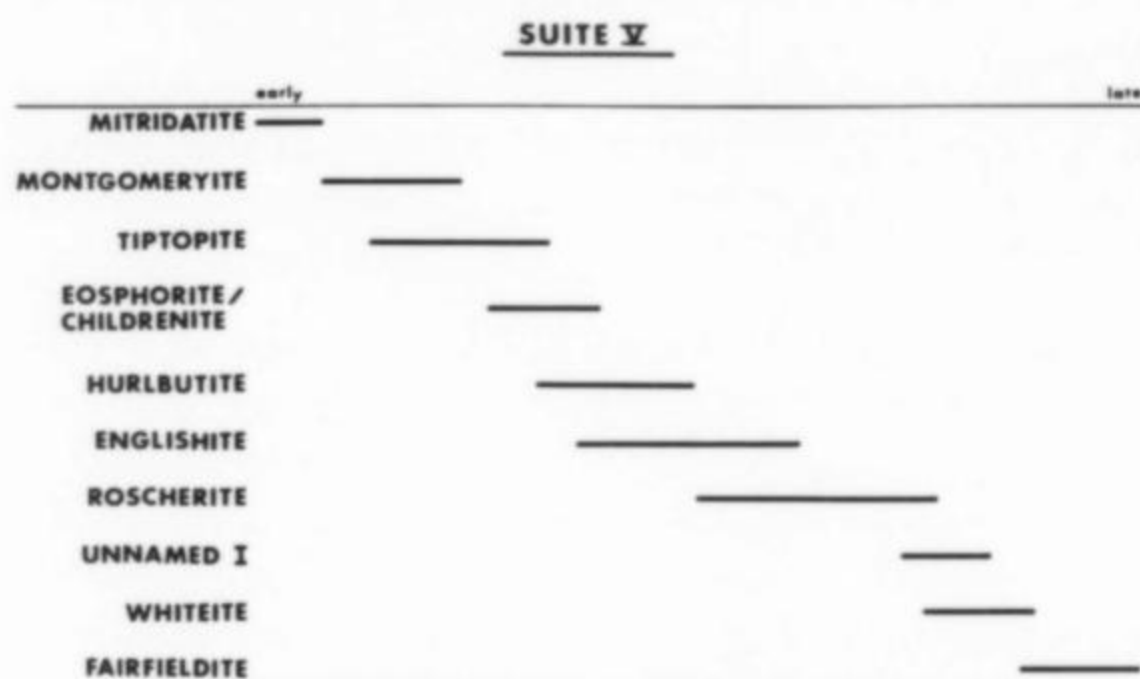


Figure 38. Paragenetic sequence of Suite V minerals.

Phosphate Suite V

Suite V consists of dark, olive-green roscherite, tan to colorless montgomeryite, fairfieldite, tiptopite, englishite, whitlockite, mitridatite, eosphorite-childrenite, hurlbutite, whiteite and the undescribed phosphate. This suite of phosphates is mineralogically similar, but distinct, compared to Suite IV, as described by Grice *et al.* (1985). An obvious difference between the two assemblages is the overall color, as exemplified by olive-green roscherite and colorless to tan montgomeryite. Another difference is that this suite contains mitridatite and eosphorite-childrenite, lacks fransoletite and does not have whitlockite as a dominant early phase. This sequence is restricted to fracture surfaces in beryl.

Dark red-brown, lamellar aggregates of mitridatite formed a crust on which montgomeryite and later-formed phosphates were deposited. After the onset of montgomeryite crystallization, sprays of tiptopite formed; the two minerals are commonly found intergrown. Montgomeryite and tiptopite also crystallized directly on beryl fracture surfaces. During and after the deposition of montgomeryite and tiptopite, eosphorite-childrenite crystals formed, followed by pale-brown spherules of hurlbutite. Toward the end of hurlbutite formation, spherulitic aggregates of englishite formed and are sometimes impaled on acicular tiptopite crystals. Roscher-

ite was deposited contemporaneously with, and was partially covered by, subsequent englishite crystallization. Crystals of the new phosphate formed with englishite. Whiteite, if present, formed after roscherite. Stacked, parallel aggregates of fairfieldite were the last phase formed and whitlockite, where present, occurs as a druse on the other minerals.

SUMMARY

The Tip Top pegmatite is the type locality for eight phosphate mineral species. These are jahnsite, robertsite, segelerite, fran-soletite, tinsleyite, tiptopite, ehrleite and a soon-to-be-described phosphate; the last five are unique to this deposit. This mine has also produced some of the most esthetic specimens of secondary phosphate minerals ever encountered in a pegmatite, with crystals ranging from sub-millimeter size up to tens of millimeters for some species. The majority of these minerals can still be collected on the dumps.

The Tip Top mine easily parallels Palermo in diversity of phosphate species and equals or surpasses it in terms of complexity of paragenesis. Through continued research, the number of phosphate species from the Tip Top pegmatite may surpass that from Palermo.

Secondary phosphates at the Tip Top formed during a moderate to low temperature (300°–25°C) hydrothermal alteration and remobilization stage. All of these secondary phosphates formed below 250°C; above this temperature H₂O ligands, which play an essential role in secondary phosphate formation, are not stable (Moore, 1973). These phosphates can be subdivided into two groups: 1) products of direct solution, oxidation and recrystallization of primary phosphates, and 2) products of the hydrothermal alteration of silicate minerals and triphylite. The majority of the metasomatic and secondary phosphate minerals are produced largely by the hydrothermal attack of primary minerals by percolating meteoric waters at depth and/or by an aqueous residual liquid remaining after core consolidation. Future work using light stable isotopes and fluid inclusion studies may give some insights to the source of the fluids and the origin of secondary phosphates at this locality.

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Known at least since the 1870s, the Stoneham barite locality has long been considered among the country's best barite occurrences. Superb blue crystals ranking with the best ever found were collected in May of last year.

LOCATION

The Stoneham deposit is located about 1 mile east and 3 miles north of the town of Stoneham, Weld County, Colorado (R56W, T8N, Sec. 16). Access is by paved road except for the final 3 miles, which are a well-maintained dirt road. The principal mineralized area covers about 20 acres, with the best collecting centered along the western edge. The dirt road passes directly in front of the property owner's home, where permission to collect must be obtained.

Early references commonly refer to the locality as Sterling, Colorado, but the town of Sterling is about 26 miles away, in Logan County. All such labels should be corrected to read Stoneham, Weld County.

HISTORY

The locality was apparently discovered in the middle 1870s; it is not mentioned in the fifth edition of Dana's *System* (1874), but was noted by Endlich (1878) in his annual report for the year 1876. Thorough descriptions have not been published, probably because the occurrence has never been commercially mined and has remained of significance to collectors only.

Since that time, the locality has remained famous as a source of fine blue barite crystals (Dana, 1892; Ellermeier, 1948; Eckel, 1961; Pearl, 1972). Over the past 25 years, it has been periodically closed to collectors, but the owners are currently granting permission provided that all workings are filled in before collectors leave the site.

In May of 1985, I visited the Stoneham locality with long-time Stoneham collectors Luke and Cathy Westervelt. During the course of 8 days, a major pocket of specimens was discovered and removed. Several of the best pieces are pictured here.

GEOLOGY

The Stoneham region consists predominantly of rocks of the Oligocene-age White River formation. Calcium-rich montmorillonite clays of the lower part of the formation are altered volcanic ash, and occur interbedded with barite-bearing shale, coarse sediments, and lenses of fresh-water limestone. Recent erosion has produced a badlands topography.

Barite was probably deposited by circulating groundwater and associated alteration of the volcanic ash. The blue color, commonly compared to that of aquamarine, is thought to be due to natural irradiation by trace elements.

Barite is found as loose crystals on the hill slopes, as cavity linings and as vein or fracture linings. Small, loose singles and fragments can be collected by the hundreds on the hillsides.

Along the western flank of the mineralized area, the silty white to tan clay layer which contains barite is overlain by 1 to 2 meters of sand, gravel and cobbles. Within the clay layer are calcite-barite seams following a joint system which dips at 15° to 45°. Pockets are found where the joints are most horizontal. These pockets are generally very tight, usually resulting in crystals which contact both the top and bottom faces of the seam. Careful collecting is therefore necessary. Pocket size reaches 2 to 5 meters.



MRL

Figure 1. Stoneham barite crystal illustrated in R. Braun's *Das Mineralreich* (1903).



Figure 2. Excellent single crystal of blue barite, 2.3 cm wide, partially removed from an overgrowth of calcite. This and the other four specimens pictured here are from the 1985 find.



Figure 3. Lustrous, well-formed group of barite crystals 6 cm long.



Figure 4. Large, tabular crystal 6.8 cm long, perched on another crystal 8 cm long.

BARITE

The most common barite crystal habit at Stoneham is moderately tabular on c with a rectangular outline elongated along b . Prism-like habits elongated on b are not uncommon among smaller crystals, but are rare in larger specimens. Common forms include $c\{001\}$, $d\{101\}$, $o\{011\}$, $m\{210\}$ and $b\{010\}$.

Color ranges from nearly colorless to gray-blue and aquamarine-blue; larger crystals appear to be a deeper blue due to their thickness, and are less transparent than smaller crystals. Crystals to 16.5 cm (6½ inches) were recovered in the recent find; earlier

references (e.g., Eckel, 1961, and Pearl, 1972) mention crystals to a maximum of 4 to 6 inches. Sizes of 1 to 2 inches are most common. Luster ranges from dull to brilliant.

Barite pockets at Stoneham also have been found rarely to contain thick druses of a pale yellow calcite. These druses may cover barite crystals partially or completely and must be carefully picked or dissolved off to reveal the crystals underneath. In other cases, barite has been found uncoated, or perched on an earlier layer of calcite.



Figure 5. Superb barite crystal group 11.5 cm (4.5 inches) long.



Figure 6. Barite group, 9 cm tall, partially coated with pale yellow calcite crystals.

Barite "roses" or rosettes also occur in the area, generally on the upper slopes of the bluffs a short distance below the top of the escarpment (Cook, 1952; Ellermeier, 1948). The rosettes are the same blue color as the single crystals commonly found but tend to be opaque. Opal, which fluoresces green, and delicate crystals of golden barite (Pearl, 1972), and yellow-green barite (Eklund, 1965) have also been found.

The Mineralogical Record, volume 17, July-August, 1986

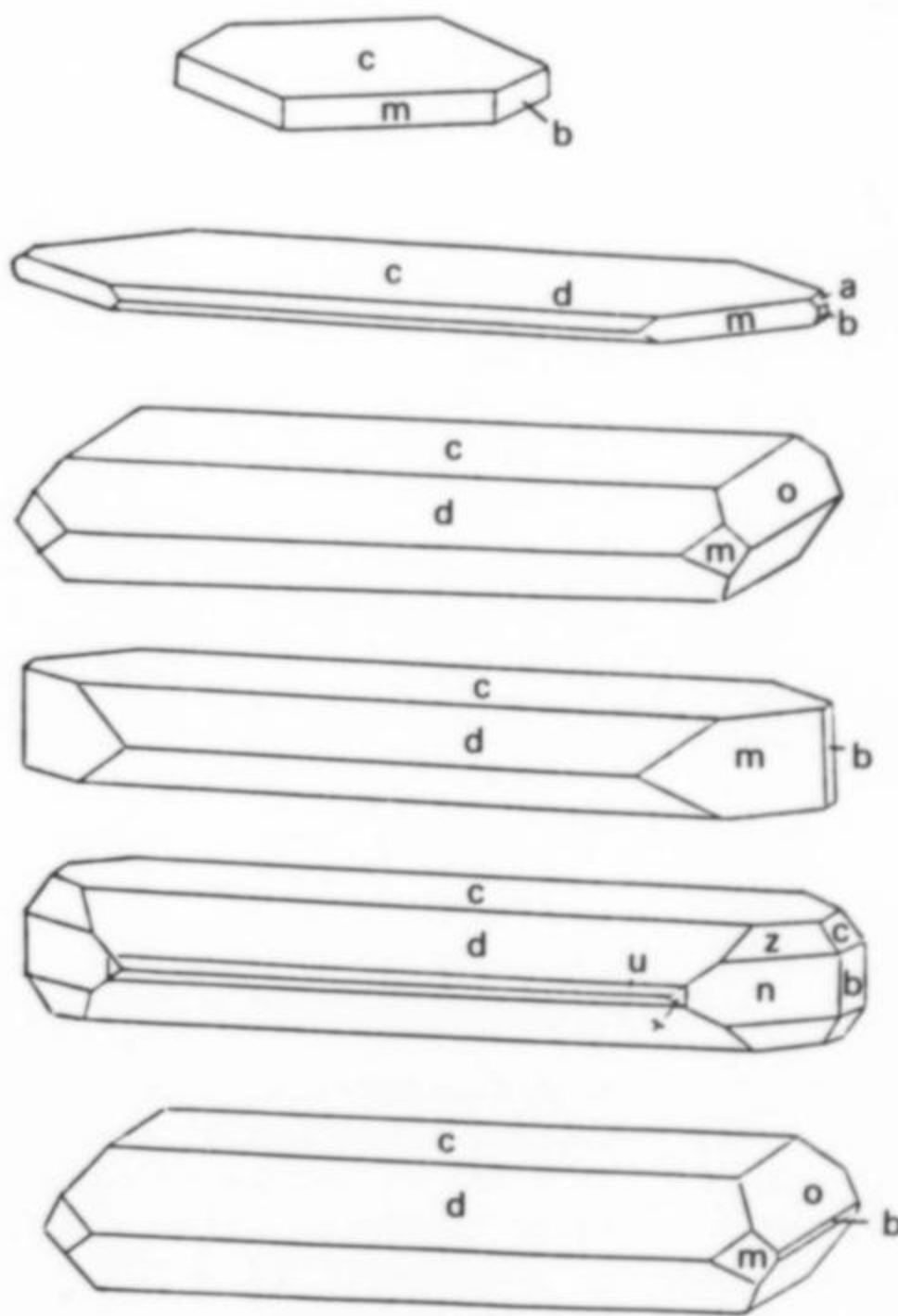


Figure 7. Barite crystal habits from Stoneham (Workman, 1964).



Figure 8. Two pockets in the calcite seam; a pinched out zone separates them.



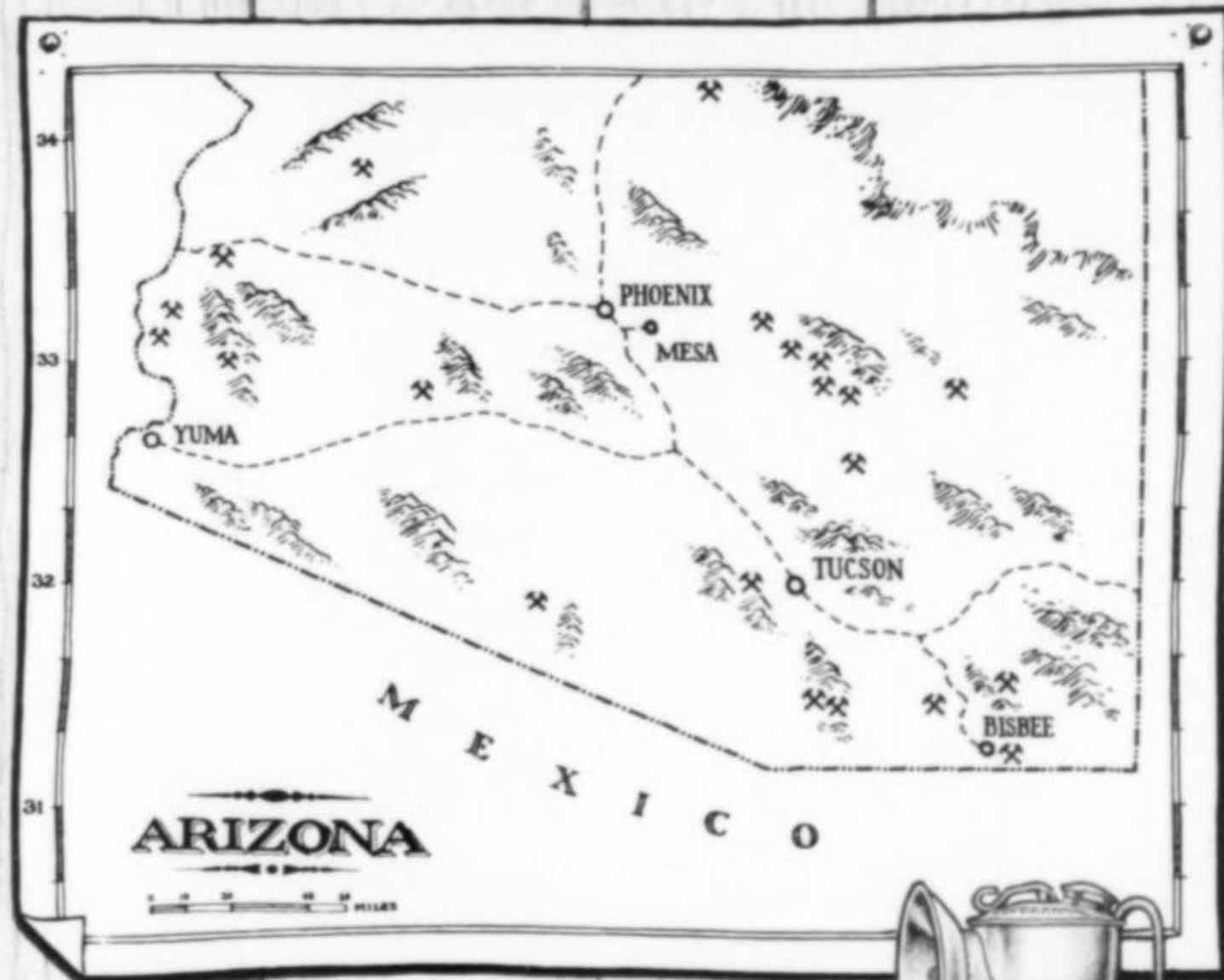
Figure 9. Excavations in the main pocket area.

CONCLUSION

Stoneham has long been a heavily collected area. Nevertheless, the barite mineralization is sufficiently extensive that much virgin ground remains. Significant finds will probably continue to be made for many years, provided that collectors maintain a good relationship with the property owner.

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The Buca della Vena mine has recently become well-known among Italian mineralogists and collectors as the type locality for apuanite and versiliaite. Several other rare and interesting species are found there as well.

INTRODUCTION

This article arose from a common interest the authors have shared in the Buca della Vena mine over a number of years. In 1976 F. Checchi, at that time a final-year student in geology, was the first to bring samples from the mine dump to the University of Pisa for identification. These had been collected as part of his thesis research dealing with geological prospecting. It was for P. Orlandi, then Curator of Minerals, the first contact with Buca della Vena minerals and the beginning of a series of studies (with S. Merlino and M. Mellini). As a result, two new species, apuanite and versiliaite, were characterized (Mellini *et al.*, 1979), and several other very rare minerals were identified including derbylite and stibivanite.

The Buca della Vena mine (formerly known only as a small iron and barite deposit) has since become a popular collecting site among local mineralogists and mineral collectors.

LOCATION

The mine is situated on the north slope of Mount Stazzema in Tuscany, Italy. Access is from the town of Pietrasanta on the Pietrasanta-Castelnuovo Garfagnana Road. Four kilometers past the village of Seravezza is a junction, in one direction leading to Stazzema-Ponte Stazzemesse; from this village, it is 1 km on the road to Cardoso, then a right turn onto a small road closed by a gate. The mine is a 20-minute hike from the gate.

HISTORY

The iron deposit was worked intermittently from Roman times up until the eighteenth century, when B. Paci reopened it on a more

continuous basis. From that time until World War II, the iron produced was used locally by various craftsmen and blacksmiths. In 1948, the Society Nuovo Pignone began a more commercial exploitation of the iron. In 1960 the Sima society began quarrying barite there for use in oil well drilling. Recent production from the mine has been used extensively in the construction of the Enel nuclear power plant. Fifteen miners are currently employed at the mine.

GEOLOGY

The Buca della Vena mine is located in the southern part of the Apuan Alps tectonic "window"; here the metamorphic units of the Apuan massif include rocks dated from Paleozoic to Oligocene, which crop out below the unmetamorphosed Tuscan Nappe and Liguride sequences.

The metamorphic basement, containing pre-Alpine deformation fabrics, is represented by phyllites and lenses of Silurian black shale and dolomite. This basement is unconformably overlain by the Permo-Triassic "verrucano" consisting of a volcanoclastic sequence, quartzitic sandstone and conglomerate, followed by a Mesozoic calcareous sequence. This latter sequence includes the Marmo formation wherein are located the famous Carrara marble quarries. A deposit of Oligocene flysch concludes the Alpine sediments.

In the metamorphic sequence two tectonic units, the Autoctone and the Massa, have been recognized (Carmignani *et al.*, 1978). The Autoctone unit is the most important, comprising the core of the Apuan massif. It is considered to be the lowermost outcrop (i.e., "window") in the entire northern Apennine Mountain belt. The

MAP OF MINES

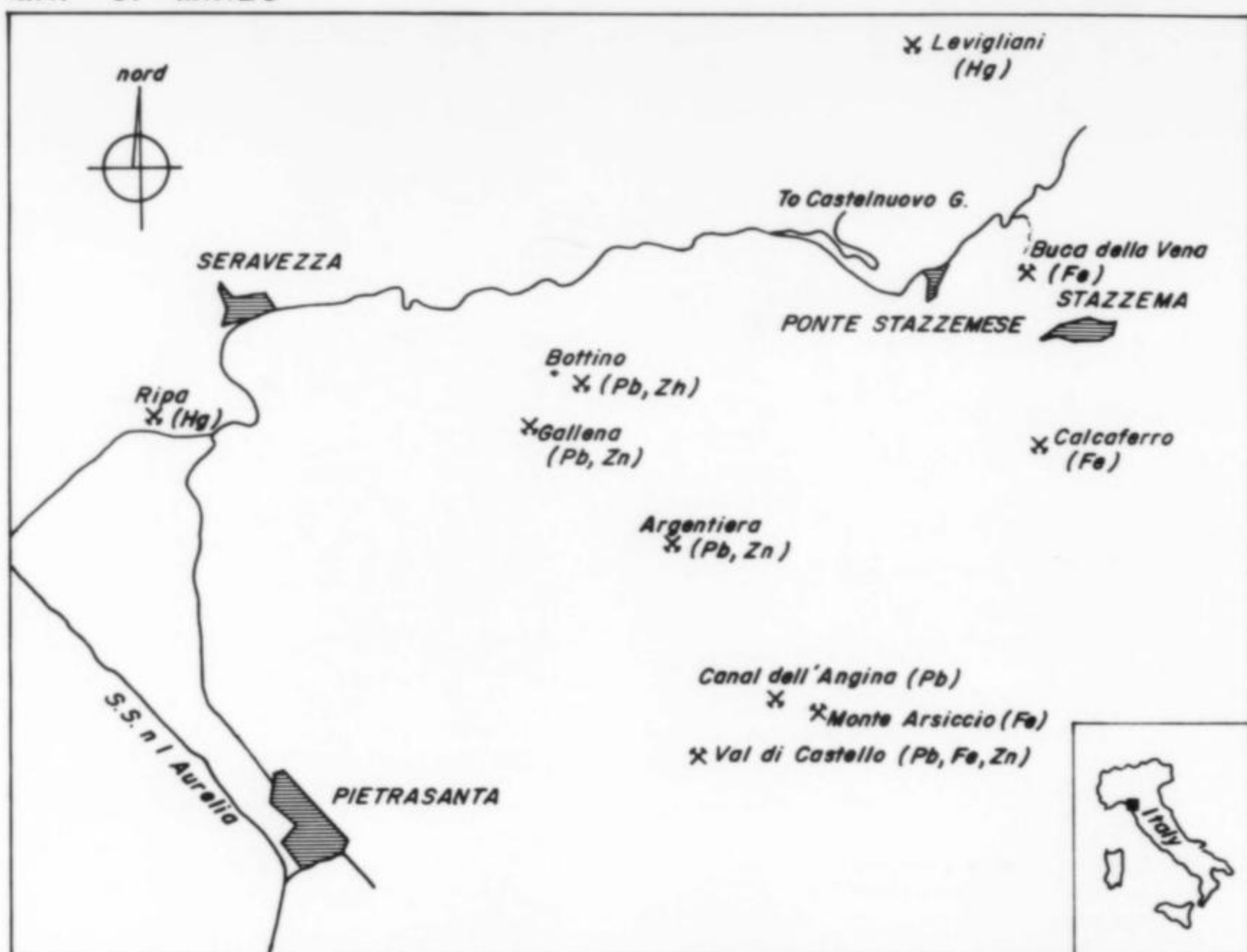


Figure 1. Location map showing the Buca della Vena and surrounding mines, Apuan Alps, northern Tuscany, Italy.

Figure 2. Miners employed at Buca della Vena. All photos by P. Orlandi.



Massa unit, originally a shallow marine sequence of middle and upper Triassic age, crops out in the western area of the window between the Autoctone rocks and the Tuscan Nappe.

In the southern part of the Apuan massif, in the Stazzema area, a number of "tectonic wedges" crop out, and in these wedges are located many of the most important ore deposits of the Apuan Alps. Unfortunately, the geological nature of these wedges has yet to be fully clarified. For a more thorough treatment of the problem, the reader is referred to the studies by Carmignani *et al.* (1978) and Ciarapica and Passeri (1982).

The ore deposit at Buca della Vena is located in the lowermost wedge formed by two carbonate lenses (dolomite and marble) interbedded with lenses of phyllite. Its top is cut by a tectonic breccia. Mineralization consists of a fine-grained mixture of barite and iron

oxides, except at the contact between phyllite and iron oxides where replacement by pyrite has taken place. Carmignani *et al.* (1976) and Checchi (1978) considered the deposit to be a metasomatic replacement of carbonate rocks. The lower of the two calcareous lenses, almost entirely replaced by barite, iron oxides and pyrite, has produced most of the ore. The upper lens is less extensively replaced and has been worked only rarely. Spheroidal relics of unaltered carbonate rock, from a few centimeters to several meters in diameter, are encountered frequently. Small amounts of granular pyrite have been found in the phyllite but have never been mined.

Recently, Cortecchi *et al.* (1984), on the basis of sulfur isotope analysis of barite and sulfides, concluded that the deposit is sedimentary in origin.

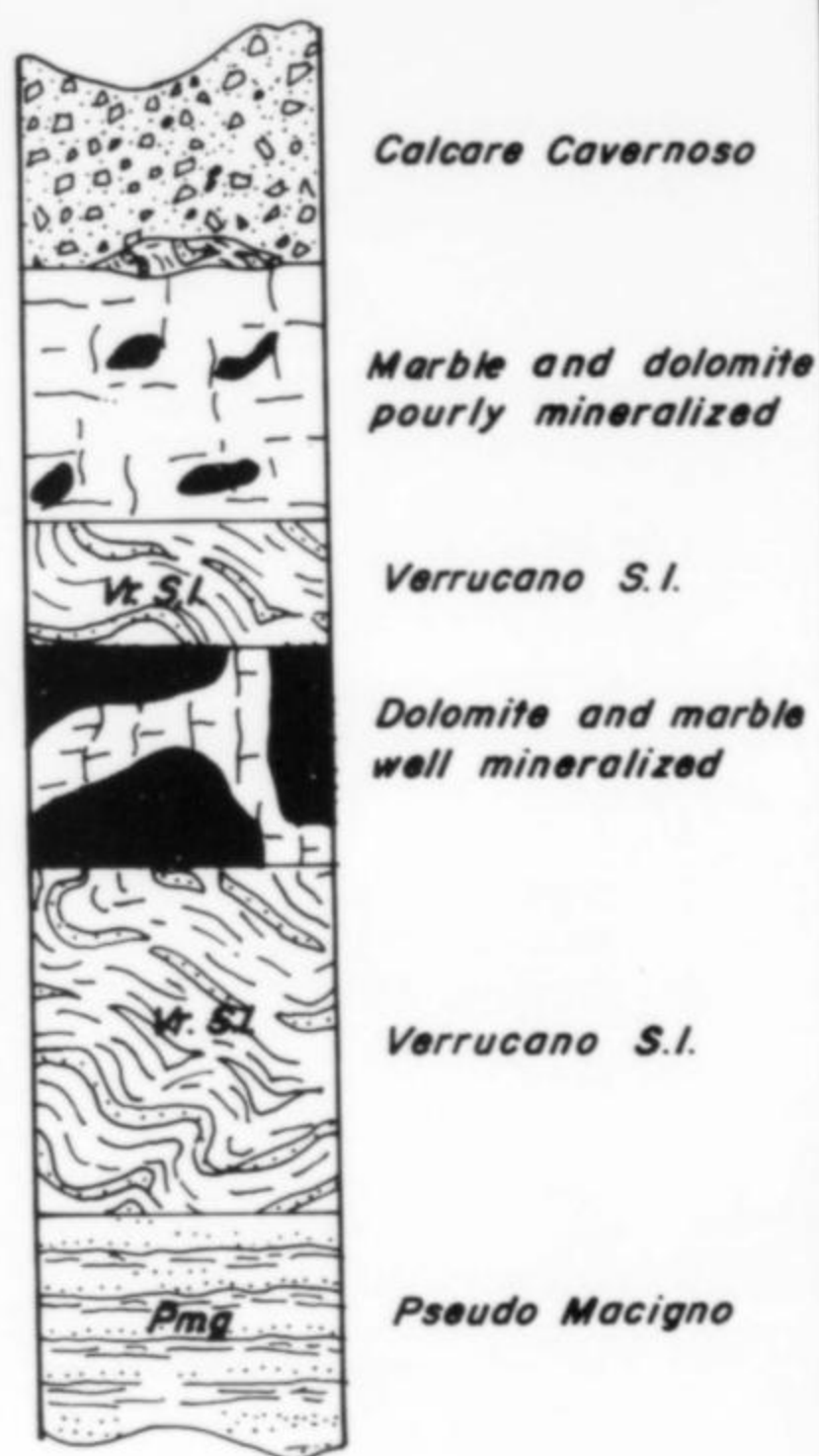


Figure 3. Relationships between the rock units of the mine area.

GEOLOGICAL MAP from Checchi 1978

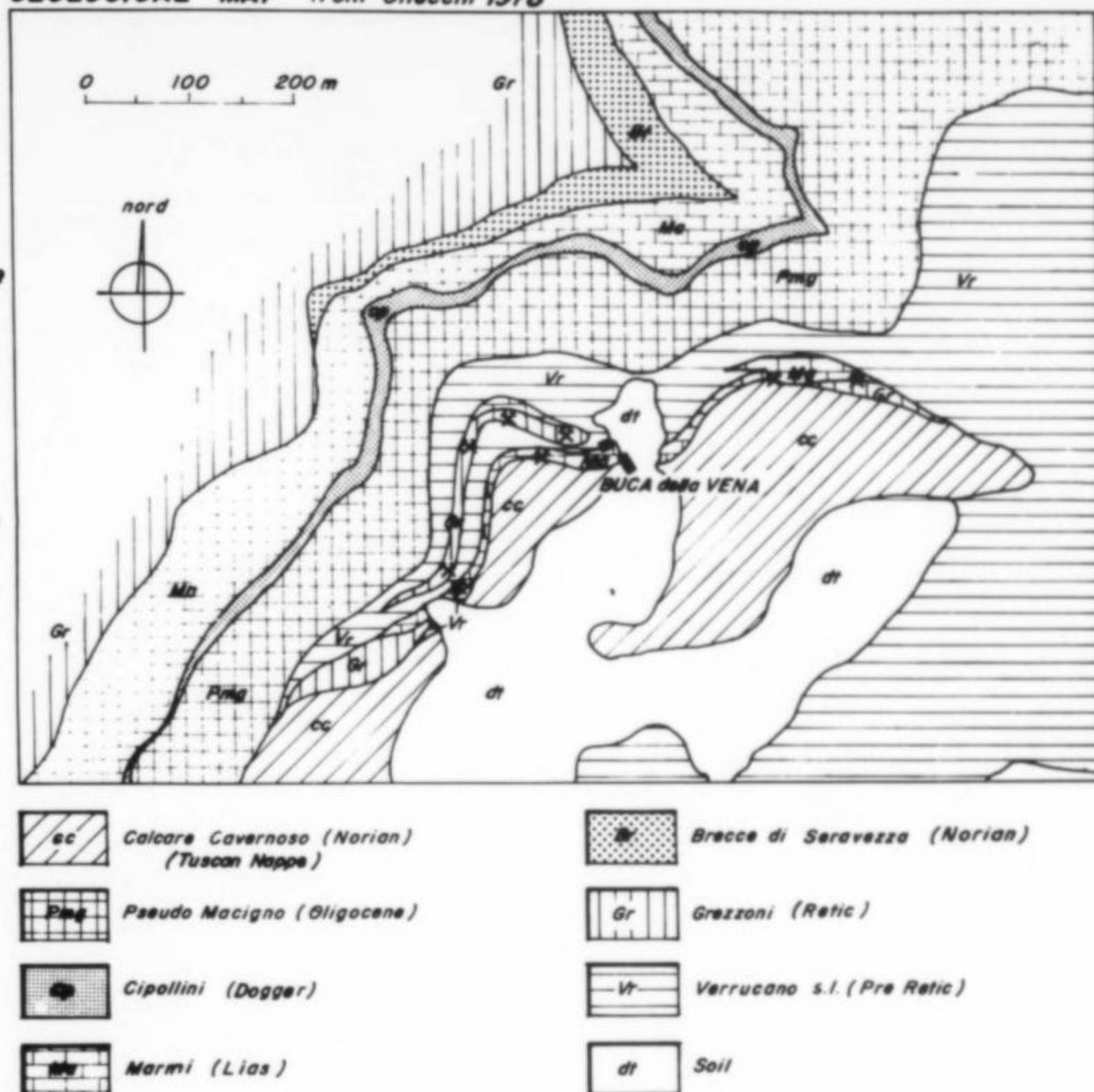


Figure 4. Geologic map (from Checchi, 1978).



Figure 5. Entrance to the mine.

MINERALS

The minerals listed below were identified by X-ray analysis (Burger, Weissenberg and Gandolfi methods) and optical study. In addition, qualitative chemical analyses by X-ray fluorescence and by SEM/EDS have been employed where necessary.

The Mineralogical Record, volume 17, July-August, 1986

Albite $\text{NaAlSi}_3\text{O}_8$

Albite occurs within the dolomite matrix but close to the contact with ore.

Allanite $(\text{Ce}, \text{Ca}, \text{Y})_2(\text{Al}, \text{Fe}^{+3})_3(\text{SiO}_4)_3(\text{OH})$

Allanite occurs as well-developed, prismatic, black crystals rarely exceeding 2 mm in narrow barite veins.

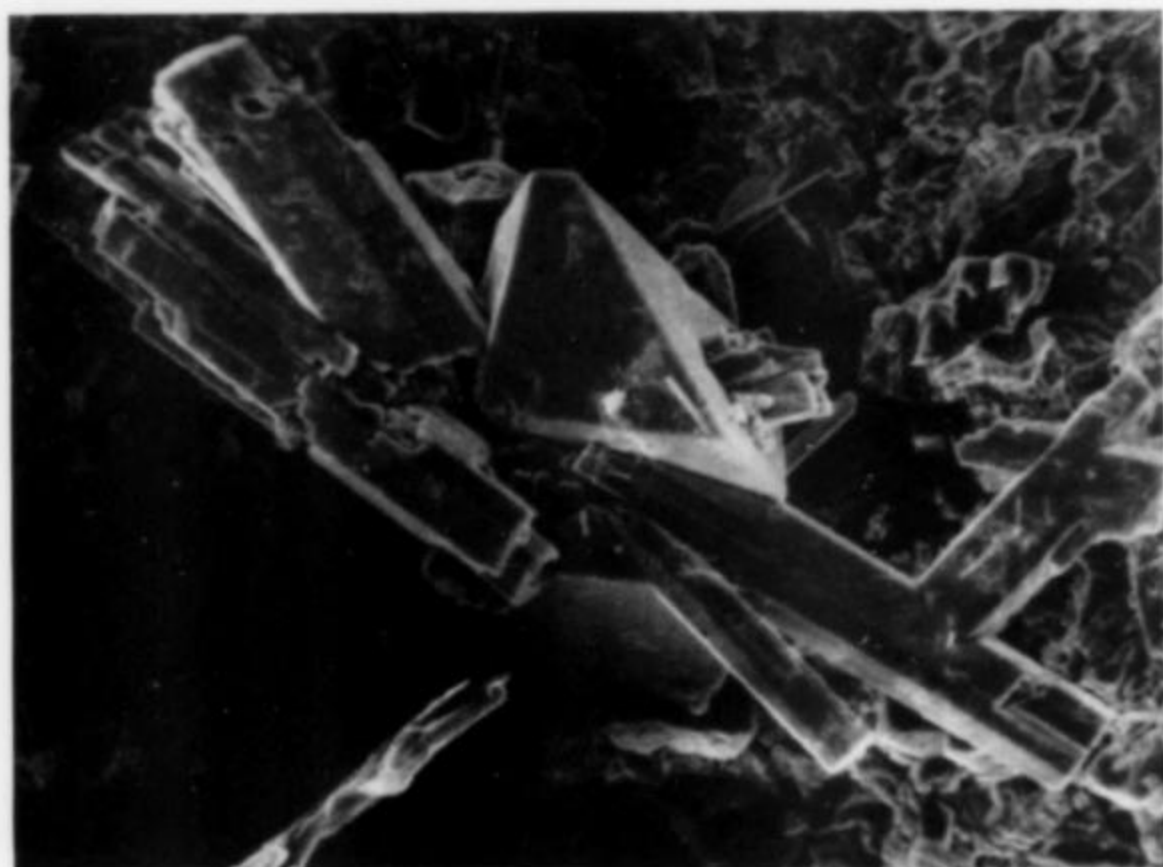


Figure 6. Prismatic, vanadium-rich derbylite with dipyramidal anatase (SEM photo, 400 X).

Anatase TiO_2

Two habits of anatase occur at Buca della Vena. One is characterized by the form {111}, in small (1-2 mm) brown crystals in cavities in dolomite. The other is characterized by {113} in black crystals in small cavities associated with vanadium-rich derbylite, stibivanite and vanadinite.

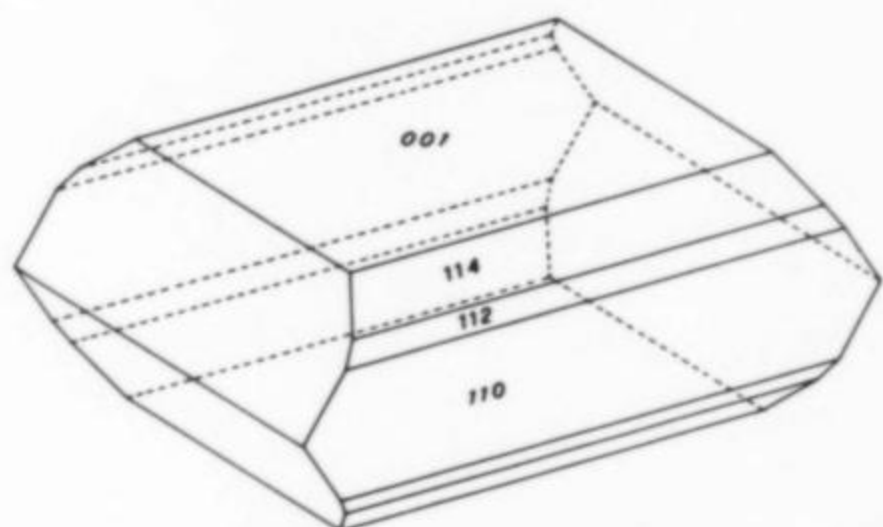


Figure 7. Apuanite crystal, idealized (Mellini *et al.*, 1979).

Apuanite $\text{Fe}^{+2}\text{Fe}^{+3}\text{Sb}_4^{+3}\text{O}_{12}\text{S}$

At present, the Buca della Vena mine is the only known occurrence of apuanite. Apuanite and versiliaite were described as new species from there by Mellini *et al.* in 1979. The mineral occurs disseminated as aggregates of crude microcrystals filling fractures and cavities in dolomite. Some of these aggregates exceed 5 cm in diameter. Single crystals are black with a metallic luster; they are prismatic in habit, consisting of the forms {001}, {110}, {112} and {114}. A few very rare crystals have been found to exceed 1 cm in size. Apuanite is more abundant than either versiliaite or schafarikite.

Barite BaSO_4

Barite is a principal constituent of the mineralization. It occurs as microcrystalline to saccharoidal masses and rarely as distinct, tabular, secondary crystals in dolomite cavities and veins.

Beryl $\text{Be}_3\text{Al}_2\text{Si}_6\text{O}_{18}$

Gemmy, prismatic, emerald-green crystals up to 2 cm in size have been found in veins in barite and dolomite. Single crystals and vein fillings have also been found near the (verrucano) schist. Associations commonly include pyrite and tetrahedrite.



Figure 8. Intergrown crystals of apuanite (SEM photo, 50 X).



Figure 9. Aquamarine beryl in prismatic crystals with quartz and hematite. Field of view: 5 mm. Marco Baldi collection.

Blue-green crystals, perfectly transparent and measuring 3 mm long by 1-2 mm in diameter, occur in small cavities in massive magnetite-hematite.

Bournonite PbCuSbS_3
Seligmannite PbCuAsS_3

Crystals in the bournonite-seligmannite series occur as prisms to 2 cm, in some cases covered by a green malachite crust. Irregular sphalerite-rich nodules have also been found.



Figure 10. Prismatic emerald-green beryl with pyrite in a barite vein. Field of view: 3 cm.

Calcite CaCO_3

Milky rhombohedral to scalenohedral crystals of calcite occur in cavities in dolomite.

Chalcostibite CuSbS_2

Only a single, elongated, tabular, millimeter-size crystal of chalcostibite has been found, in a small cavity in a dolomite vein.

Cinnabar HgS

Cinnabar is very rare at Buca della Vena, usually associated with tetrahedrite in veins of secondary dolomite.

Chlorite group

Platy crystals of chlorite occur in dolomite and as vermicular aggregates of thin, green, tabular crystals in cavities.

Derbylite $\text{Fe}_4^{+3}\text{Ti}_3\text{Sb}^{+3}\text{O}_{13}(\text{OH})$

Derbylite is very rare at Buca della Vena (Mellini *et al.*, 1983). It occurs as lustrous, metallic, black crystals with striated faces, in secondary barite-dolomite veins. Crystals reach 2 mm in size; trilling twins are common.

Recently, vanadium-rich and arsenic-rich varieties of derbylite have been found. The vanadium-rich crystals are acicular-prismatic, very small ($\ll 1$ mm), and always associated with black pyramidal crystals of anatase in cavities lined with black dolomite crystals.

Dolomite $\text{CaMg}(\text{CO}_3)_2$

Clear, rhombohedral dolomite crystals 2–3 cm on an edge occur in veins and cavities in dolomitic rocks.

Hematite Fe_2O_3

Hematite is a major constituent of the replacement bodies. It occurs as microcrystalline lamellae and large (to 1 cm on edge) rhombohedral crystals in secondary veins in dolomite.

Magnetite $\text{Fe}^{+2}\text{Fe}^{+3}\text{O}_4$

Magnetite occurs as granular, microcrystalline lenses and irregular masses chaotically associated with irregular blocks of dolomite. Single crystals have been found in secondary veins in dolomite and barite.

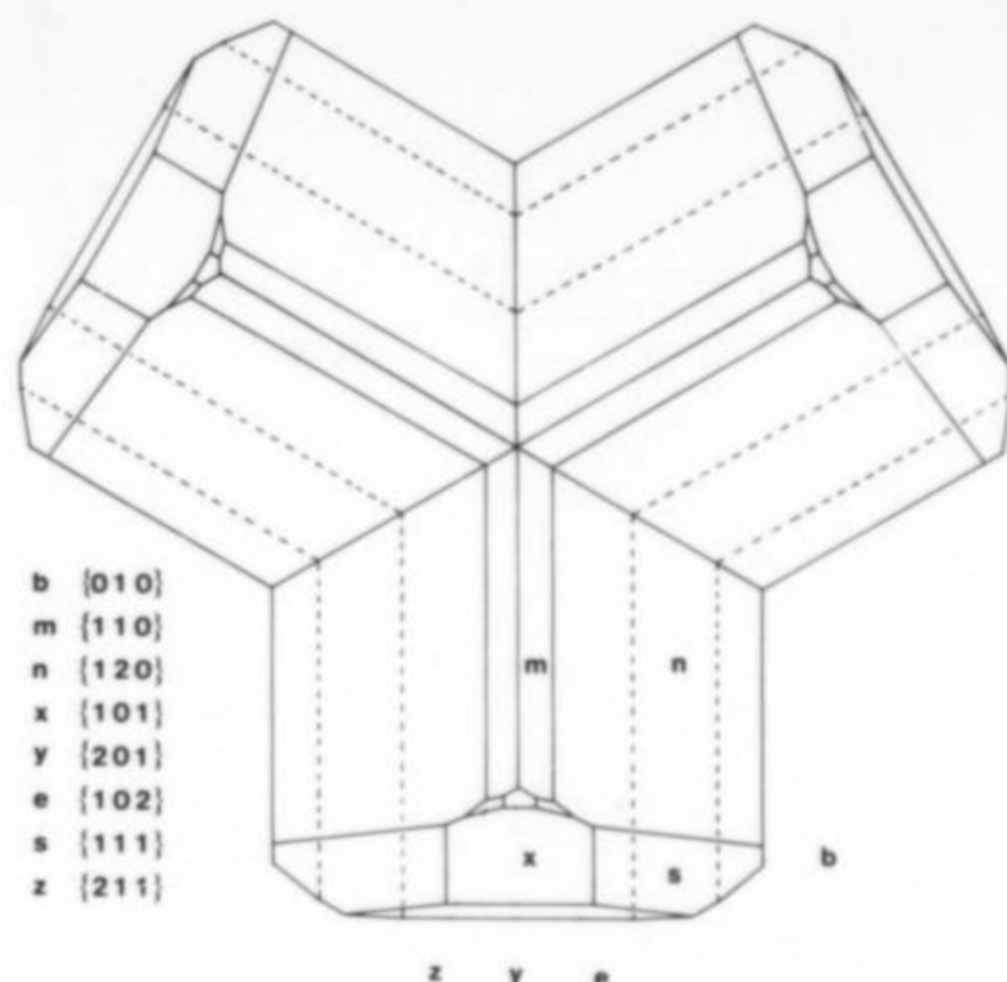


Figure 11. Derbylite: idealized drawing showing typical trilling twin.

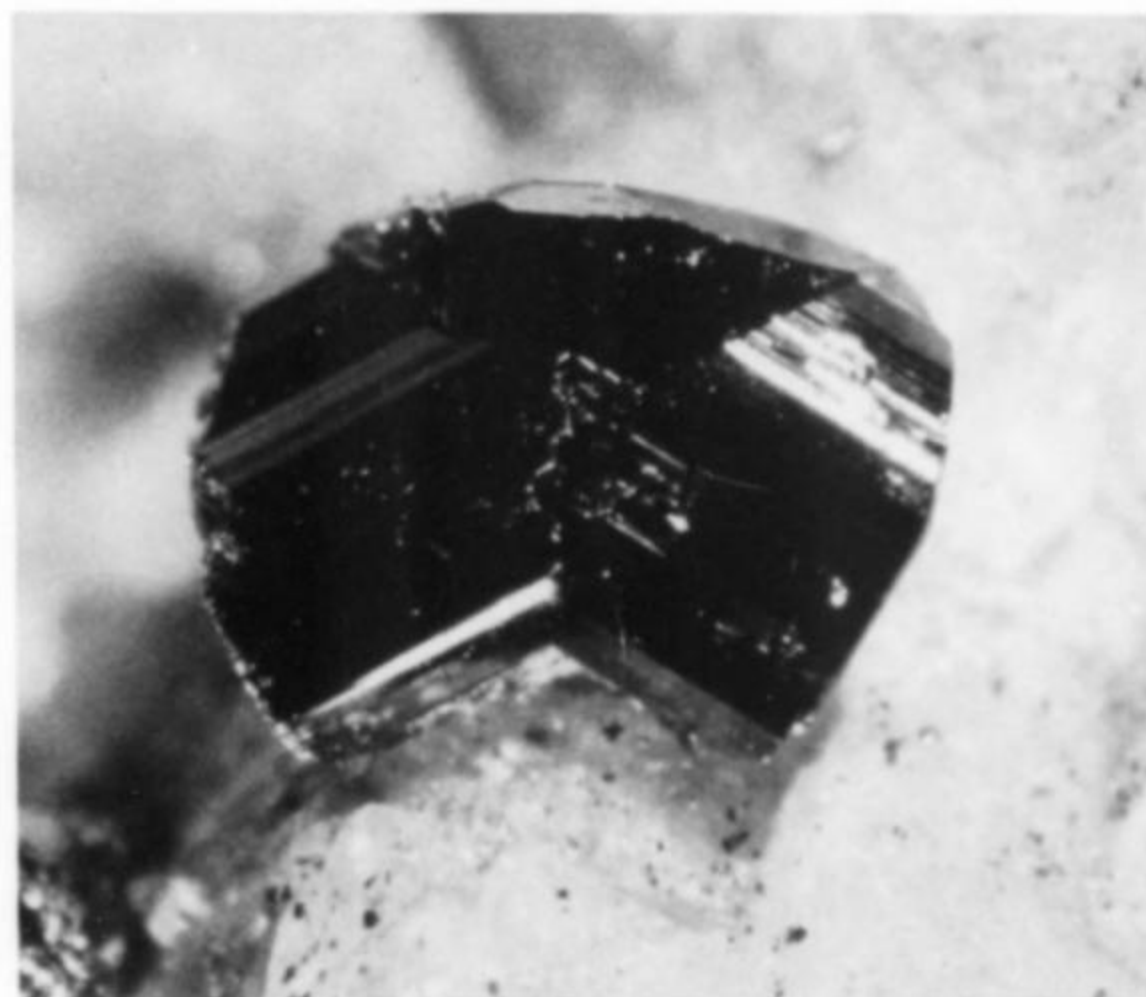


Figure 12. Superb, 1.5-mm derbylite crystal. University of Pisa collection.

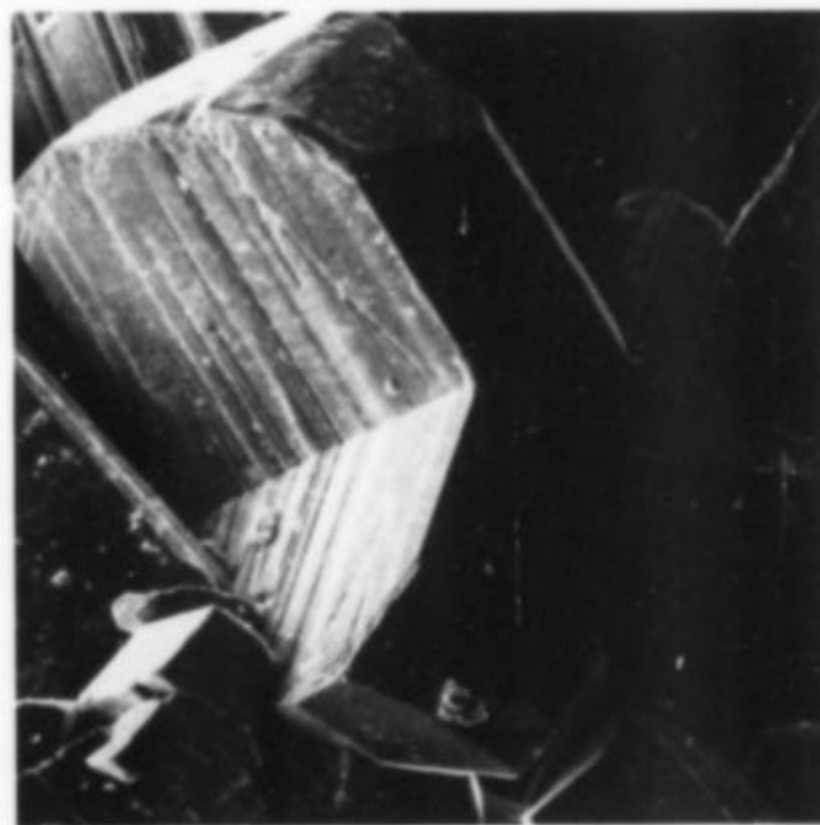


Figure 13. Derbylite twin (SEM photo, 150 X). Riccardo Mazzanti collection.

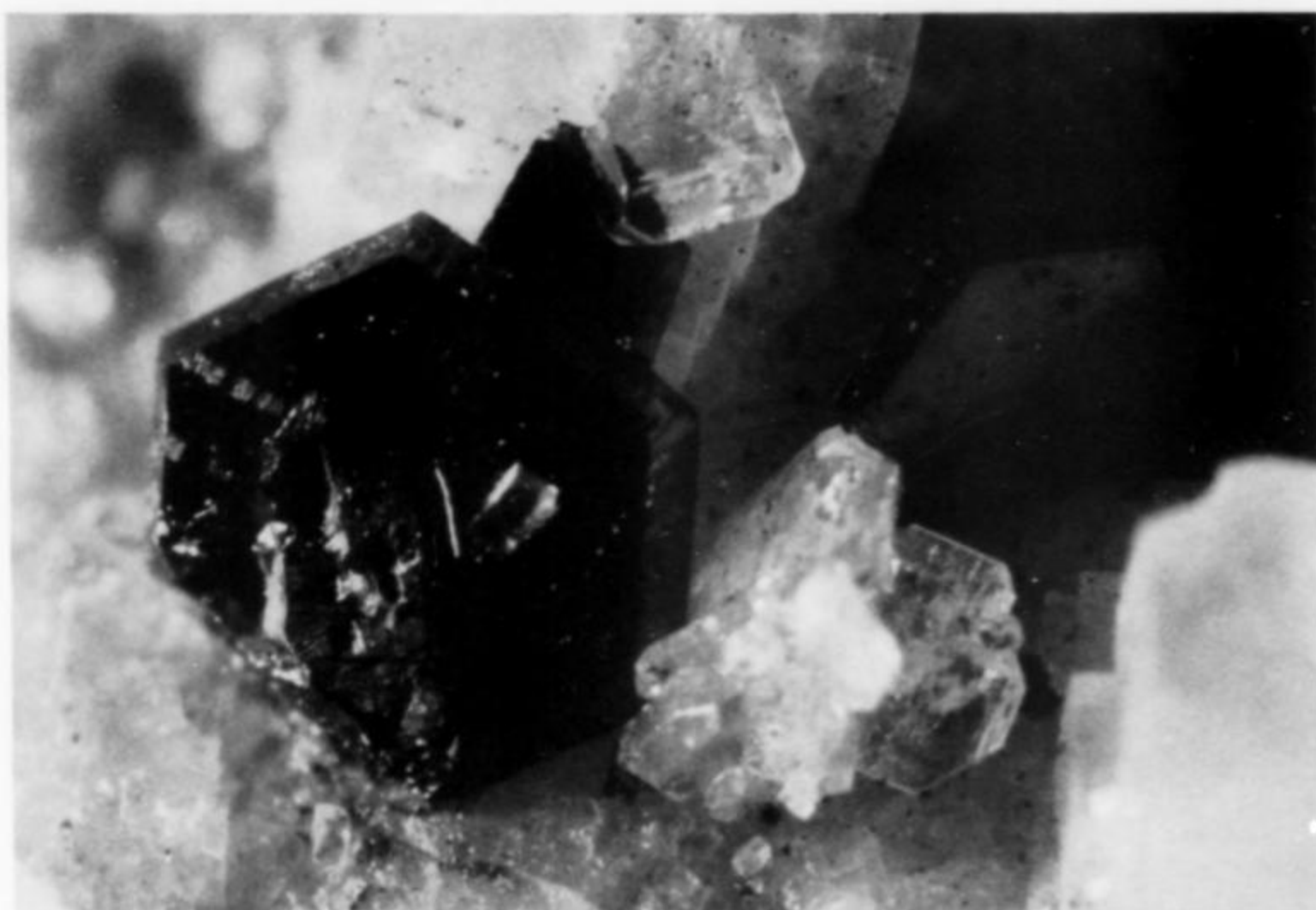


Figure 14. Flattened, rhombohedral hematite crystal 0.5 mm on edge.

Figure 15. Octahedral magnetite crystal in a vein with barite and dolomite. Field of view: 1 cm.

Malachite $\text{Cu}_2(\text{CO}_3)(\text{OH})_2$

Malachite occurs as disseminated or spherical aggregates of acicular crystals on bournonite and tetrahedrite crystals.

Muscovite $\text{KAl}_2(\text{Si}_3\text{Al})\text{O}_{10}(\text{OH},\text{F})_2$

"Phengite," a silica-rich variety of muscovite, occurs associated with albite and chlorite in dolomite near the contact with replacement bodies.

Pyrite FeS_2

Pyrite is the most abundant sulfide in the deposit. With barite, it forms microcrystalline lenses near the schist contact. Sporadic well-formed crystals (pyritohedrons) occur in cavities and in massive dolomite and calcite.

Quartz SiO_2

Rare, small, prismatic crystals occur associated with calcite and dolomite.

Rutile TiO_2

Rare, black, prismatic crystals of rutile occur in cavities in dolomite and marble.

Schafarzikite $\text{Fe}^{+2}\text{Sb}_2^{+3}\text{O}_4$

Schafarzikite is particularly rare at Buca della Vena, rarer than versiliaite and much rarer than apuanite. Of the three species, schafarzikite has the dullest luster.

Siderite $\text{Fe}^{+2}\text{CO}_3$

Siderite forms veins in the iron ore and also occurs as well-developed crystals in secondary veins in dolomite.

Spessartine $\text{Mn}_3\text{Al}_2(\text{SiO}_4)_3$

Spessartine occurs as millimeter-size, orange-brown, trapezoidal crystals in marble.

Sphalerite $(\text{Zn},\text{Fe})\text{S}$

Sphalerite is generally found as rare anhedral individuals in ore, with small pyrite crystals associated. Rare, honey-yellow, tetrahedral crystals occur in cavities in dolomite, and tiny, ruby-red crystals have been observed with altered aggregates of apuanite.

Stibivanite $\text{Sb}_2^{+3}\text{V}^{+4}\text{O}_5$

Stibivanite is extremely rare, forming as a green felt of acicular crystals in cavities in pinkish dolomite. The largest crystal seen measures more than 3 mm; most are much smaller (Merlino and Orlandi, 1983).



Stibnite Sb_2S_3

Metallic, gray, prismatic, striated crystals of stibnite occur associated with hematite in dolomite veins.

Stolzite PbWO_4

We have observed only three crystals of stolzite since beginning this study. They are yellow-brown with typical resinous luster and occur associated with hematite in veins.

Tetrahedrite $(\text{Cu},\text{Fe})_{12}\text{Sb}_4\text{S}_{13}$

Tennantite $(\text{Cu},\text{Fe})_{12}\text{As}_4\text{S}_{13}$

Tetrahedrite-tennantite occurs associated with beryl and pyrite as lustrous black crystals in cavities in dolomite. The habit consists of a very large number of small faces of different forms.

Vanadinite $\text{Pb}_3(\text{VO}_4)_3\text{Cl}$

Small ($\ll 1$ mm), lustrous, yellow, prismatic crystals of vanadinite occur with aggregates of altered apuanite.

Versiliaite $\text{Fe}_2^{+2}\text{Fe}_4^{+3}\text{Sb}_6^{+3}\text{O}_{16}\text{S}$

Buca della Vena is the type locality for versiliaite. The mineral is very rare and occurs intimately associated with the much more

Figure 16. Superb platy, green stibivanite crystal, 3 mm across. Pierluigi Pierotti collection.

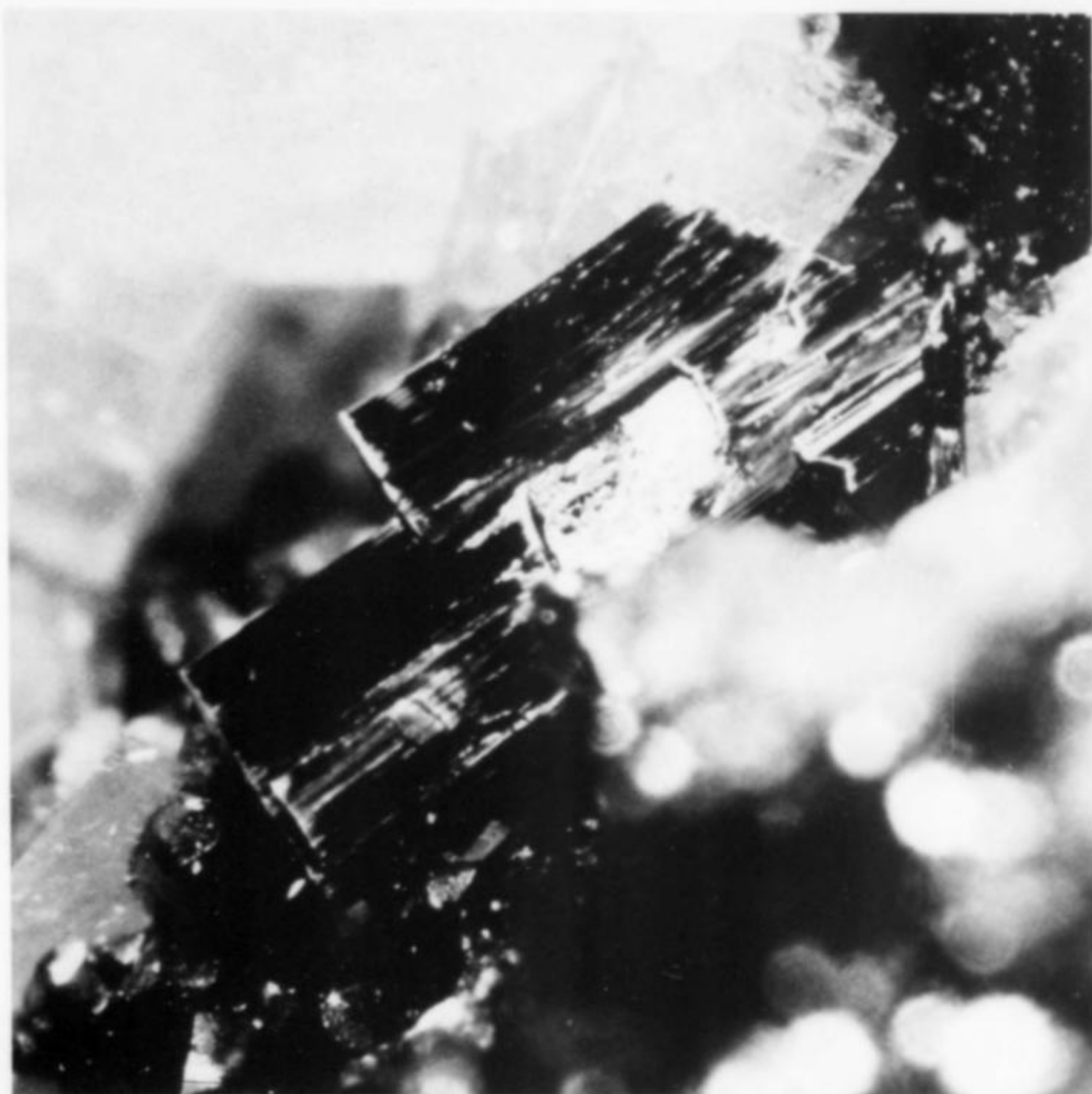


Figure 17. Prismatic vanadinite crystals with anatase (SEM photo, 240 X).

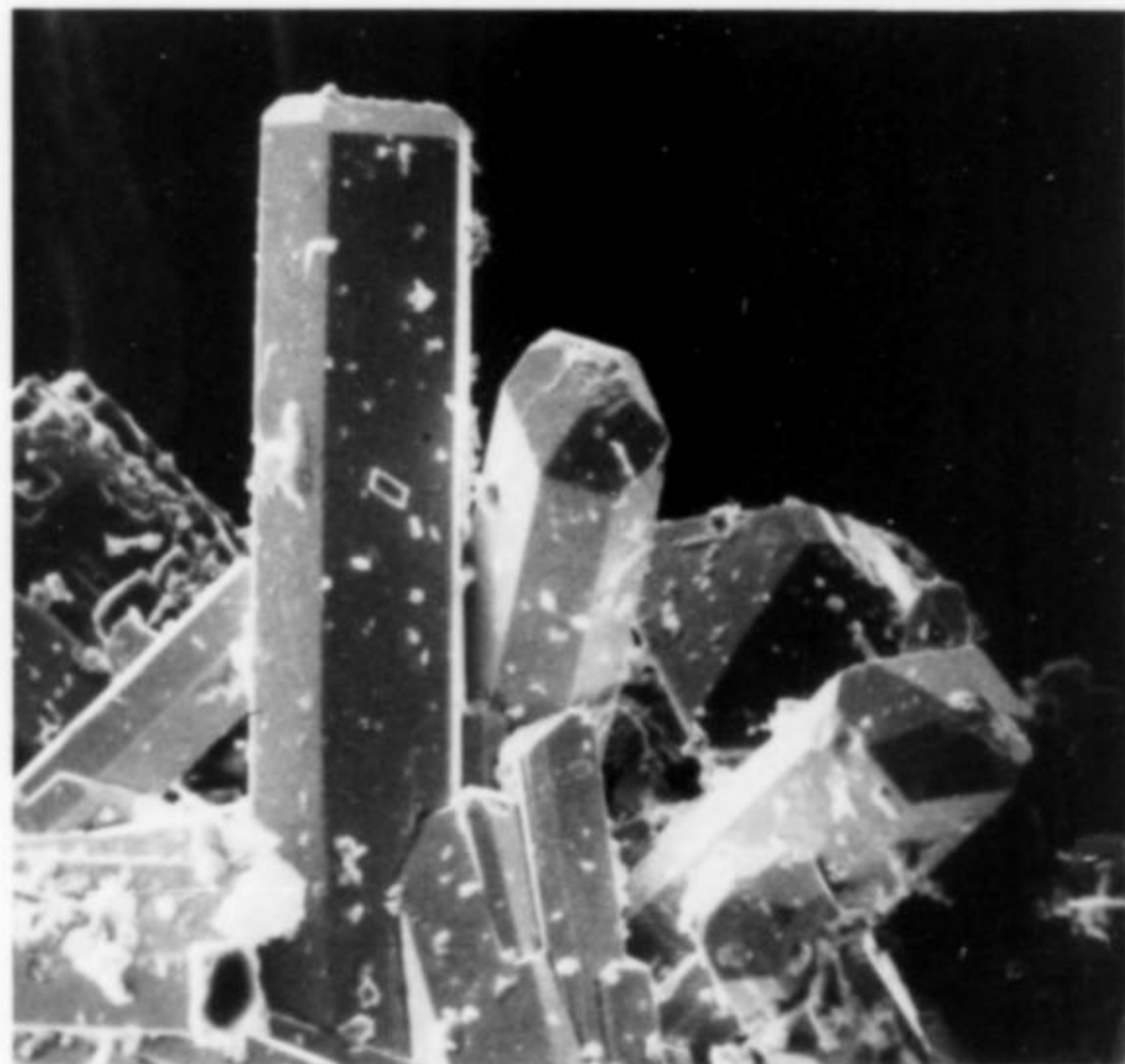


Figure 18. Large crystal of versiliaite, 5 mm on edge. Checchi collection.

abundant apuanite. A TEM study (Mellini *et al.*, 1981) has shown versiliaite, apuanite and schafarzikite intergrown down to the unit cell scale.

Vivianite $\text{Fe}_3^{+2}(\text{PO}_4)_2 \cdot 8\text{H}_2\text{O}$

Small, millimeter-size aggregates of granular, apple-green vivianite crystals occur associated with hematite and sphalerite.

Acicular Unknowns

Two different (new?) species, occurring in platy-acicular, metallic, black crystals, have been under preliminary X-ray study. A qualitative SEM/EDS analysis revealed the presence of Sb, Pb and S in both phases. A full description is in preparation.

PARAGENESIS

All of the minerals described here were collected mostly from the dump. However, a careful study of the interior surfaces of the mine itself has helped to deduce the original location of the various species.

Minerals of the schafarzikite group appear to have occupied a very restricted area near the zone where the upper verrucano lens becomes thinner and the two carbonate lenses come in close contact. Of the three minerals, schafarzikite contains Fe^{+2} only, whereas the others contain Fe^{+2} and Fe^{+3} . The $\text{Fe}^{+3}/\text{Fe}^{+2}$ ratio of apuanite exceeds that of versiliaite. These minerals are very well-crystallized in veins and in massive carbonate rock. Therefore, we think that the conclusions drawn by Carmignani *et al.* (1976) re-

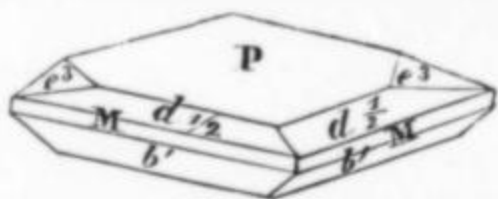
garding siderite, magnetite and hematite are relevant to the schafarikite group as well. Carmignani *et al.* observed siderite in the nucleus of magnetite crystals, and hematite along fractures in magnetite, suggesting the sequence of formation/alteration: siderite → magnetite → hematite (i.e., $Fe^{+2} \rightarrow (Fe^{+2} + Fe^{+3}) \rightarrow Fe^{+3}$). We think the sequence schafarikite → versiliaite → apuanite took place in parallel, as a function of the iron oxidation state in these minerals.

Beryl and allanite, normally found in pegmatites, occur at Buca della Vena in veins and along the contacts surrounding carbonate relics in the lower ore lens. Rare titanium and vanadium minerals have also been observed, suggesting that fluids responsible for replacement of the carbonate lenses leached Be, Ti and V (originally present only to traces in the carbonate rocks) and precipitated the rare minerals. A detailed study of the paragenetic interrelationships of *all* species at Buca della Vena might allow an expansion of the genetic model of Carmignani *et al.* (1976).

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Microminerals

Bill Henderson

This column will be a veritable witch's brew, with a bit of this and a smitch of that. However, I hope it will be a bit more palatable. First, a few items which will no doubt appear in future editions of the *Guinness Book of Records*. Figure 1 is a black and white print of a hand-colored illustration from *Unterhaltung aus der Naturgeschichte*, by Gottlieb Tobias Wilhelm (*Das Mineralreichs*, 2 volumes, Vienna, 1825-1828). Sent me by Carlo Cassinelli, it undoubtedly shows the world's first micromounter, for why else would the gentleman in the top hat be employing a magnifying glass?

The colorless and very shapely epistilbite crystal in Figure 2 ranks as the oldest micro specimen in the author's collection. As the labels in Figure 3 show, the specimen was collected in 1877, and passed through the University of Copenhagen.

The bertrandite in Figure 4 was the very first micromount in the author's collection, and thus might also qualify for the *Book of Records*. Decades ago, in 1957 to be exact, I didn't know beans about minerals. All I had had were two or three university courses in the subject. At that time, while a poor graduate student at Yale, I was given to traveling as far as I could on weekends in search of new specimens. At the Strickland quarry in Portland, Connecticut, I collected some tiny crystals, diamond shape in cross section, and very thin. They were transparent and a very faint pink in color when viewed on edge. Their identity was a mystery to me, and also to the faculty at Yale, so I took them one night to Neal Yedlin. Neal not only identified them on sight as bertrandite but also mounted them in one of his own cardboard boxes as shown. This first meeting was the beginning of a friendship lasting many years, and was the starting point for my own micromounting activities.

Turning to other topics, it is interesting how often a mineral can be either missing from a micromounter's collection for years or else represented by only mediocre specimens, and then, within the space of a month or so, he receives not one but several much finer specimens from two or more locations. Such happened to me recently with respect to thomsonite, an unusual but not rare member of the zeolite family. Recently, I received really superb crystals of thomsonite such as those shown in Figure 5 from Jon Mommers (P.O. Box 219, Sunshine, Victoria 3020, Australia). The square, tabular crystals from the basalts of Simmons Bay are much larger and better formed than any I had seen previously, and are frequently associated with what appears to be natrolite in epitaxial growth. Only months later, equally fine single crystals (Fig. 6) and groups (Fig. 7) of thomsonite were received from Domenico Forloni (Via Trento 19, 20017 Passirana-Rho [Milano], Italy).



Figure 1. Reproduction of a hand-tinted illustration showing, perhaps, the world's first micromounter. From a book by Gottlieb Tobias Wilhelm published in 1825.



Figure 2. Colorless, 5-mm crystal of epistilbite from the Teigarhorn, Iceland, collected in 1877.

Epistilbit.
Caroc Teigarhorn.
 1877.2626. Island.

FROM THE MINERALOGICAL AND GEOLOGICAL MUSEUM OF
 THE UNIVERSITY OF COPENHAGEN.

.....
 EPISTILBITE

TEIGARHORN
 ICELAND

(15)

Figure 3. Earlier label and University of Copenhagen label for the specimen shown in Figure 2.



Figure 4. Bertrandite, faintly pink, 2-mm crystals from the Strickland quarry, Portland, Connecticut, in Neal Yedlin paper micromount box.

These crystals from the slopes of Monte Somma are even larger than those from Australia, and it is very hard to decide which are the finer specimens.

Jean Jenks (Prices Road, RD 2, Upper Moutere, Nelson, New Zealand), has over the years sent me a number of very nice micromineral specimens. Recently, she sent some distorted pyrite associated with chabazite and a most interesting rhombohedral carbonate, all in basalt from Lyttelton Borough, New Zealand. The



Figure 5. Colorless, transparent fans of square, tabular thomsonite crystals. From Simmons Bay, Flinders, Victoria, Australia; the largest group is 3.5 mm across.

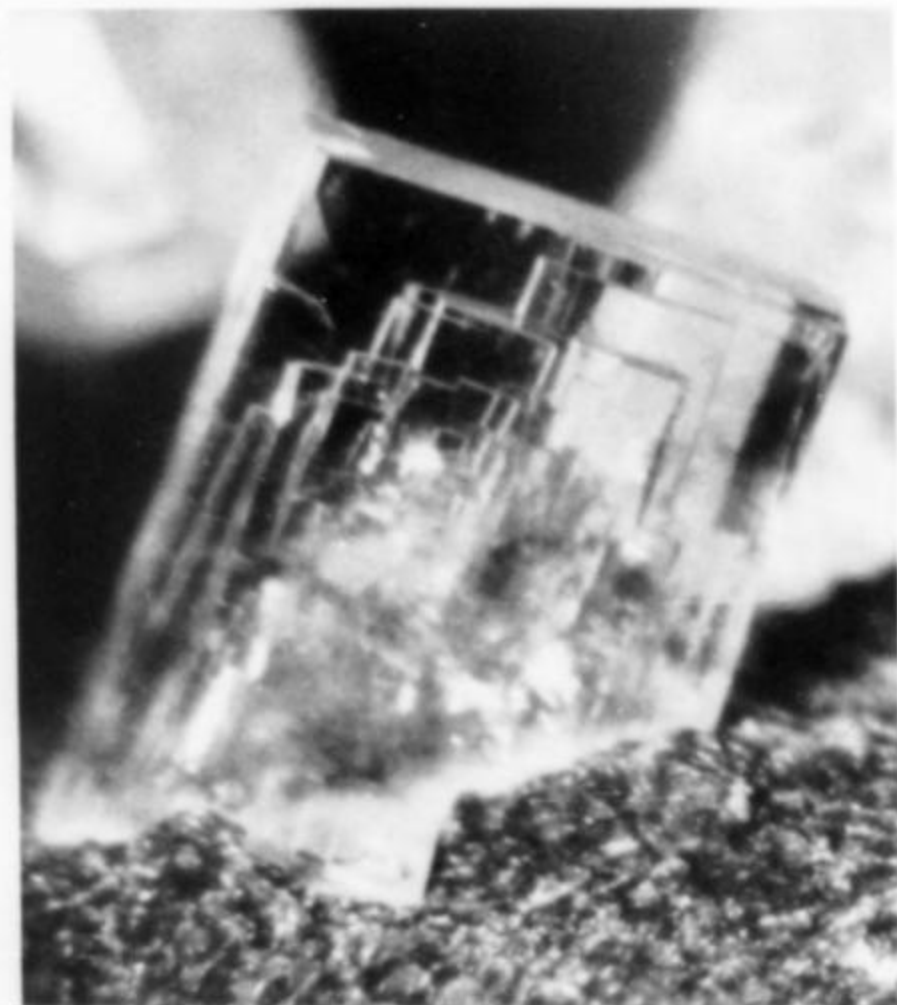


Figure 6. Transparent single thomsonite crystal, 1.5 mm high, from Monte Somma, Vesuvius, Italy.

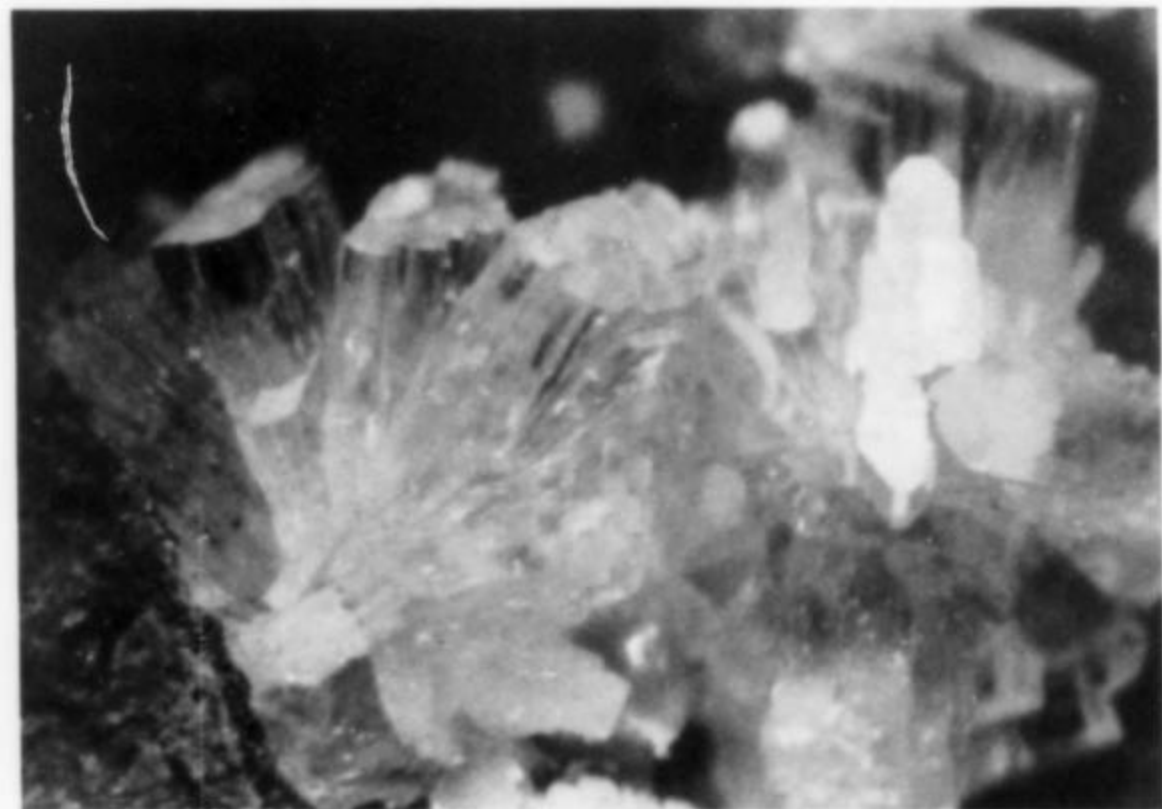


Figure 7. A 2.5-mm group of thomsonite crystals from Monte Somma, Vesuvius, Italy.

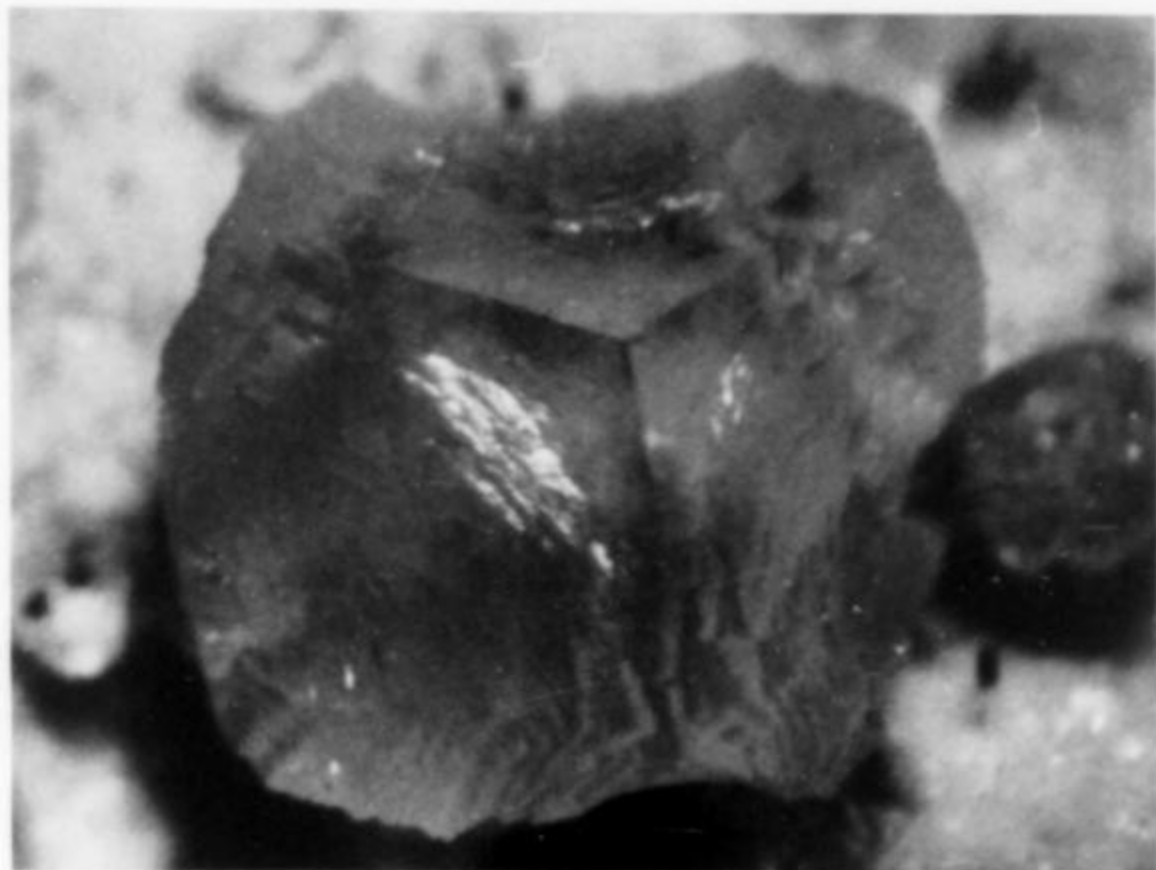


Figure 8. A 3-mm crystal of calcite showing both rhombohedral and saddle-shape crystal forms, from Lyttelton Harbor Board quarry, Lyttelton Borough, New Zealand.



Figure 9. White needles of mesolite on orange heulandite in a bean-shape, 4-mm cavity. From Rock Creek, Stevenson, Skamania County, Washington.

Figure 11. Extremely large (0.5 mm) transparent, colorless crystal of paulingite with yellow inclusions; Three Mile School, near Ritter, Grant County, Oregon.

carbonate (Fig. 8) looks for all the world as though it began growth as a simple rhomb such as one sees in calcite, but it then changed habits to the saddle shape characteristic of dolomite and, to a lesser extent, rhodochrosite. The crystals are colorless to white, and their optics fit perfectly for calcite. To be certain of its identity, a crystal was analyzed by microprobe, and found to contain more than 98% calcium (carbonate) plus minor amounts of magnesium and iron. The saddle-shaped crystals are therefore definitely calcite of a fairly unusual habit.

Returning to the wonderful world of zeolites, readers might like to inspect the beautiful mesolites shown in Figures 9 and 10. These were sent to me by Janice M. Healy (P.O. Box 5414, Aloha, Oregon 97006) (Yes, Oregon, not Hawaii). They occur on a white

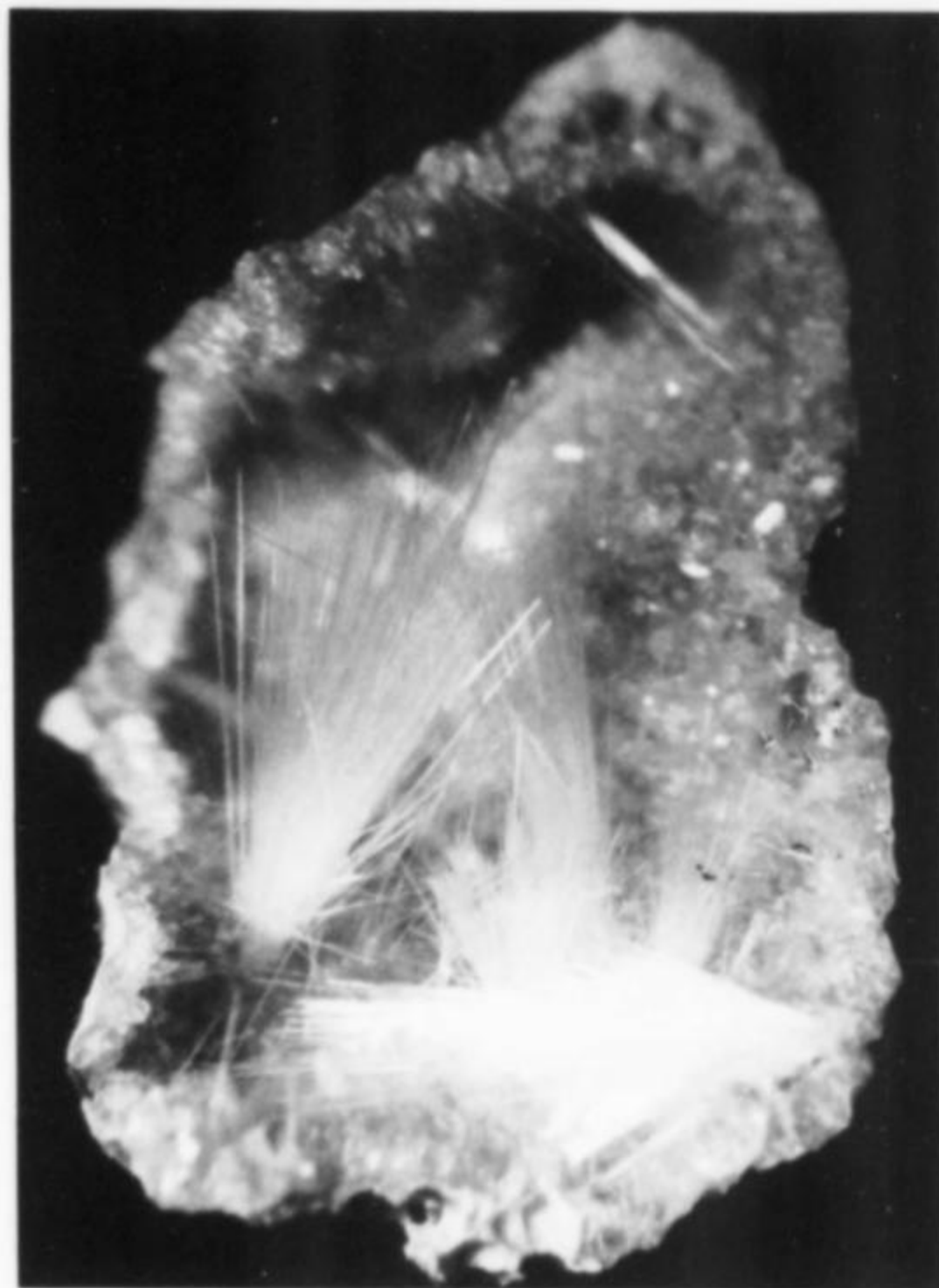
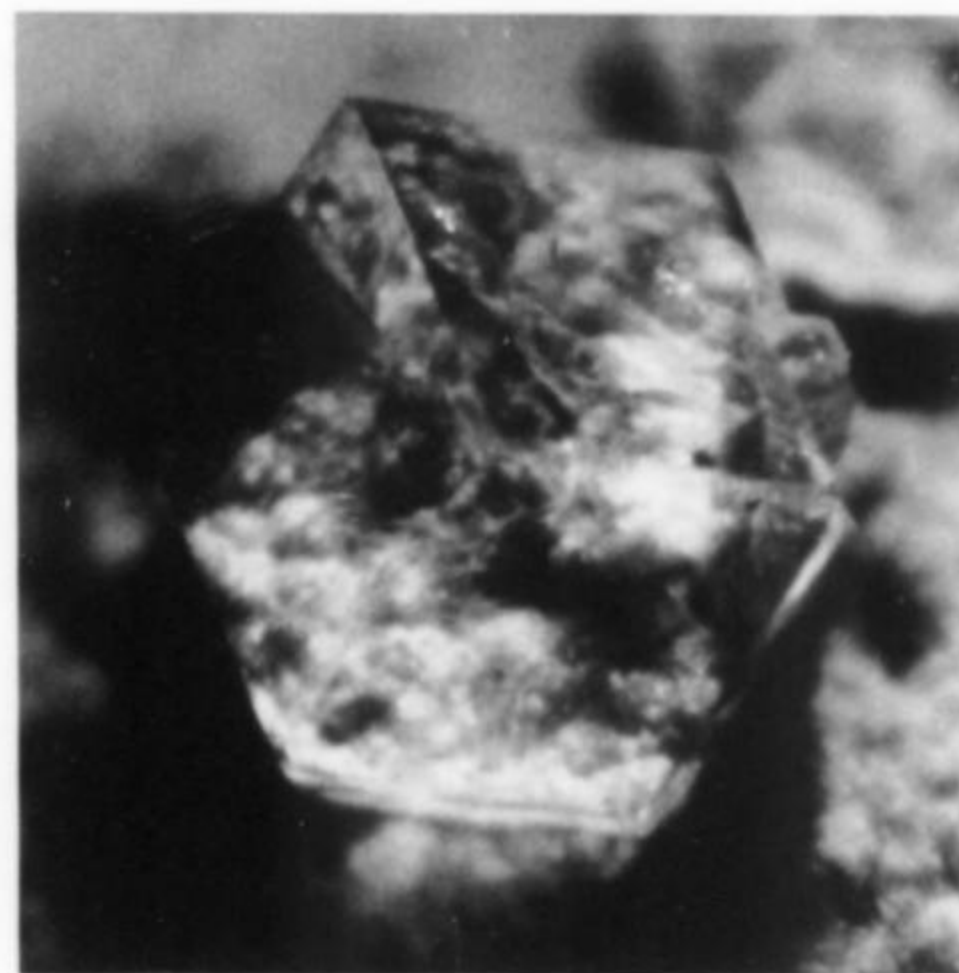


Figure 10. A 6-mm group similar to that in Figure 9.



to pink background of minute heulandite crystals in tiny, loose, bean-shaped vugs, and are from Rock Creek, Stevenson, Washington. Janice would be pleased to exchange these "beans" for micros from your area.

Before leaving zeolites, I have to mention the colorless, absolutely transparent paulingites from Three Mile School, near Ritter, Oregon, sent me by Vi Frazier (P.O. Box 27, Point Arena, California 95468). These are nice not only for the yellow inclusions visible at the point of attachment of the crystals (see Figs. 11 and 12) but also for their size. The largest crystals are 0.4–0.5 mm across, and those familiar with the mineral will understand that these are truly humongous giants for the species. Vi has sent me many other zeolites over the years, including a very interesting suite from



Figure 12. Two 0.4-mm crystals of paulingite; Three Mile School, near Ritter, Grant County, Oregon.



Figure 13. Pale violet, 3-mm crystals of creedite from the Liberty mine, Tonopah, Nye County, Nevada.

basalts associated with the borates at Boron, California.

Walter Lombardo (4728 Elm Avenue, Las Vegas, Nevada 89110) is a collector-dealer who recently sent me some superb micros of creedite from the Liberty mine, Tonopah, Nye County, Nevada. The crystals (Fig. 13) are very sharp, transparent, and vary from colorless to pale amethyst. While not as large as the ones recently found in Santa Eulalia, Mexico, the sharpness and transparency of the crystals makes them very worthwhile. Walter has other choice micro species such as vanadinite from the Ruth mine, Goodsprings, Nevada, in crystals which rival those from Arizona for color; and blue celestite from the Billie mine, Death Valley, California.

Wayne DeBrusk, *the Rare Earth*, has been advertising 3-5 mm cubanite crystals from the Sudbury, Ontario, area in recent issues

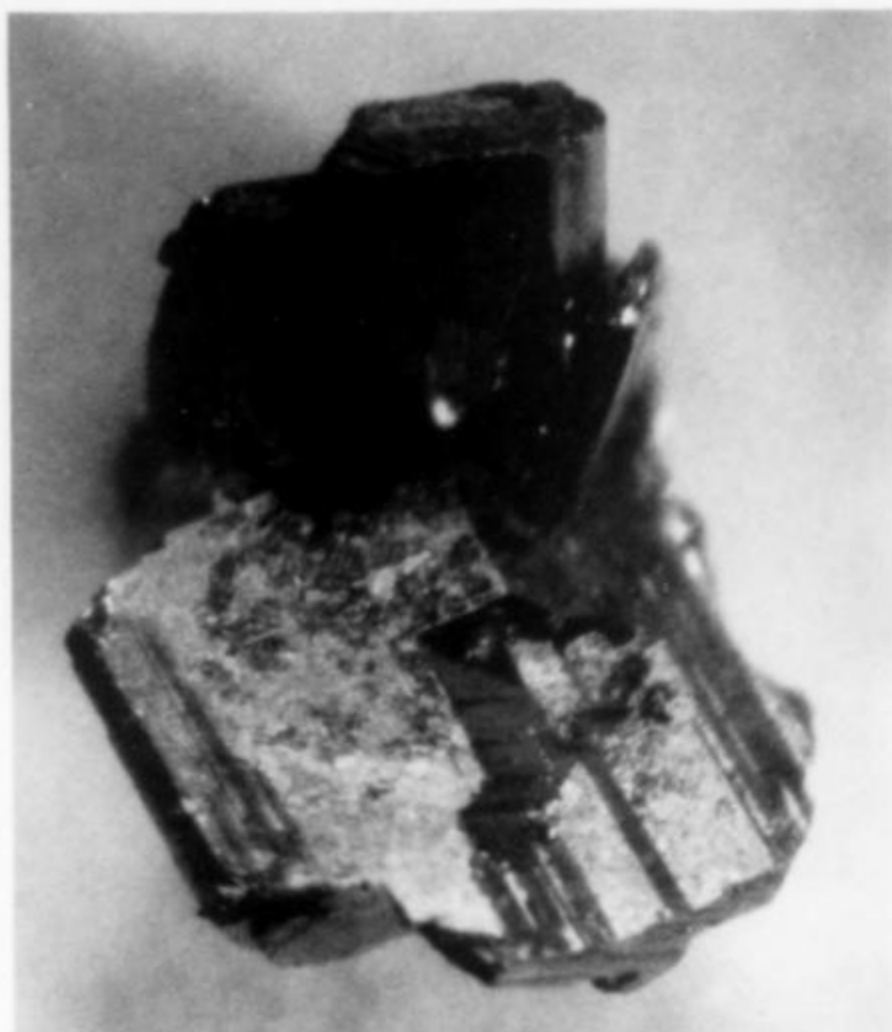


Figure 14. A 2-mm group of brass-yellow cubanite crystals; locality, the Strathcona mine, Onaping, Ontario, Canada.

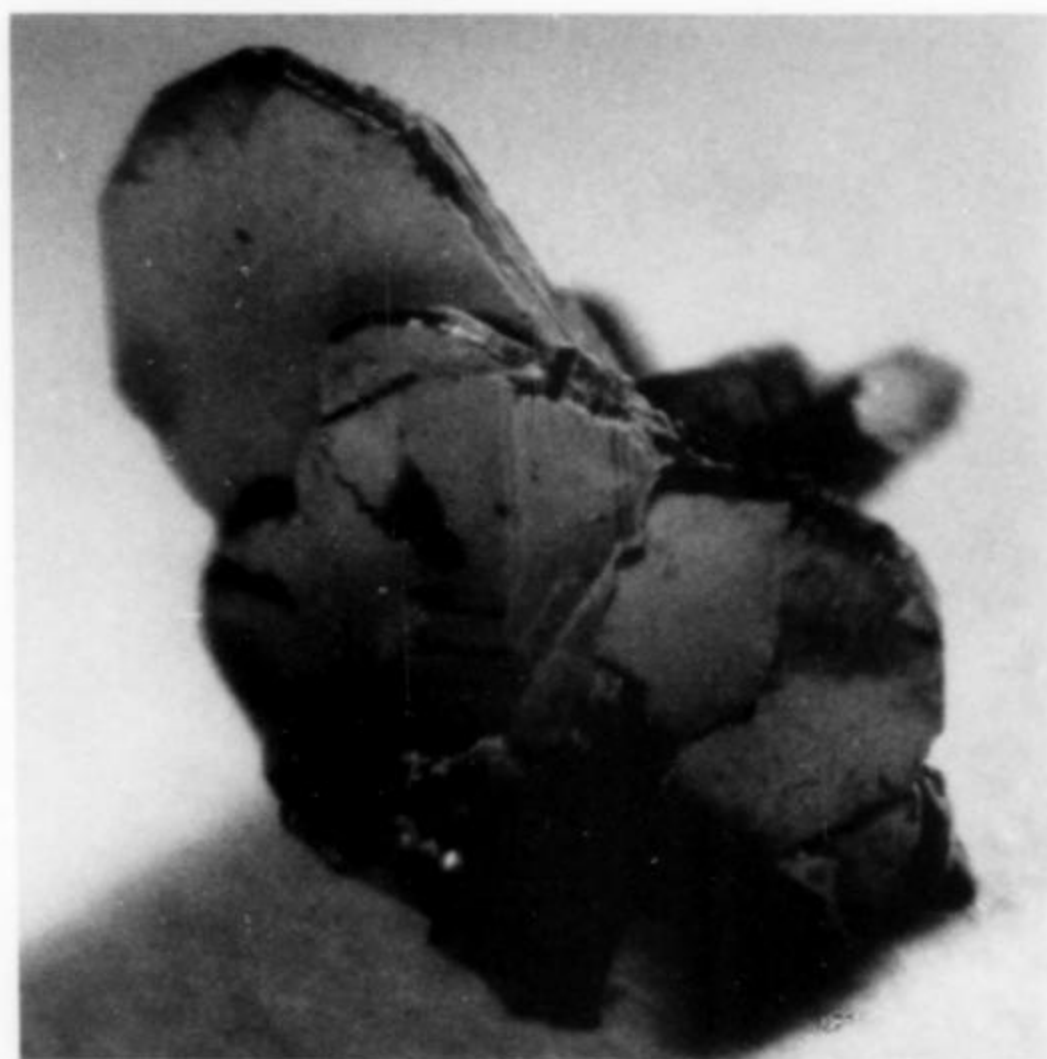


Figure 15. Tabular and rounded brass-yellow crystals of cubanite to 1.5 mm across, from the Strathcona mine, Onaping, Ontario, Canada.

of the *Mineralogical Record*. Rather than receive a single rather large crystal as advertised, I asked him to send two smaller, micro size crystals. As can be seen in Figures 14 and 15, the crystals of this rare mineral are indeed excellent, and I urge you to write for some while they last.

Many micromounters think that, because the specimens they work with are so small, they can do little in the way of chemical testing to prove their identity. This is hardly the case, and a small number of advanced collectors have developed a variety of spot tests and other analytical methods that are elegant for very small fragments of crystals. Dana's *Manual of Mineralogy* and *Textbook of Mineralogy* list simple tests for the elements, and also describe diagnostic chemical tests for the more common minerals. Not long ago, I had occasion to use some simple closed tube tests to distinguish between various suspected minerals. While Dana

describes the use of 1/8 inch diameter tubes or larger, I used melting-point capillaries. These closed-end tubes, about 1 mm in inside diameter, are readily available from scientific supply houses.

The first specimen to be checked was one from the Noyes mine, Madoc, Ontario, and might have been arsenopyrite (FeAsS) or marcasite (FeS_2) on the basis of its twinned crystals. Dana's *Text-book* says of arsenopyrite, "In the closed tube, gives . . . a conspicuous sublimate of metallic arsenic which is of bright grey crystals near the heated end and of a brilliant black amorphous deposit farther away." For marcasite in the closed tube, Dana's *Manual* describes a copious yellow sublimate of sulfur. To test for this pair, tiny fragments of known arsenopyrite, marcasite and the unknown were placed in three melting-point capillaries. A microscope was needed even to see that the fragments were indeed in the capillaries! Always, always, always, when able, run known minerals alongside the unknowns. It's a lot safer, as tests sometimes don't match those described in the literature, and another element sometimes masks or confuses tests for the element being sought. Since both arsenopyrite and marcasite are of low melting point, it was sufficient to heat the capillaries in nothing more than a candle flame. After a few seconds, the resulting sublimates were those shown in Figure 16, showing conclusively that the unknown gave a sublimate of sulfur like that of marcasite, and completely unlike the arsenic crystals and mirror of arsenopyrite.

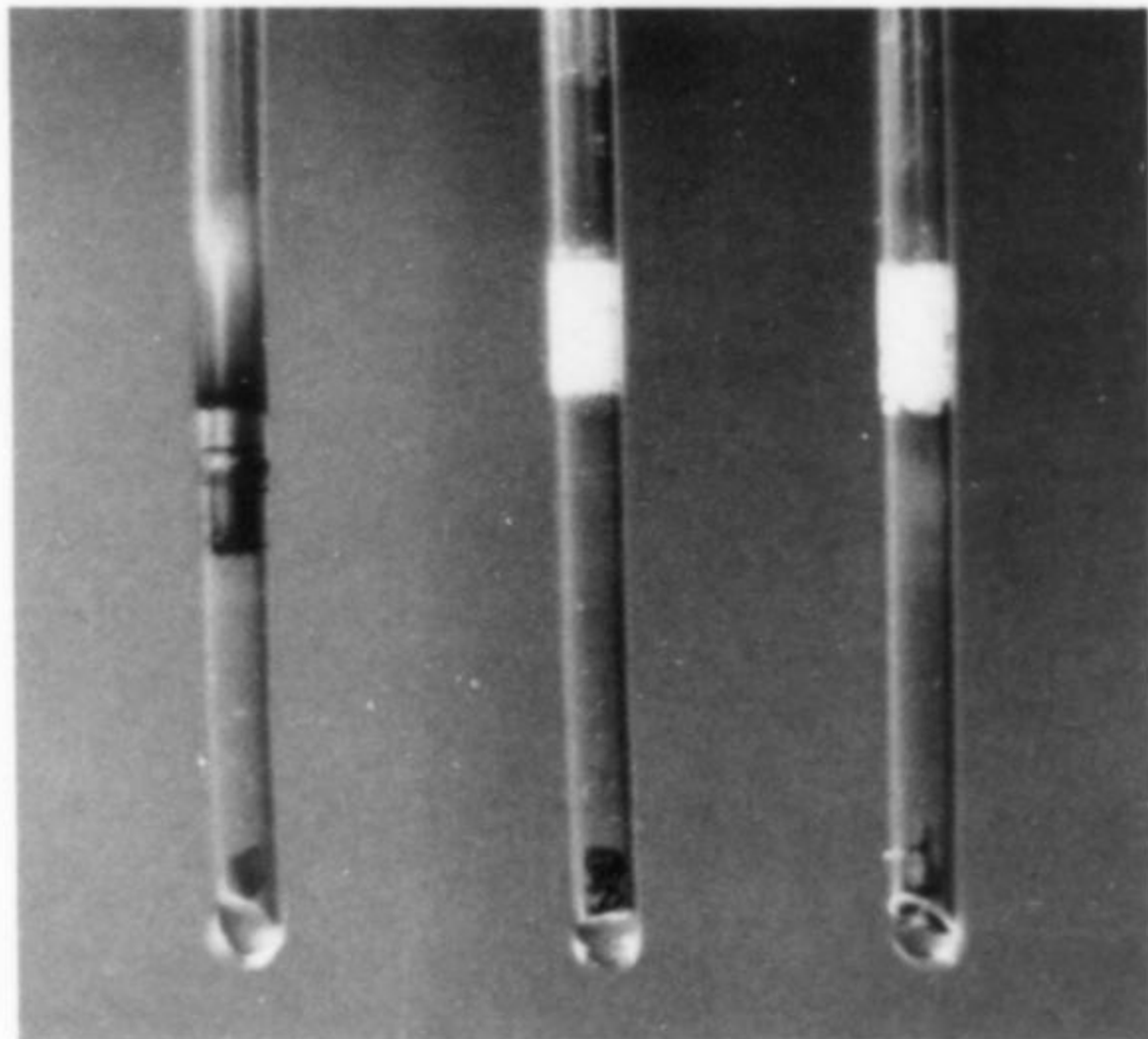


Figure 16. Micro closed-tube chemical tests showing: (left) a shiny black mirror above black crystals, characteristic of arsenic in known arsenopyrite; (center) light yellow sublimate from an unknown mineral; and (right) light yellow sublimate of sulfur from known marcasite.

In the second case, I was concerned with a specimen showing black, acicular crystals from Arizona. These were variously described as being groutite ($\text{MnO}(\text{OH})$), a very difficultly fusible mineral, or stibnite (Sb_2S_3), a mineral which fuses very readily. A similar series of tests using a candle flame and capillary tubes gave the results shown in Figure 17. While known groutite and the unknown failed to fuse and gave no sublimate, known stibnite gave the red antimony oxy-sulfide and yellow sulfur sublimates described in the textbooks for stibnite. Again, the distinction was unmistakable, and, given a choice of these two species only, the mineral was taken as being groutite.

The Mineralogical Record, volume 17, July-August, 1986

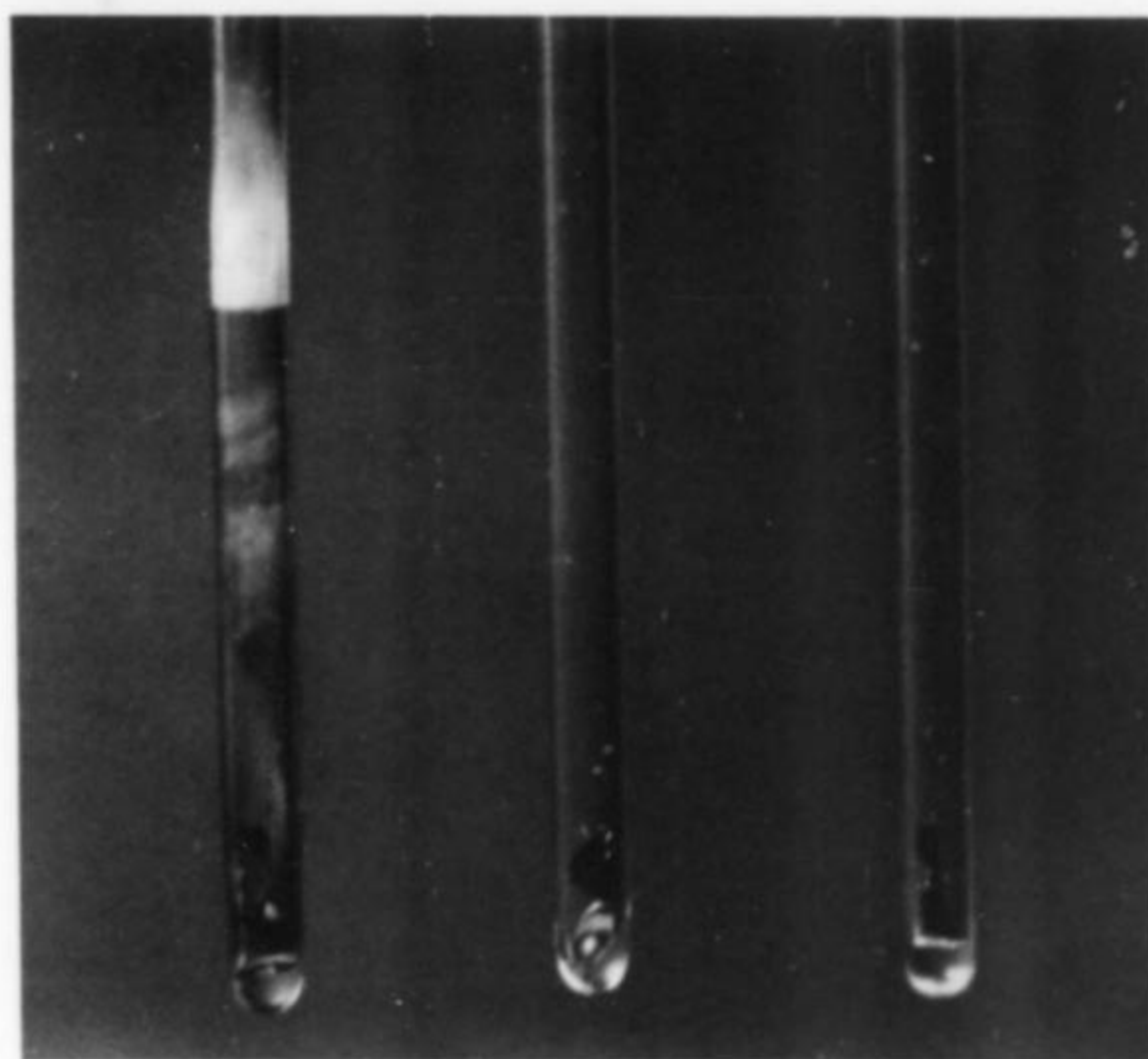


Figure 17. Micro closed-tube chemical tests showing: (left) light yellow sublimate of sulfur above a less distinct band of red antimony oxy-sulfide from known stibnite; (center) no reaction from infusible unknown; and (right) no reaction from known groutite.

These tests are highly useful, and readers will have fun devising diagnostic micro tests of their own. Samples as small as those shown in Figure 18 or even smaller are all that are required, and almost every micro specimen has an inconspicuous or broken crystal which can be sacrificed for such testing. As Neal Yedlin used to say, "Read a good mineral book," and learn to do your own identifications.

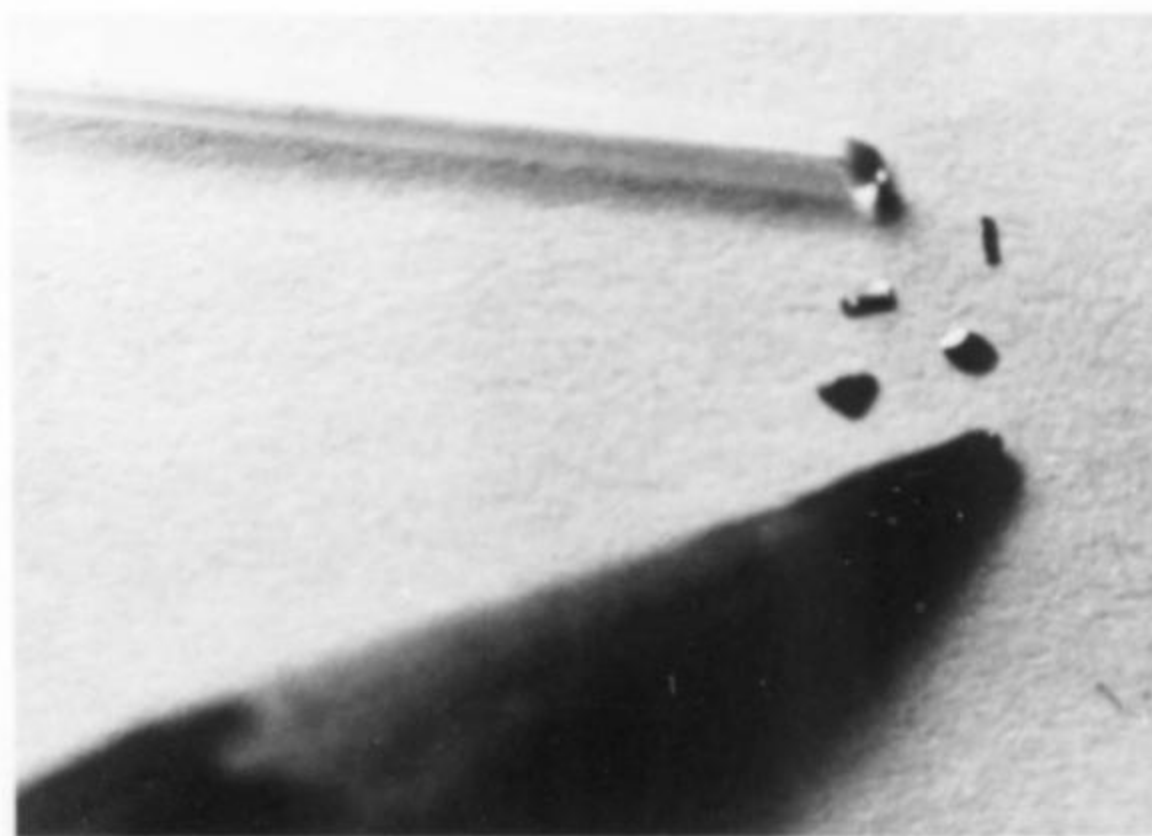


Figure 18. Size comparison of a pencil point with samples for closed tube chemical tests in melting-point capillaries.

Not long ago, I was sorting tiny crystal groups found in gravel collected at Mont St-Hilaire, Quebec in 1981. Found in the gravel were pink ancylite "asters," extremely thin rosettes of catapleiite, pink rhombs of rhodochrosite, colorless epididymites, fragile needles of aegirine, natrolite, fluorite and sphalerite. The groups were so tiny, though, that they were frequently crushed and broken or else dropped while transferring them to medicine capsules. What to do? The solution was to glue two short sections of thin rubber

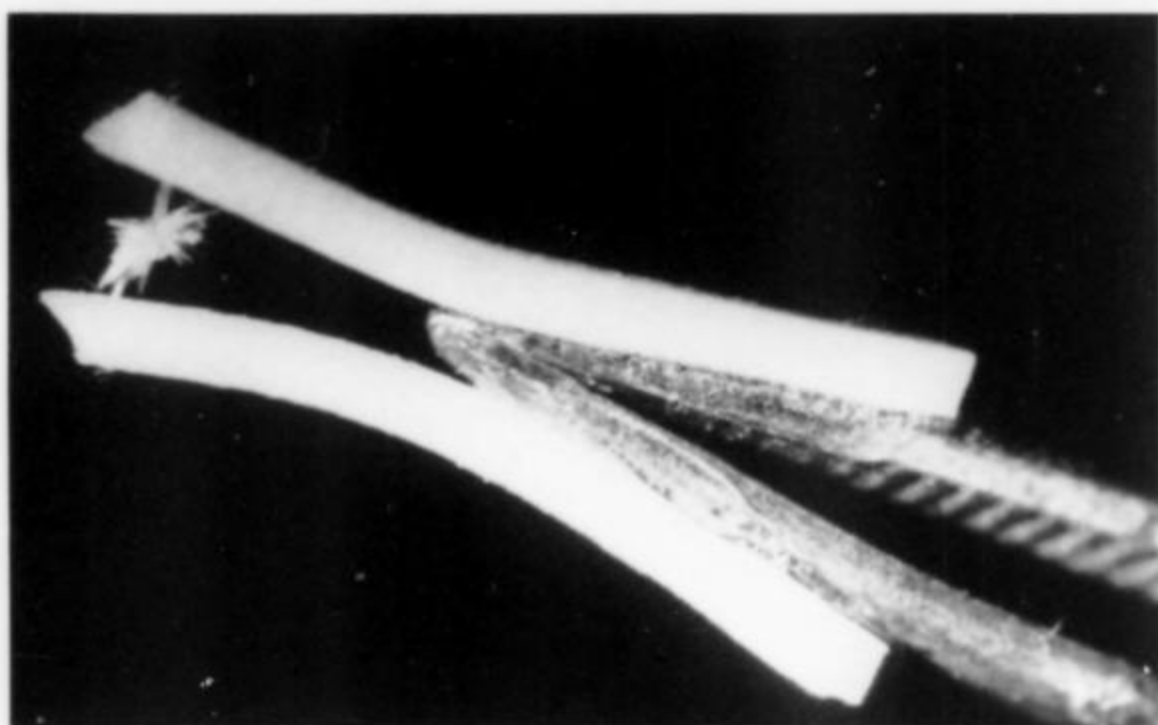


Figure 19. Design of metal tweezers tipped with portions of rubber bands for handling very fragile crystal groups. Being held: a 0.5-mm group of very pale pink ancylites from Mont St-Hilaire, Quebec, Canada.

bands to the *outside* edge of the tips of a pair of fine tweezers. As shown in Figure 19, it was then possible to pick up and transfer the smallest crystal groups without crushing or dropping them. Further, it was found that even quite large and blocky crystals could be transferred with ease, as the rubber tips had amazing sticking power. Putting the rubber bands on the outside of the tweezer tips means that pressure can be exerted to completely close the metal tweezer tips without either deforming and bending the rubber tips or applying excess pressure to the specimen. While almost certainly a patentable idea, I bequeath this freely to the micromineralogical fraternity. On this note of outstanding generosity, the column is closed.

Happy hunting!

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Letters

PEGMATITIC DANBURITE

Regarding Lanny Ream's letter (vol. 17, no. 2, p. 145) in which he states that Butte, Montana, and Anjanabonoina, Madagascar are the only known pegmatitic occurrences of danburite, I can point out two more. Large, prismatic crystals of danburite up to 15 cm in size were found in a pegmatite, with green tourmaline, near La Huerta, Baja California, Mexico (personally collected by Josephine Scripps and myself in 1969). Danburite has also been found in the J-27 pegmatite near Riverside, California.

Al Ordway
Bloomington, CA

There is a pegmatite occurrence of danburite in Norway. Specimens collected from the Tangen quarry near Kragerø, in southern Norway, were sent to the Mineralogy-Geology Museum in Oslo for identification in 1906-1908. They were provisionally identified as topaz, but were recognized as danburite by I. Oftedal in 1963. The specimens measure up to 20 cm in size and are all fragments of two large crystals.

Alf Olav Larsen
Stathelle, Norway

John Sinkankas pointed out another locality listed in his *Gemstones of North America*: the Chuqui pegmatite mine near the El Fenomeno tungsten mine near Rosa de Castilla, Baja.

Lanny Ream

EXCHANGES

I am an advanced systematic collector and a *Mineralogical Record* subscriber since 1979. My own collection consists of about 10,000 specimens representing 2244 valid species (as listed by Fleischer). Roughly 180 square meters (1900 square feet) of my house has been taken over by the collection.

I'd be most appreciative if you could publish my name and address for the purpose of exchanging specimens. I am seeking rare minerals of systematic interest, any size, from localities worldwide. I specialize in minerals from Vesuvius; a list of all specimens available for trade, as well as a list of satisfied trading partners, can be provided.



(Enclosed is a photo showing about a third of my showcases.)

G. Galvani
Via de Amicis 35
20123 Milano, Italy

MISLABELING

We applaud the recent article on mislabeling on mineral specimens (vol. 17, no. 2, p. 99). Correct and specific locality documentation is important not only as a guide to further collecting but as a key to understanding the geochemical conditions which contributed to a mineral's crystallization, and thereby perhaps being able to predict where else it might occur. Consequently, we are particularly attentive to locality data accompanying specimens in the Harvard collection. We adopted the British Museum's square bracket convention for indicating attributions several years ago. We also like your suggested attribution labels and are having some printed for use in the museum.

Carl A. Francis
William C. Metropolis
Harvard Mineralogical Museum

The article on mineral specimen mislabeling was most interesting. I was surprised to see no comments on the American Federation of Mineral Societies practice of replacing chemical formulas with "simplified compositions" (for labels in competitive displays).

T. J. Schmierer
Albuquerque, NM

Personally I have always found that to be an odious practice which encourages mental laziness and sloppiness. "Simplified" compositions are by definition incorrect in most cases because of the omissions; how is the viewer to know what has been left out? Encouraging the use of simplified formulas merely panders to those collectors who are determined to avoid learning anything (such as the meaning of a few simple symbols). The AFMS would do better by its constituents if it set a higher standard.

Ed.

GENSTAR QUARRY

This is to inform your readers that the Genstar Medford quarry in Carroll County,

Maryland (see vol. 17, no. 2, p. 151-152) is effectively closed to collectors; to my knowledge, no entry has been granted this year, even to clubs which previously enjoyed a good relationship with Medford. This is supposedly due to some incidents in which collectors refused to leave a dangerous area when asked to do so by quarry officials, or who caused problems in other ways.

Fred J. Parker
Columbia, MD

SAINTE-MARIE aux MINES

Regarding Werner Lieber's brief reference to Sainte-Marie aux Mines, France, in the recent Silver Issue (vol. 17, no. 1, p. 6), I would like to offer some additional information. Although native silver and proustite from this area are today rather rare in collections, the early written records are quite informative.

In 1530 (not 1539) a mass of silver weighing 50 kg was broken loose in a stope; two pieces of it were presented as a gift to Emperor Charles the Fifth of Spain. In 1755 one of the mines yielded specimens of such beauty that they were sold as collection pieces. The most spectacular find, however, was made on October 17, 1581, by a miner named Klaus Shirbald; he found a block of native silver weighing 592 kg (1304 pounds!). The Chronicles describe it as a mass of "fir branches, feathers and curled grass."

Altogether more than 120 species are

known from Sainte-Marie aux Mines, including seven new species and many of exceptional quality. The deposits were apparently first worked in the 8th century. By the 19th century approximately 800 named mines and concessions had been exploited in the district. In recent years, archeological work at Sainte-Marie aux Mines has revealed much about the history and mining technology of the 16th century.

Alain Martaud
Bersee, France

MAILERS & MAILBOXES

Thank you for acting on Henry Fisher's suggestion to add "Please Do Not Fold" to the mailers. I was amused by your suggestion to consider buying a larger mailbox; I did so three years ago, with that very purpose in mind. It stands to reason, of course, that my latest issue was delivered to my neighbor (in a small mailbox) by mistake.

John Jaszczak
Columbus, OH

EPITAXIAL or EPITACTIC?

I was interested to read about marsturite epitaxial on rhodonite in the latest *Mineralogical Record* (vol. 17, no. 2, p. 123-125). I write to point out, in the true fashion of a nit-picking editor sent on this earth to torment authors, that use of the word *epitaxial* is inappropriate. There is nothing "axial" about the relationship

described. In fact, the root, which gives rise to the word *tactile*, refers to a surface or plane. I would advise a switch to one of the two adjectival forms, *epitactic* or *epitaxic*, proposed by International Mineralogical Association and the International Union of Crystallography (see, for example, *Canadian Mineralogist*, 16, 113-117, and in particular point M on p. 115).

Robert F. Martin
Editor, *Canadian Mineralogist*
Montreal, Quebec

RESPONSE FROM NATIONAL GEOGRAPHIC

I just recently had the opportunity to see some issues of the *Mineralogical Record*; it is handsome and well written. I particularly appreciate the locality descriptions that present the site history and geology as well as the mineralogy. As proof of my sincerity I enclose a check for a subscription.

Now about that statement (vol. 16, no. 6, p. 440), "Try sending *National Geographic* an old back issue in trade for a full year's subscription and see what they say!" We suggest you ransack your attic for a copy of *National Geographic*, vol. 1/no. 1 (1888) . . . it's currently worth about \$7000!

Best of luck with your very fine publishing venture.

Thomas Y. Canby
Science Editor
National Geographic Magazine
Washington, DC

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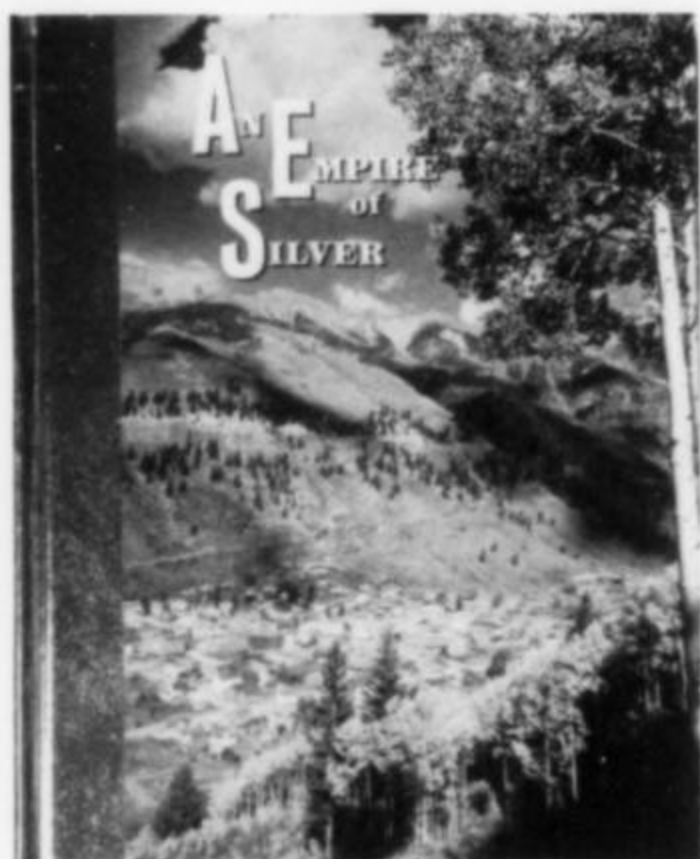
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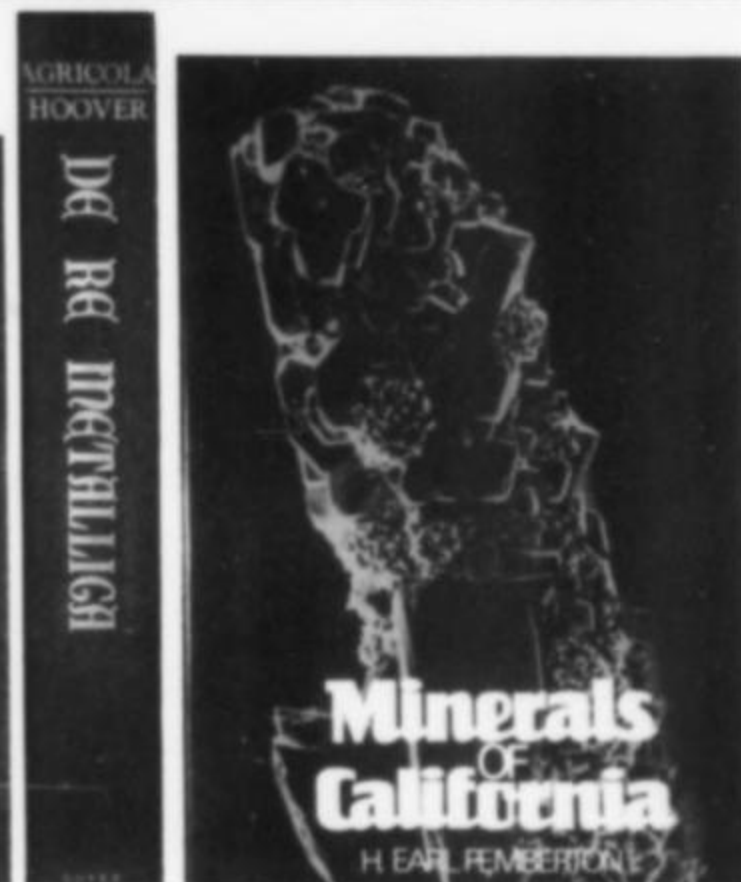
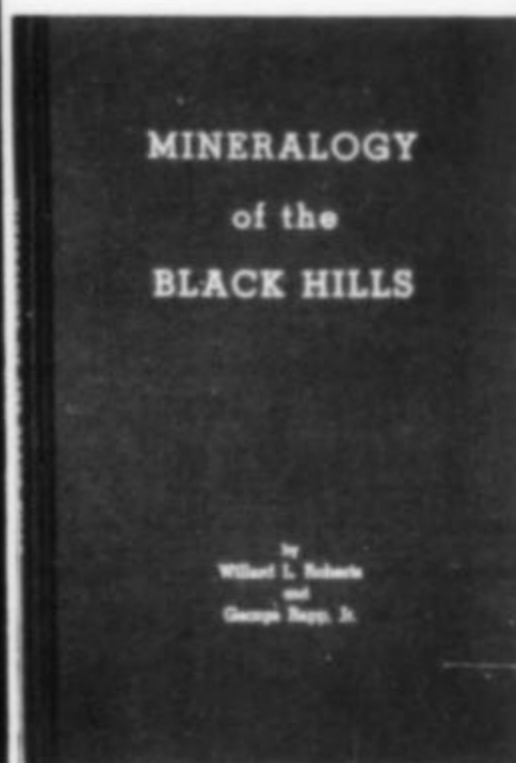
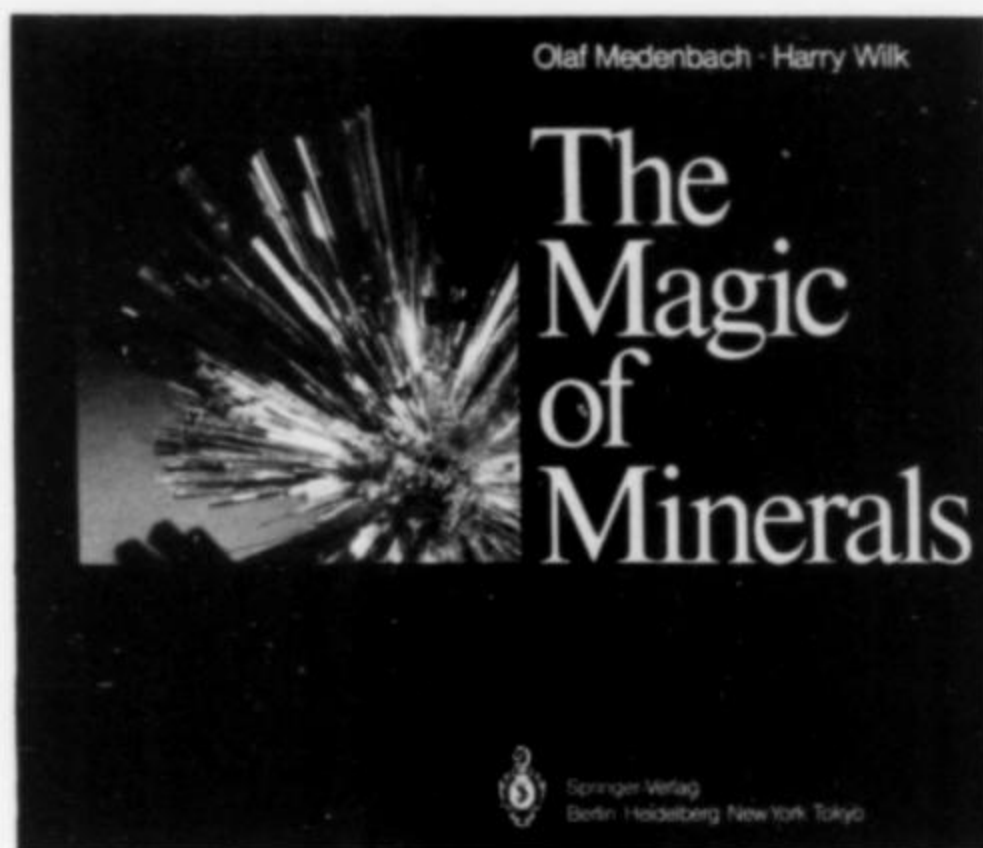
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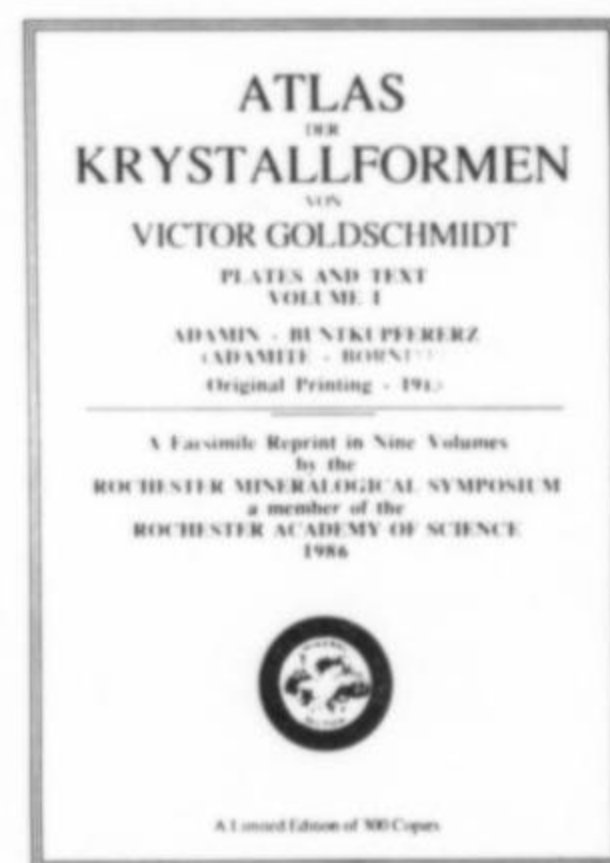
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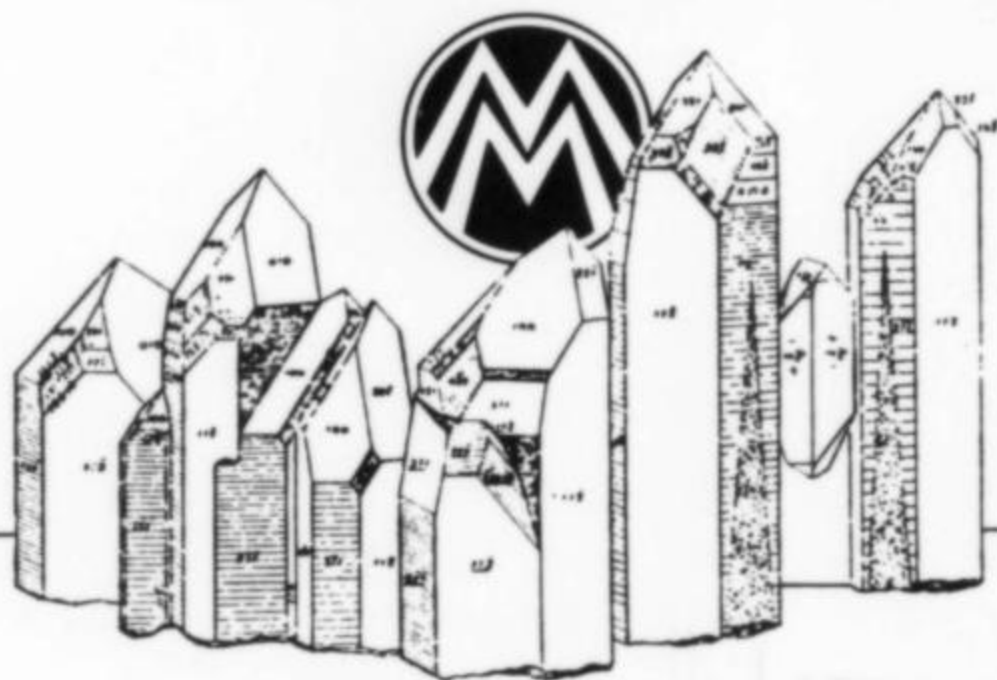


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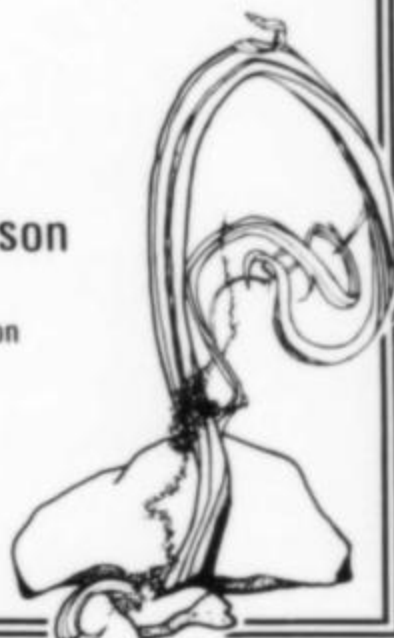
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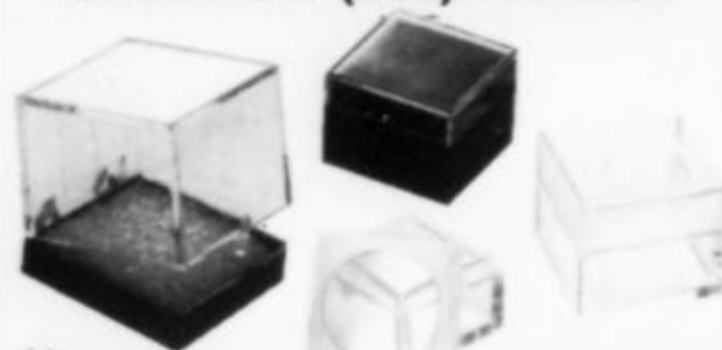
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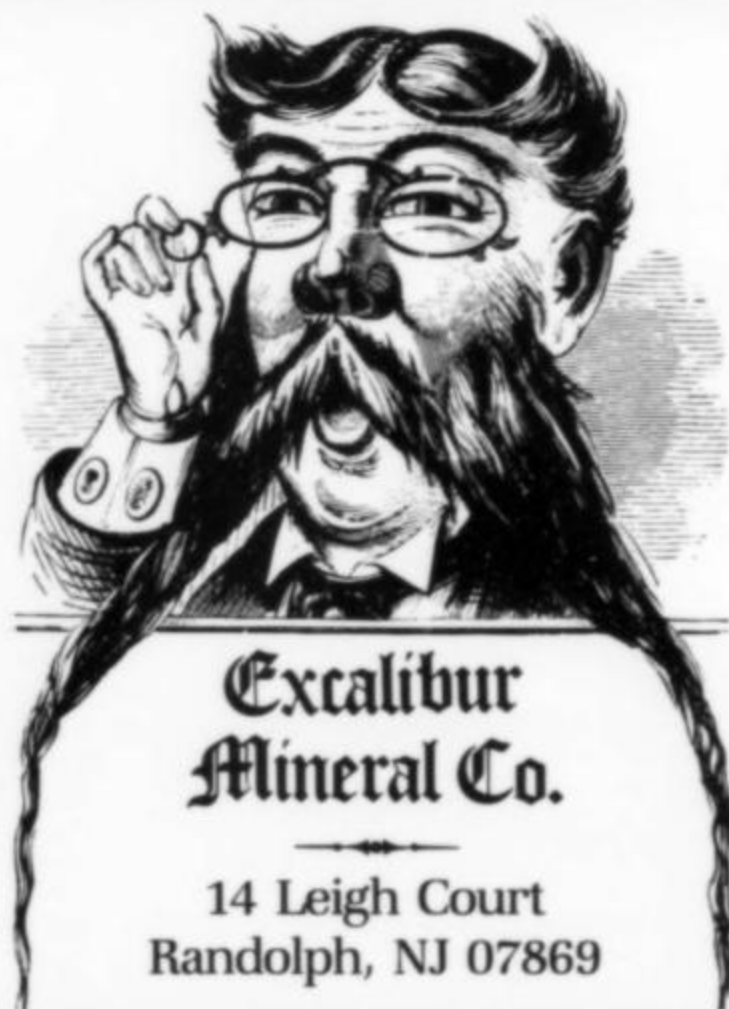
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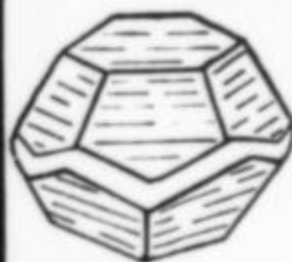
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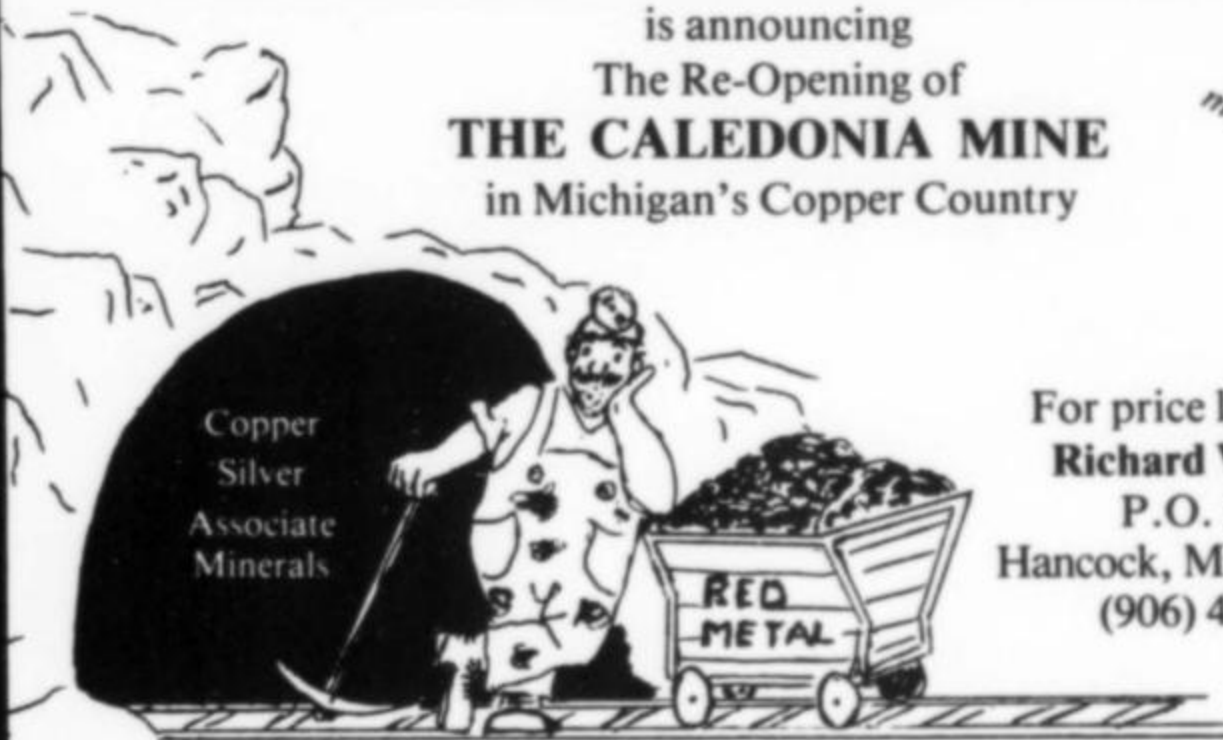
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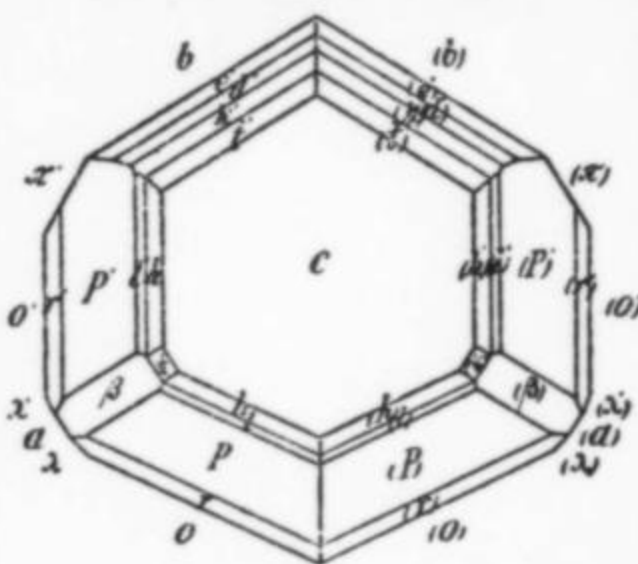
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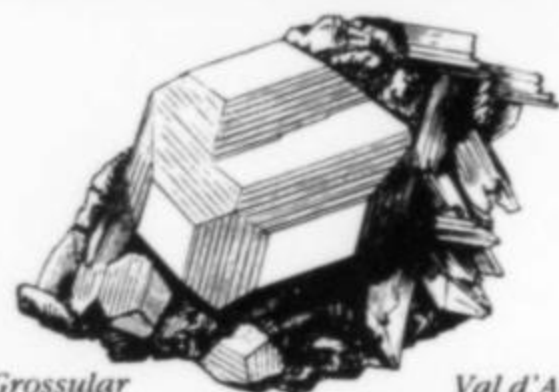
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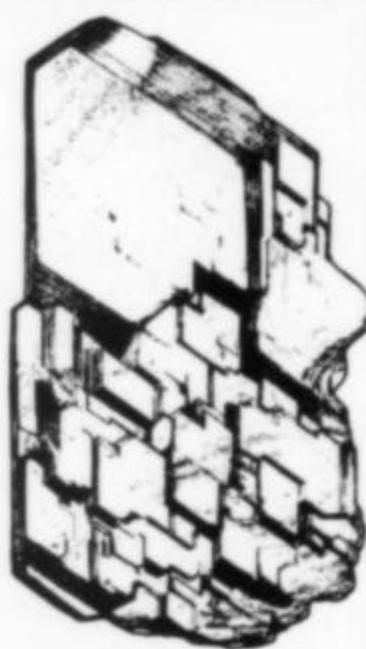


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Next issue: Fun on the way to the Munich Show.

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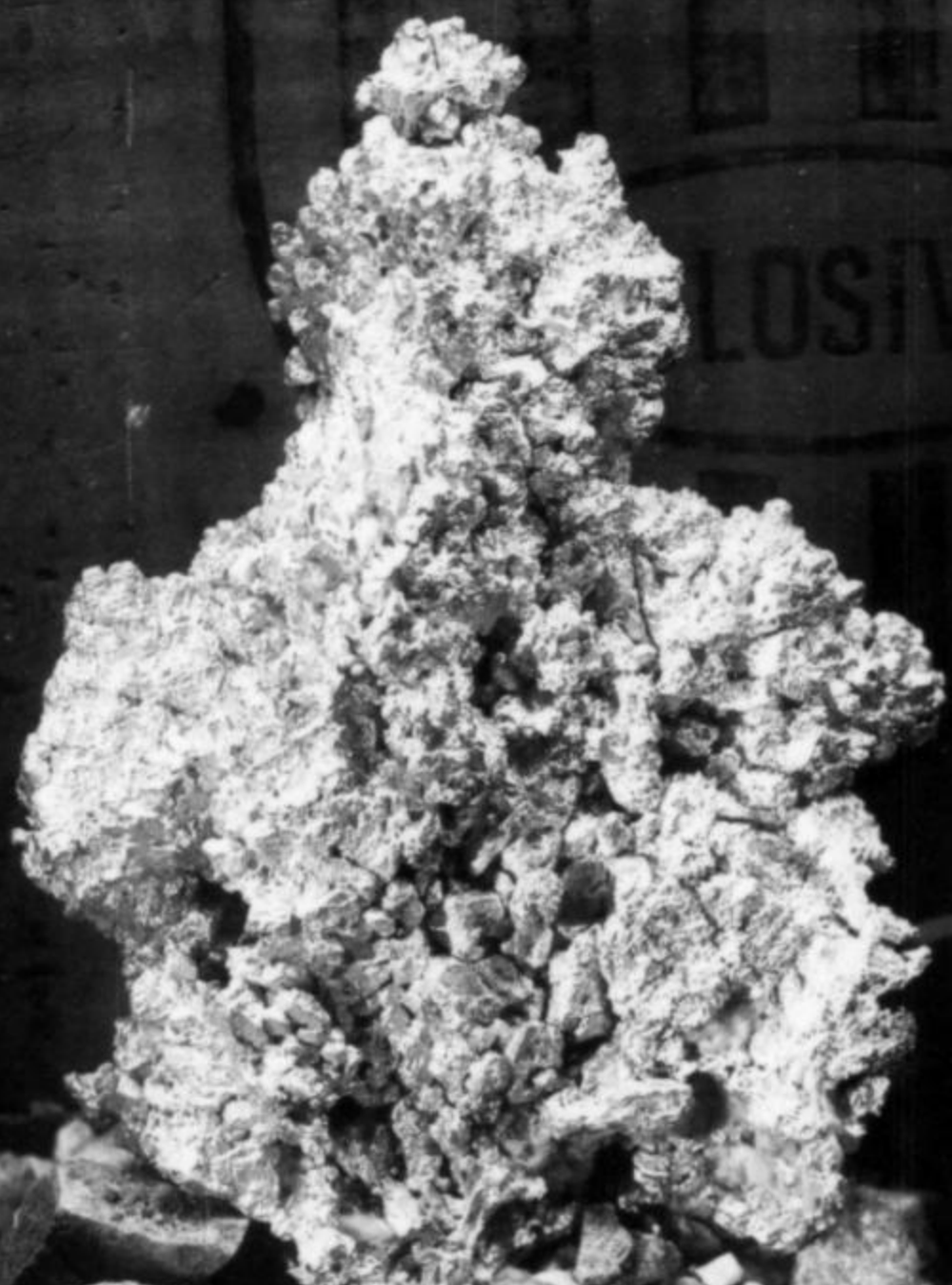


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